Development and application of the Model of Aerosol Dynamics, Reaction, Ionization, and Dissolution (MADRID)

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[1] A new aerosol model, the Model of Aerosol Dynamics, Reaction, Ionization, and Dissolution (MADRID) has been developed to simulate atmospheric particulate matter (PM). MADRID and the Carnegie-Mellon University (CMU) bulk aqueous-phase chemistry have been incorporated into the three-dimensional Models-3/Community Multiscale Air Quality model (CMAQ). The resulting model, CMAQ-MADRID, is applied to simulate the August 1987 episode in the Los Angeles basin. Model performance for ozone and PM is consistent with current performance standards. However, organic aerosol was underpredicted at most sites owing to underestimation of primary organic PM emissions and secondary organic aerosol (SOA) formation. Nitrate concentrations were also sometimes underpredicted, mainly owing to overpredictions in vertical mixing, underpredictions in relative humidity, and uncertainties in the emissions of primary pollutants. Including heterogeneous reactions changed hourly O3 by up to 17% and 24-hour average PM2.5, sulfate2.5, and nitrate2.5 concentrations by up to 3, 7, and 19%, respectively. A SOA module with a mechanistic representation provides results that are more consistent with observations than that with an empirical representation. The moving-center scheme for particle growth predicts more accurate size distributions than a typical semi-Lagrangian scheme, which causes an upstream numerical diffusion. A hybrid approach that simulates dynamic mass transfer for coarse PM but assumes equilibrium for fine PM can predict a realistic particle size distribution under most conditions, and the same applies under conditions with insignificant concentrations of reactive coarse particles to a bulk equilibrium approach that allocates transferred mass to different size sections based on condensational growth law. In contrast, a simple bulk equilibrium approach that allocates transferred mass based on a given distribution tends to cause a downstream numerical diffusion in the predicted particle size distribution. INDEX TERMS: 0305 Atmospheric Composition and Structure: Aerosols and particles (0345, 4801); 0345 Atmospheric Composition and Structure: Pollution—urban and regional (0305); 0365 Atmospheric Composition and Structure: Troposphere—composition and chemistry; 0368 Atmospheric Composition and Structure: Troposphere—constituent transport and chemistry; KEYWORDS: aerosols and particles, air quality, three-dimensional model


1. Introduction

[2] The demonstration of attainment of the National Ambient Air Quality Standards (NAAQS) (24-hour average and 3-year average concentrations) for particulate matter with aerodynamic diameter less than 2.5 μm (PM2.5) and progress under the Regional Haze rule in the United States will require the use of three-dimensional (3-D) air quality models to evaluate the effect of various emission management options on PM2.5 concentrations. These models will also be used in other areas due to the recent promulgation of a Canada-wide air quality standard for 24-hour average PM2.5 concentrations and the forthcoming consideration of PM2.5 standards by the European Union. Major episodic PM models include urban-scale models, mesoscale models and urban-through-global
models. Examples of the urban-scale models include the Urban Airshed Model with Aerosols (UAM-AERO) [Lurmann et al., 1997], the California/Carnegie-Mellon Institute of Technology (CIT) model [Meng et al., 1998] and the Urban Airshed Model with the Aerosol Inorganics Model (UAM-AIM) [Sun and Wexler, 1998a, Sun and Wexler, 1998b]. Examples of the mesoscale models include the Regional Particulate Model (RPM) [e.g., Binkowski and Shankar, 1995], the Denver Air Quality Model- Version 1 and 2 (DAQM and DAQM-V2) [Middleton, 1997; Regional Air Quality Council (RAQC), 1999], the Models-3 Community Multiscale Air Quality (CMAQ) Modeling System [Byun and Ching, 1999] and the San Joaquin Valley Air Quality Study (SIVAQ) and the Atmospheric Utility Signatures Predictions and Experiments Study (AUSPEX) Regional Modeling Adaptation Project (SARMAP) Air Quality Model with Aerosols (SAQM-AERO) [Dabdub et al., 1998; Pai et al., 2000]. An example of the urban-through-global models is GATOR [Jacobson, 1997a, 1997b]. These models have been applied to various airsheds including the Los Angeles Basin, CA (e.g., CIT, GATOR, SAQM-AERO, UAM-AERO, UAM-AIM and Models-3/CMAQ), Denver, CO (e.g., DAQM-V2) and the eastern North America (e.g., RPM and Models-3/CMAQ). Recent reviews of the current status of 3-D air quality models for PM [Seigneur et al., 1999; Seigneur, 2001; Seigneur and Moran, 2003] have suggested that existing 3-D models have several limitations in their treatment of aerosols that should be addressed before they can provide reliable results in a policy or regulatory context. For example, areas of improvements include the treatment of secondary PM formation (e.g., sulfate, nitrate and secondary organic aerosol (SOA)) and subgrid-scale plume treatment [Seigneur et al., 1999]. In this work, we present the development of a new model for the treatment of PM processes and its incorporation into a 3-D host model, the U.S. Environmental Protection Agency (EPA) Models-3/CMAQ. [5] CMAQ is a 3-D grid-based air quality model that can be applied to simulate ozone (O₃) and other photochemical oxidants, PM, and the deposition of pollutants such as acidic and toxic air pollutants. CMAQ was selected as the 3-D host air quality model following a review of several existing 3-D models by Seigneur et al. [2000a]. Its original formulation has been described by Byun and Ching [1999]. The version used here is the August 2000 version released by EPA. The new modules incorporated into CMAQ include the Model of Aerosol Dynamics, Reaction, Ionization, and Dissolution (MADRID) and the Carnegie-Mellon University (CMU) bulk aqueous-phase chemical kinetic mechanism (hereafter referred to as the CMU bulk aqueous-phase chemical mechanism). In addition, some existing modules of CMAQ were modified either to be compatible with the new modules or to improve the representation of atmospheric processes. These new or modified modules are incorporated into CMAQ as options. The resulting model, CMAQ-MADRID, is applied to simulate the 27–28 August 1987 episode of the Southern California Air Quality Study (SCAQ) in the Los Angeles Basin. Its performance is evaluated for O₃, PM₂.₅, particulate matter with aerodynamic diameter less than 10 μm (PM₁₀) and PM chemical components, and compared with previous performance evaluations conducted with other PM models. Sensitivity studies are conducted to evaluate the sensitivity and the sources of uncertainties in model predictions. Sensitivity simulations include those with and without heterogeneous reactions and those with different modules/algorithms for SOA formation, particle growth due to condensation (or shrinkage due to volatilization) and aqueous-phase chemistry as well as gas/particle mass transfer. CMAQ-MADRID was provided to EPA for public utilization in October 2002.

[4] We describe in section 2 the formulation of MADRID including the treatment of aerosol thermodynamics, dynamics and processes that govern the chemical composition and size distribution of PM in MADRID. These processes include new particle formation, condensational growth (or shrinkage by volatilization), and mass transfer between the bulk gas phase and PM. We describe in section 3 the treatment of cloud processes (e.g., particle scavenging, aqueous-phase chemistry, and particle formation after cloud evaporation), heterogeneous reactions taking place at the surface of particles or droplets, and dry/wet deposition of particles and condensible organic species. The application of CMAQ-MADRID along with results from base and sensitivity simulations is presented in section 4. Finally, the results are summarized along with discussions on further model improvements in section 5.

2. Formulation of MADRID

[5] MADRID is developed to simulate important microphysical processes that govern the chemical composition and size distribution of PM. We conducted comprehensive reviews of existing modules available to simulate PM thermodynamics and dynamics [Zhang et al., 1999, 2000; Seigneur, 2001]. On the basis of these reviews, a set of modules or algorithms that provides the best compromise between numerical accuracy and computational efficiency was selected to simulate those processes. The selected modules were then integrated and, if warranted, modified, to constitute a coherent framework for the simulation of atmospheric PM. The formulation of MADRID is described below according to the major PM processes.

2.1. Chemical Composition of PM

[6] PM consists of primary components that are emitted directly into the atmosphere and secondary components that are formed in the atmosphere by nucleation or condensation of gaseous species. We consider organic and inorganic PM of anthropogenic and biogenic origin. The chemical composition of PM is governed by the mass transfer between the bulk gas phase and the surface of the particles and the phase transition at the surface. The rate of mass transfer depends strongly on the composition at thermodynamic equilibrium between the bulk gas phase and the particles. The timescale for establishing thermodynamic equilibrium is sufficiently small for particles of small size; thermodynamic equilibrium between the gas phase and fine particles is therefore justified [e.g., Meng and Seinfeld, 1996; Fahey and Pandis, 2001]. The thermodynamic equilibrium cannot always be assumed for coarse particles (except for liquid water), however, and all other species are usually not in equilibrium. This complicates the aerosol modeling, as the dynamics of mass transfer have to be explicitly treated, at least for the larger particles with high concentrations. We describe below
the treatment of gas-particle thermodynamics. The treatment of mass transfer under nonequilibrium conditions is described in section 2.2.4.

2.1.1. Thermodynamics for Inorganic Species

[1] A comprehensive review of the existing algorithms available to simulate the gas/particle thermodynamic equilibrium of inorganic species was conducted by Zhang et al. [2000] and updated as candidate algorithms underwent further improvements. Six modules that simulate the gas/particle partitioning of inorganic species were compared using 400 different case studies. These six modules included MARS-A [Binkowski and Shankar, 1995], SEQUILIB [Pilinis and Seinfeld, 1987], SCAPE2 [Kim and Seinfeld, 1995; Meng et al., 1995], EQUISOLV II [Jacobson, 1999], AIM2 [Clegg et al., 1998a, 1998b] and ISORROPIA [Nenes et al., 1998, 1999]. All modules treat sulfate, nitrate, ammonium and water. Except for MARS-A, all modules also treat sodium and chloride. In addition, SCAPE2 and EQUISOLV II treat crustal soluble species: calcium, magnesium, potassium and carbonate. AIM2 does not simulate alkaline systems and was therefore not considered for incorporation into MADRID. MARS-A was the default module of CMAQ but it does not treat sodium chloride (NaCl). Among the computationally efficient modules that treat NaCl, ISORROPIA was judged superior to SEQUILIB in terms of numerical accuracy and stability. For a comprehensive treatment of the aerosol system, both SCAPE2 and EQUISOLV II were considered suitable. SCAPE2 is more computationally demanding than ISORROPIA (e.g., by factors of 10 to 440 under the conditions tested by Nenes et al. [1998]). EQUISOLV II can be applied to solve aerosol thermodynamic equilibrium at one time for either a single grid cell or multiple grid cells with a vectorized approach that can be applied to scalar or vector machines. We have tested the computational time of EQUISOLV II and ISORROPIA. For the tests, both modules consider the same total number of equilibrium equations but with different sets of equilibria, and ISORROPIA solved fewer equations in some concentration/relative humidity (RH) regimes. Under these test conditions, ISORROPIA required less computational time in a single grid cell than did EQUISOLV II for the same convergence criteria. As the number of grid cells increased, EQUISOLV II became more computationally efficient than ISORROPIA because the computer time required per grid cell decreases on a scalar machine due to an array-referencing minimization technique, and additional speedup occurs with multiple cells on a vector machine due to vectorization [Zhang et al., 2000]. However, it is not a trivial effort to implement the vectorized version of EQUISOLV II into CMAQ for solving aerosol thermodynamics at one time for multiple grid cells. Moreover, PM emission inventories do not yet include the chemical composition of crustal species that are treated in EQUISOLV II. Therefore ISORROPIA was selected to simulate the thermodynamics of inorganic PM species in MADRID. The latest version of ISORROPIA (i.e., v1.6) is currently used in MADRID.

[8] The thermodynamic equilibria that are simulated by ISORROPIA are presented by Nenes et al. [1998]. The entire concentration domain is divided into subdomains such as sulfate very rich (free acid), sulfate rich (non free acid), sulfate poor and sodium poor, and sulfate poor and sodium rich. The systems of nonlinear equations for each subdomain were ordered and manipulated so that analytical solutions can be obtained for as many equations as possible. Adopting this approach, most cases can be solved using only one level of iteration, which increases computational efficiency considerably. Significant speedup is also obtained by minimizing the number of calls to subroutines that calculate activity coefficients. In addition, some speedup can be gained under some conditions by reducing the number of equations solved for a given RH/concentration regime and/or using precalculated activity coefficient tables. The bisection method is used to obtain the solution. ISORROPIA provides options to treat particles to be either in a thermodynamically stable state, where particles can be solid, liquid or both, or in a metastable state, where particles are always an aqueous solution. The first option (i.e., all states are treated) is used in MADRID.

2.1.2. Thermodynamics for Organic Species

[9] Accurate simulation of SOA formation requires a gas-phase mechanism that treats all important semivolatile or nonvolatile organic species and their oxidation products. Such a detailed representation of organic species and their chemistry, however, usually cannot be incorporated in 3-D air quality model, due mainly to computational constraints and current knowledge gaps in the atmospheric chemistry of organic compounds. Different VOC-lumping approaches are therefore used in existing gas-phase chemical mechanisms to provide a simplified treatment for VOC speciation and chemistry. These include the lumped structure approach (e.g., the Carbon-Bond Mechanism Version IV (CBM-IV) of Gery et al. [1989]); the lumped species approach (e.g., the Regional Acid Deposition Mechanism Version 2 (RADM2) of Stockwell et al. [1990] and the Statewide Air Pollution Research Center gas-phase mechanism (SAPRC-99) of Carter [2000]); and the lumped surrogate species (e.g., the Caltech Atmospheric Chemistry Mechanism (C ACM) of Griffin et al. [2002]). C ACM contains 361 reactions of 191 species and provides detailed descriptions of several generations of products from alkanes (3 classes), alkenes (2 classes), aromatics (2 classes), alcohols (3 classes), isoprenes and terpenes (2 classes) [Griffin et al., 2002]. This mechanism is uniquely suitable for simulating SOA formation because it explicitly treats 42 condensable second- and third-generation products. Although many of the reactions for organic species are generic and some organic species are nonexistent, SAPRC-99 includes the product yield coefficients and rate constants for over 100 individual organic species [Carter, 2000]. It is therefore suitable for simulating SOA formation. Its predecessors (i.e., SAPRC-90, SAPRC-93 and SAPRC-97) have been applied for simulating SOA formation in several airsheds [e.g., Pandis et al., 1992, 1993; Bowman et al., 1995]. SAPRC-99 is available in Models-3/CMAQ since June 2002. On the other hand, both CBM-IV and RADM2 are computationally more efficient than both C ACM and SAPRC-99. They contain only a few explicit and lumped organic species; thus they may not be well suited for a detailed representation of SOA formation. For example, since individual organic compounds are lumped based on carbon-bond structure in CBM-IV, they are often disaggregated and assigned to more than one mechanism species. Many of the organic mechanism species thus contain frag-
ments of molecules and the identity of the original organic compound is lost. It is therefore not possible to track the amount of some SOA precursors (e.g., long-chain alkanes and alkenes) reacting in CBM-IV.

[10] CMAQ-MADRID includes two SOA formulations: one uses an empirical representation of SOA formation (referred to as MADRID 1) that is based on data obtained in smog chamber experiments [Odum et al., 1997; Griffin et al., 1999]; the other uses a mechanistic representation of SOA formation (referred to as MADRID 2) that simulates an external mixture of hydrophilic and hydrophobic particles [Pun et al., 2002]. Different gas-phase chemical mechanisms are used in MADRID 1 and MADRID 2 to accommodate the different SOA speciations simulated. MADRID 2 uses CACM gas-phase chemistry, MADRID 1 uses either CBM-IV or RADM2, which are the only two gas-phase chemical mechanisms available in the August 2000 version of Models-3/CMAQ. MADRID 2 includes 10 surrogate compounds, grouped according to their affinity for water (5 surrogate species for 28 explicit hydrophilic compounds and 5 surrogate species for 14 hydrophilic compounds), origin (anthropogenic versus biogenic), size (number of carbons), volatility, and dissociation properties [Pun et al., 2002]. Hydrophilic organic compounds include those with a short carbon chain (≤7 carbons; or ≤10 carbons with three or more functional groups), high solubility (≥1 g solute/100 g water), and a high effective Henry’s law constant (≥1 × 10^4 M atm^−1). Hydrophobic compounds are identified by their estimated octanol-water partition coefficients. A detailed description of MADRID 2 has been presented elsewhere [Pun et al., 2002] and is not repeated here. For MADRID 1, the formulation of the SOA module requires some additional precursors and condensable organic products that are not explicitly treated in CBM-IV or RADM2 (a detailed list of SOA precursors and products used in CBM-IV and RADM2 can be found in Byun and Ching [1999]). Therefore it is necessary to add several organic species and reactions to these two mechanisms to make them compatible with the SOA formulation of MADRID 1. Zhang et al. [2002a] and Pun et al. [2003] have presented the chemical speciation used for the formulation of MADRID 1. We present below the SOA formulation in MADRID 1, along with modifications to the original CBM-IV and RADM2 gas-phase chemical mechanisms.

[11] The MADRID 1 formulation for SOA includes 2 anthropogenic VOC precursors, 4 surrogate anthropogenic species representing their condensable products, 12 biogenic VOC (BVOC) precursors and 34 surrogate biogenic species representing their condensable products (22, 8 and 4 surrogate species resulted from the OH, O3 and NO3 oxidation reactions, respectively). The two anthropogenic precursors are assumed to be aromatics and are characterized as one with low SOA yield and one with high SOA yield. The high-yield aromatic species include those containing no more than one methyl substituent and no more than one ethyl substituent (i.e., toluene, ethylbenzene and ethyltoluenes) as well as n-propylbenzene. The low-yield aromatic species include those containing two or more methyl substituents (i.e., xylene, trimethylbenzenes, dimethylbenzenes and tetramethylbenzenes). The existing CBM-IV and RADM2 lumped aromatic species toluene (TOL) and xylene (XYL) represent aromatics with 7- and 8-carbon ring structures in CBM-IV or less and more reactive aromatics in RADM2, respectively. These two lumped species (i.e., TOL and XYL) were selected to represent the high-yield and low-yield anthropogenic precursors, respectively. Other anthropogenic SOA precursors, such as long-chain (>C8) alkanes, long-chain internal alkenes and cresol and phenols, are not considered in MADRID 1 (they are treated in MADRID 2). Four anthropogenic condensable products were added to the existing reactions of TOL and XYL: TOLAE1 and TOLAE2 (high aerosol yield products) for TOL oxidation by OH and XYLAE1 and XYLAE2 (low aerosol yield products) for XYL oxidation by OH, using the experimentally determined stoichiometric coefficients of Odum et al. [1997]. The 12 biogenic precursors in MADRID 1 do not appear explicitly in CBM-IV and RADM2 because they are decomposed into their functional groups (i.e., ALD2, OLE and PAR) in CBM-IV or assigned to surrogate species in RADM2. Therefore we added those species and their corresponding oxidation reactions which lead to biogenic SOA (BSOA) formation using the experimentally determined stoichiometric coefficients of Griffin et al. [1999] and the kinetic rate constants compiled in the work of Lamb et al. [1999].

[12] Since BVOC are already represented in the original lumped structure (or species) formulation of CBM-IV (or RADM2) to simulate O3 formation, it is important that the reactions added for BSOA formation do not alter O3 chemistry. Accordingly, we added oxidation reactions of BVOC in a format such that an oxidant (e.g., OH, O3 and NO3) is treated as both a reactant and a product following the approach of Gipson and Young [1999], namely:

\[
\text{BVOC} + \text{OH} \rightarrow b \text{BSOA} + \text{OH}
\]

where \( b \) is the stoichiometric coefficient for the biogenic product BSOA. In this way, the oxidant mixing ratio and thus O3 chemistry are unaffected by this new reaction, which affects only the precursor and the condensable organic product.

[13] The following equation governs the gas/particle partitioning of each of the condensable species in MADRID:

\[
K_i = \left( \frac{A_i}{M_{\text{sum}} G_i} \right),
\]

where \( K_i \left( \text{m}^3/\mu \text{g}^{-1} \right) \) is the partition coefficient obtained from the smog chamber experiments, \( A_i \) and \( G_i \) (\( \mu \text{g} \text{m}^{-3} \text{air} \)) are the mass concentrations of species \( i \) in the particulate- and gas-phase, respectively, and \( M_{\text{sum}} \left( \mu \text{g} \text{m}^{-3} \text{air} \right) \) is the sum of primary organic carbon (OC) (nonvolatile) and secondary OC (semivolatile) in the particulate phase that serve as the organic absorbing medium. The SOA yields and gas/particle partition coefficients at experimental temperatures are based on Odum et al. [1997] and Griffin et al. [1999]. Griffin et al. [1999] conducted outdoor chamber experiments on aerosol formation under both daytime conditions in the presence of OH, O3 and NO3 and nighttime conditions with either O3 or NO3. Since the OH radical is the primary oxidant that oxidizes the largest fractions of BVOC during the daytime [Griffin et al., 1999], we assume that the 22 surrogates from
the daytime photooxidation experiments of Griffin et al. [1999] resulted from the OH oxidation. The daytime SOA yields and gas/particle partition coefficients from Griffin et al. [1999] are thus used for these OH oxidation reactions in MADRID 1.

[15] The smog chamber experiments from which $K_i$ and stoichiometric coefficients were derived were conducted at temperatures higher (301–316 K) than the typical ambient temperatures. Following Sheehan and Bowman [2001], the temperature dependence of $K_i$ can be accounted for as follows:

$$K_i(T) = K_i^* \frac{T}{T^*} \exp \left[ \frac{\Delta H_{\text{vap},i}}{R} \left( \frac{1}{T} - \frac{1}{T^*} \right) \right].$$

where $K_i(T)$ and $K_i^*$ are the partition coefficients at temperature $T$ and a reference temperature $T^*$, respectively. $R$ is the ideal gas constant $(8.2 \times 10^{-5} \text{ m}^3 \text{ atm mol}^{-1} \text{ K}^{-1})$. $\Delta H_{\text{vap},i}$ is the enthalpy of vaporization of the pure species $i$. The value of $\Delta H_{\text{vap},i}$ is assumed to be 88 kJ mol$^{-1}$ for condensable products from terpenes and aromatics (<C10) and 175 kJ mol$^{-1}$ for those from sesquiterpenes. The values for $\Delta H_{\text{vap},i}$ correspond to the arithmetic mean of available literature data for <C10 and >C10 compounds [Tao and McMurtry, 1989; Bilde and Pandis, 2001].

### 2.2. Aerosol Dynamics

#### 2.2.1. Particle Size Distribution

[15] Two major approaches have been commonly used to represent the particle size distribution: the modal and the sectional approaches. In the modal approach, the particle size distribution is represented by several modes (e.g., Aitken, accumulation and coarse modes) and an analytical function (typically, a lognormal distribution) is used to represent the particle size distribution of each mode. The aerosol dynamic processes that govern the evolution of the particle size distribution can then be solved analytically. In the sectional approach, the particle size distribution is approximated by a discrete number of size sections. Some properties of the particle size distribution (e.g., mass of individual chemical species) are then assumed to be uniform within each size section and to be conserved as the aerosol general dynamic equation is solved. The modal approach offers computational advantages over the sectional approach, but has inherent limitations in representing a wide variation of the observed aerosol size distributions and their physical and chemical processes [e.g., Zhang et al., 2002b]. The modal approach is used in the original CMAQ. The sectional approach is used in MADRID to represent the particle size distribution. The processes that govern aerosol dynamics include coagulation, nucleation (i.e., the formation of new particles), growth due to condensation (or shrinkage due to volatilization) and the mass transfer of chemical species between the bulk gas phase and the particle surface. All processes except for coagulation are treated explicitly in MADRID. Coagulation is not included in the current version of MADRID, although it may have a large effect on fine particle number concentrations, particularly near emission sources, and it may affect the mass and number concentrations of fine particle via internally mixing particles over the entire size distribution [Jacobson, 2002]. For this particular study with a simulation period of a few days, the effect of neglecting coagulation on the overall PM mass predictions may be negligible, because the timescale for coagulation is long compared to that of other processes such as condensation. We do not attempt to evaluate the predicted PM number concentrations in this study, because coagulation may have a significant impact on fine particle number concentrations, the predicted PM number concentrations for ultra-fine particles (i.e., those with a diameter less than 0.1 μm) may not be accurate.

[16] Either two or multiple particle size sections can be used to represent the particle size distribution in MADRID. For the 2-section representation (i.e., fine and coarse), particle growth by condensation and shrinkage by volatilization are not simulated because there is minimal exchange via growth/shrinkage between the fine and coarse particle sections. For a multisectional representation, a minimum number of 8 sections is recommended to provide sufficient resolution of particle sizes for the meaningful simulation of aerosol dynamic processes. New particle formation, growth by condensation, shrinkage by volatilization and gas/particle mass transfer are simulated. We describe below the formulations used in MADRID to simulate these processes.

[17] The general dynamic equation for the multisectional representation of PM can be expressed as follows.

$$\frac{\partial q_i(d_{p,j},x,t)}{\partial t} = -\mathbf{u}(x,t) \cdot \nabla q_i(d_{p,j},x,t) + \nabla \cdot \mathbf{K}(x,t) \nabla q_i(d_{p,j},x,t) + V_d(x,t) \nabla q_i(d_{p,j},x,t) + H_i(d_{p,j},x,t) q_i(d_{p,j},x,t) - \frac{\partial (mq_i d_{p,j},x,t)}{\partial m} H_i d_{p,j},x,t + E_i(d_{p,j},x,t) + F_i(d_{p,j},x,t)$$

where $q_i$ is the mass of species $i$ in size section $j$ with a characteristic diameter $d_{p,j}$, $x$ is the position vector of the corresponding grid cell, $t$ is the time, $\mathbf{u}$ is the resolved wind vector, $\mathbf{K}$ is the eddy-diffusivity tensor, $V_d$ is the vertical deposition velocity that includes gravitational settling, $H_i$ is the condensational growth rate, $m$ is the mass of particles with a diameter $d_{p,j}$, $E_i$ is the emission rate of particulate species $i$ in size section $j$ and $F_i$ is the rate of new particle mass formation for species $i$ in the lowest size section.

[18] The first two terms on the right-hand side are solved using the host model transport algorithms [Byun and Ching, 1999]. The other terms are discussed below. Note that MADRID uses the Stokes particle diameter in its formulation whereas the PM$_{2.5}$ and PM$_{10}$ regulations use the aerodynamic diameter. These two diameters are related by the square root of the particle density. For example, for a particle density of 1.35 g cm$^{-3}$, the aerodynamic diameter of 2.5 μm corresponds to a Stokes diameter of 2.15 μm.

#### 2.2.2. Formation of New Particles

[19] Four algorithms [Pandis et al., 1994; Wexler et al., 1994; Fitzgerald et al., 1998; Harrington and Kreidenweis, 1998] currently used in 3-D air quality models to calculate the absolute rate of particle nucleation were compared by Zhang et al. [1999]. Although all algorithms were formulated from the same theoretical basis, they gave highly
variable results under typical conditions. Therefore the use of these parameterizations of the absolute nucleation rates is associated with significant uncertainties at present. Consequently, Zhang et al. [1999] recommended a method that calculates the relative rates of new particle formation and condensation on existing particles [McMurry and Friedlander, 1979] instead of calculating the absolute rate of nucleation. This method was selected for the treatment of new sulfate particle formation in MADRID.

[20] The new particle parameterization of McMurry and Friedlander [1979] is computationally demanding because it involves iteration among several equations. Consequently, a parameterized version that uses a look-up table with precalculated rates of new particle formation is used. An option for neglecting the calculation of new particle mass formation is also provided.

2.2.3. Condensational Growth

[21] Condensational growth (or shrinkage due to volatilization) is a process that allows particles to grow (or shrink) upon condensation of the condensable species (or evaporation of the volatile species). The simulation of condensation/volatilization is challenging with a sectional representation because numerical diffusion may result from the solution of the governing equation in 3-D simulations. Three basic approaches have been used to simulate condensational growth:

[22] 1. Semi-Lagrangian techniques that calculate the mass (or number) flux from one section to the next. The basic finite difference method [e.g., Seigneur, 1982] is the simplest example of a semi-Lagrangian technique. Bott’s scheme [Bott, 1989], the scheme of Chock and Winkler [2000] and that of Nguyen and Dabdub [2002] are more advanced examples of semi-Lagrangian techniques.

[23] 2. Lagrangian techniques that calculate the movement of the section boundaries according to the growth law and redistribute the resulting sectional distribution onto the fixed sectional representation. The UAM-AERO scheme [Lurmann et al., 1997] is an example of a Lagrangian technique where a spline function is used for the redistribution of the sectional representation.

[24] 3. The moving-center technique of Jacobson [Jacobson, 1997a] where the diameter representative of the section moves according to the growth law. It contains features of both Eulerian and Lagrangian schemes, since it uses fixed boundaries and allows movement within and across the boundaries.

[25] Zhang et al. [1999] compared four condensational growth algorithms: one modal approach that is used in Models-3/CMAQ [Binkowski and Shankar, 1998] and three sectional approaches: the Bott’s scheme used in the 1998 version of the CIT model [Meng et al., 1998], the UAM-AERO scheme and the moving-center scheme. The Bott’s and the UAM-AERO schemes were shown to lead to significant numerical diffusion and the moving-center scheme appeared to be the most accurate among the algorithms tested. Consequently, the moving-center scheme is used to simulate condensational growth in MADRID with more than two particle size sections. The sensitivity of the predicted total particle mass and size distributions to the moving-center scheme and a simple finite difference scheme are evaluated in the sensitivity study.

[26] The following growth law is used in MADRID to simulate the flux between the gas phase and particles [Capaldo et al., 2000],

\[
J_{ij} = 2\pi dp_j D_{gi} N_j (C_{\infty,i} - C_{uij}) \frac{1 + K_{uij}}{2 K_{uij} (1 + K_{uij})} \]

where \(J_{ij}\) is the growth/evaporation rate of species \(i\) in a particle in size section \(j\), \(N_j\) is the number density of particles in section \(j\), \(D_{gi}\) is the molecular diffusivity of species \(i\) in air, \(C_{\infty,i}\) and \(C_{uij}\) are the concentrations of species \(i\) in the bulk gas phase and at the surface of particles in section \(j\), respectively. \(K_{uij}\) is the Knudsen number, \(K_{uij} = 2 \lambda_i / dp_j\), \(\lambda_i\) is the mean free path of species \(i\), and \(\alpha_{uij}\) is the accommodation coefficient for species \(i\) on the particle in section \(j\), a value of \(\alpha_{uij} = 0.1\) is assumed for all species and all size sections.

[27] Each section center (i.e., \(dp_j\)) moves according to the change in mass in the section. As a section center reaches one of the boundaries of the section (upper boundary in the case of condensation; lower boundary in the case of volatilization), the particulate mass contained in that section moves into the adjacent section. This technique minimizes numerical diffusion across size sections since particulate mass is transferred from one section to the next only in the case where a section center reaches one of the section boundaries. It is important to note that this technique allows the simultaneous tracking of PM mass and number concentrations. In the most common implementation of the sectional approach in 3-D models, only PM mass is tracked. PM number is generally not conserved, since the PM number concentrations are diagnosed from the predicted PM mass and the fixed PM mean diameters (i.e., the so-called single-moment algorithm).

[28] In a 3-D air quality model, PM populations that are mixed within a given grid cell are likely to originally have different section centers. We used a mixing approach that is similar to that of Jacobson [1997a] for mixing PM populations within a given grid cell via advection, turbulent diffusion, convection, emissions and sedimentation. In this approach, an initial particle size distribution is assumed for newly emitted particles to calculate the number concentrations of the emitted particles based on the emitted mass concentrations. The emitted particles are then placed in the section surrounding their diameter. Since the changes in both mass and number concentrations of particles due to various atmospheric processes (e.g., emission, advection, turbulent diffusion and nucleation) are explicitly treated in CMAQ-MADRID, the new common section center for the mixed particles in a given size section can then be calculated using the particle mass and number concentrations \((m_j\) and \(n_j\) respectively) in the same size section:

\[
d_{p,ij} = \sqrt[3]{\frac{6m_j}{\pi \cdot n_j \cdot \rho_p}} \]

where \(\rho_p\) is the density of the particle.

2.2.4. Gas/Particle Mass Transfer

[29] Gas/particle mass transfer is a process that transfers mass of condensable species from bulk gas phase to the
particle surface. The timescale for the diffusion of a molecule from the bulk gas phase to the surface of a particle increases with the diameter of the particle. Therefore fine particles will tend to reach equilibrium rapidly whereas coarse particles can remain in nonequilibrium conditions [e.g., Wexler and Seinfeld, 1990; Dassios and Pandis, 1999]. Three basic approaches have been developed to treat the gas/particle mass transfer:

1. Dynamic approach that explicitly simulates gas/particle mass transfer for each size section by solving the equation for mass fluxes between the bulk gas-phase and individual particles (or particles in a given size range). Chemical concentrations in the bulk gas phase and in the particles in a given size section may or may not be in equilibrium. Examples of the dynamic approaches for multicomponent aerosols include those of Meng et al. [1996], Meng et al. [1998], Jacobson [1997a, 1997b], Sun and Wexler [1998a, 1998b] and Pilinis et al. [2000].

2. Equilibrium approach that assumes an instantaneous chemical equilibrium between the bulk gas phase and the whole particulate phase. This approach can be further divided into two categories: bulk equilibrium approach and nonbulk equilibrium approach. In the bulk equilibrium approach, all particles over size sections are assumed to have the same chemical composition. In the nonbulk equilibrium approach (also referred to as the size-resolved equilibrium approach; see Moya et al. [2002]), particles in different size sections may have different chemical compositions. The bulk equilibrium approach of Binkowski and Shankar [1995] and the simple bulk equilibrium approach of Hudischewskyj and Seigneur [1989] and Seigneur et al. [1997] (the latter approach was used in the sensitivity study in this work and is referred to as a simple bulk equilibrium approach hereafter) are examples of simple bulk equilibrium approaches, in which the transferred material is allocated to the particle size distribution using weighting factors that are derived based on either initial particle mass/surface area or a given distribution. In more advanced bulk equilibrium approaches such as those used in UAM-AERO [Lurmann et al., 1997] and CIT [Meng et al., 1998] (the latter approach is referred to as the CIT bulk equilibrium hereafter), the weighting factors are calculated based on condensational growth law using diffusion-limited assumptions, accounting more or less for the nonequilibrium between the bulk gas phase and particles in a given size range. Examples of nonbulk equilibrium approach include those of Pilinis and Seinfeld [1987]; Kleeman et al. [1997]; Jacobson et al. [1996a]; Jacobson [1999] and Moya et al. [2002], in which a system of nonlinear algebraic equations is solved for each size range to determine the partitioning of semivolatile species, while the mass transfer between the bulk gas-phase and bulk particulate phase is still considered to occur at an instantaneous thermodynamic equilibrium.

3. Hybrid approach that combines both dynamic and equilibrium approaches. The CMU hybrid approach of Capaldo et al. [2000] is an example of such a hybrid approach, in which the mass transfer is treated explicitly for the coarse particles and the gas/particle equilibrium is assumed for the fine particles.

The dynamic approach provides the most accurate representation of interphase partitioning of semivolatile species in theory but its use in 3-D air quality models is limited by its large computational expenses. In the dynamic approaches that are used in current 3-D models (e.g., CIT, GATOR and UAM-AIM), particles are usually assumed to be internally mixed (i.e., all particles within a given size range have the same chemical composition) and are distributed according to size sections. Therefore the mass transfer equation is solved between the bulk gas phase and the surface of the particles. On the other hand, the equilibrium approach is computationally efficient and has been used extensively in many 3-D models. Zhang et al. [1999] compared the CIT bulk equilibrium approach and the simple bulk equilibrium approach with the CIT dynamic approach in a box model. They found that the simple bulk equilibrium approach is inaccurate under many ambient conditions, whereas the CIT bulk equilibrium approach is appropriate when chloride and carbonate concentrations are insignificant. While the bulk equilibrium approach introduces errors in the partitioning calculation, particularly for cases with highly reactive coarse particles, the nonbulk equilibrium approach provides a more accurate representation of the interphase partitioning. However, the nonbulk equilibrium approach may lead to infinite solutions for solids and the equilibrium assumption is usually not valid for coarse particles. The novel hybrid approach combines merits of both dynamic and equilibrium approaches; therefore it provides the best compromise between numerical accuracy and computational speed.

Accordingly, we selected the hybrid approach of Capaldo et al. [2000] (referred to as the CMU hybrid approach) to treat gas/particle mass transfer in MADRID and modified it as discussed below.

Capaldo et al. [2000] recommended the use of a threshold size of 1 μm, which is a cut-off size between the equilibrium approach and the dynamic approach. We tested a threshold of 2.15 μm under typical urban aerosol conditions in a box model to estimate the sensitivity of model results to the selection of the cut-off size. The maximum difference in predicted PM concentrations in a given size section with the two thresholds under the tested conditions was only 3%, although errors in predicted size distribution may be higher under other conditions or small errors may propagate into larger errors over long simulation periods in a 3-D model. Nevertheless, a threshold value of a Stokes diameter of 2.15 μm represents a suitable compromise between computational efficiency and accuracy and is therefore used when 2 or 8 size sections are selected in MADRID. For a different section number and/or distribution, a cut-off size between 1 and 2.15 μm should be selected by the user to be the threshold value.

For particles in size sections above the threshold diameter, the mass transfer flux equation is solved according to the growth law of equation (4). The chemical concentration $C_{ij}$ is calculated by solving the gas/particle thermodynamic equilibrium knowing the particulate chemical concentrations. This calculation is conducted using ISORROPIA for wet particles and the Multicomponent Aerosol Dynamic Model (MADM) for dry particles [Pilinis et al., 2000; Capaldo et al., 2000]. If the bulk gas-phase chemical concentration exceeds the concentration at the surface of the particle, the mass transfer occurs from the bulk gas phase toward the particle, and vice-versa. The mass
transfer and thermodynamic equilibrium equations are solved iteratively until convergence is attained.

\[ C_{a,j} = C_{\infty,j} \]  

Thus ISORROPIA is used to calculate the bulk particulate concentrations given the bulk gas-phase concentrations. The material transferred between phases is distributed over the fine particle size sections by using weighting factors that are based on the surface area of particles in each section [Capaldo et al., 2000].

[37] As chemical species either condense onto or volatilize from the particles, these particles grow or shrink accordingly. Therefore in the case where more than 2 sections are selected, the hybrid mass transfer algorithm must be coupled with a growth/shrinkage algorithm. The original formulation of the mass transfer algorithm in the CMU hybrid approach used a finite difference scheme to treat growth/shrinkage of the particles. As discussed above, such schemes can lead to substantial numerical diffusion. Therefore the finite difference scheme was replaced by the moving-center scheme described above.

[38] CMAQ-MADRID offers two other options for gas-to-particle mass transfer: a simple bulk equilibrium approach in which equation (6) applies to all size sections and the weighting factors are calculated based on sulfate concentrations; and the CIT bulk equilibrium approach of Meng et al. [1998]. The CIT bulk equilibrium approach distributes the particle mass changes over the size sections based on a particle surface area weighting that is similar to that of the CMU hybrid approach. These two options are computationally more efficient than the CMU hybrid approach. The sensitivity of the predicted particle mass concentrations and size distribution to the three different gas/particle mass transfer approaches is presented in section 4.5.4.

3. Formulation of Other Processes Related to PM

[39] Besides the processes that directly govern the chemical composition and size distribution of PM, other processes that also affect PM size distributions must be considered. Such processes include cloud processes that lead to additional PM mass via aqueous-phase reactions in cloud droplets and subsequent evaporation; heterogeneous reactions at the surface of cloud droplets and particles that produce additional PM; and deposition processes that remove particles of different sizes. We describe the modifications that were made to CMAQ to account for those additional processes in conjunction with the incorporation of MADRID into CMAQ.

3.1. Cloud Processes

[40] The original CMAQ cloud module was modified to include a more comprehensive aqueous-phase chemical mechanism and to provide treatment of cloud processing of aerosols (e.g., particle activation and scavenging, particle formation from droplet evaporation) that is based on a sectional size representation. We describe these modifications below.

3.1.1. Aqueous-Phase Chemistry

[41] The original CMAQ includes a simplified aqueous-phase chemical mechanism that is based on the RADM mechanism [Walcek and Taylor, 1986]. It includes only 10 gas-aqueous equilibria, 9 aqueous equilibria and 5 kinetic reactions for the oxidation of SO$_2$ to sulfate. Because the solubility of SO$_2$ and the oxidation rates of dissolved SO$_2$ species depend on the acidity of the cloud or fog droplet, the aqueous-phase chemical mechanism needs to include a fairly long list of species that can affect the acidity of atmospheric droplets. The requirements for such an aqueous-phase chemistry module are (1) a robust and efficient numerical solver and (2) a relatively complete representation of the aqueous chemistry of sulfur and nitrogen species. We selected a more comprehensive chemical kinetic mechanism, i.e., the CMU bulk aqueous-phase chemical mechanism [Pandis and Seinfeld, 1989; Seinfeld and Pandis, 1998; Fahey and Pandis, 2001] and incorporated it as an option into CMAQ. The CMU mechanism includes 17 gas-aqueous equilibria, 17 aqueous equilibria, and 99 aqueous-phase kinetic reactions among 18 gas-phase species and 27 aqueous-phase species. Like all available aqueous-phase mechanisms, it is designed to simulate sulfate production from SO$_2$ in atmospheric liquid water and includes the three dominant oxidation pathways for S(IV) by H$_2$O$_2$, O$_3$, and O$_2$ catalyzed by Fe$^{3+}$ and Mn$^{2+}$. In addition, it includes other reactions for the oxidation of S(IV) to sulfate, the oxidation of nitrogen species to nitrate, and reactions for carbonate, chlorine, organic and oxygen species that are involved in the formation of sulfate and/or nitrate species.

[42] The mechanism was implemented with options that allow the user to reduce the number of aqueous-phase chemical reactions used in a particular simulation. The user can include or exclude (1) the chlorine chemistry, (2) the radical chemistry, and (3) the Fe$^{3+}$ and Mn$^{2+}$ catalyzed oxidation of dissolved SO$_2$. These options are provided because these portions of the chemistry may have only small effects on the results and sometimes present very stiff numerical conditions that require significant amounts of computer time. Neglecting the radical chemistry in polluted environments improves computational efficiency without introducing significant errors. The concentrations of radical species (i.e., OH, HO$_2$, O$_2$, NO$_3$, ClO$^-$, SO$_4$ and CO$_3^-$) become zero in the aqueous phase when the radical chemistry is turned off. In this case, the CMU mechanism consists then of 14 gas-aqueous equilibria, 16 aqueous equilibria, and 32 aqueous kinetic reactions among 18 gas-phase species and 19 aqueous-phase species.

3.1.2. Aerosol Activation and Scavenging

[43] Aerosol activation and scavenging contribute to the species concentrations in cloud droplets. The activated or scavenged fraction is influenced by many factors including the types of clouds, cloud supersaturation, the characteristics of aerosols (e.g., number concentration, size distribution, chemical composition and solubility), and updraft velocity. Both empirical and mechanistic parameterizations have been used to simulate aerosol activation by cloud droplets in 3-D models. In the interest of computational efficiency, we developed an empirical parameterization for the sectional size representation based on available observations, as shown in Table 1.
For a relatively fine resolution of the particle size representation (i.e., 6 or more size sections between 0.02 and 10 \( \mu m \)), particles with aerodynamic diameter greater than 0.35 \( \mu m \) are assumed to be 100\% activated (this diameter corresponds to a Stokes diameter of 0.3 \( \mu m \) for a density of 1.35 g cm\(^{-3}\)). The use of 0.35 \( \mu m \) as an activation cutoff diameter is a reasonable approximation because particles larger than this size are expected to require a supersaturation less than 0.03\% for activation, a prevailing condition for most clouds [Gillani et al., 1995]. Note that a larger activation cutoff diameter may be needed for episodes with heavy fogs. For example, few particles smaller than 1 \( \mu m \) are activated by fogs (because of much smaller supersaturations) in Los Angeles or San Joaquin Valley of California. Particles with aerodynamic diameter greater than 0.1 \( \mu m \) (i.e., a Stokes diameter of 0.086 \( \mu m \) for a density of 1.35 g cm\(^{-3}\)) but smaller than 0.35 \( \mu m \) are assumed to be partially activated. Both observations and theoretical calculations have shown that the fraction of activation for particles in this size range can vary from 0 to 1 [Gillani et al., 1995; Leaitch, 1996; Liu et al., 1996]. We assume that 50\% of the particle mass in this size range is activated completely into cloud droplets, and that the remaining 50\% of particle mass is slowly scavenged according to \( \exp(-\beta \tau) \), where \( \beta \) is the mass scavenging coefficient for particles with \( dp \leq 2.5 \mu m \), \( \tau \) is the cloud lifetime. It is determined by simulated cloud properties (e.g., droplet size distribution, liquid water content and settling velocity) and aerosol properties (e.g., particle size distribution and polydispersity diffusivity) based on the equation of Pruppacher and Klett [1980] [Binkowski, 1999]. \( \tau \) is the cloud chemistry time step for grid-resolved cloud and the cloud lifetime for subgrid convective clouds. A 1-hour lifetime is assumed for subgrid convective clouds [Byun and Ching, 1999]. Particles with aerodynamic diameter less than 0.1 \( \mu m \) are assumed to remain as interstitial particles with 0\% activated. For a coarse size resolution (i.e., 2 to 5 size sections between 0.02 and 10 \( \mu m \)), an activation cutoff diameter of 2.5 \( \mu m \) is used and the activated fraction for particles smaller than 2.5 \( \mu m \) is assumed to be 80\%.

### 3.1.3. Formation of Particles After Cloud Evaporation

The particle concentrations after cloud evaporation are calculated as follows. The change in mass concentration of individual particulate components during the cloud lifetime (i.e., between cloud formation and evaporation) is calculated first. This change is then added to the activated particle size distribution using a uniform relative change across the particle size distribution. For example, if the total change in sulfate concentration is a 10\% increase, the sulfate concentrations in each particle size section are increased by 10\%. The activated mass along with its changes due to aqueous-phase chemistry is then added back to the interstitial mass size distribution to obtain the total particle mass size distribution after cloud evaporates.

If a 2-section particle size representation is used, no particle mass movement is simulated between the 2 size sections. For a multisection particle size representation, however, it is necessary to account for particle growth that occurs due to the increase in mass in each size section. The moving-center scheme is used to calculate the new particle size distribution that results from this particulate mass increase.

### 3.2. Heterogeneous Chemistry

Jacob [2000] conducted a review of heterogeneous chemistry and recommended that the heterogeneous reactions of \( \text{HO}_2 \), \( \text{NO}_2 \), \( \text{NO}_3 \) and \( \text{N}_2\text{O}_5 \) on the surface of aqueous particles and cloud droplets be parameterized by a simple reaction probability in 3-D \( \text{O}_3 \) models. In this parameterization, the uptake of a gas-phase species by condensed phases is considered as an irreversible loss process with a first-order heterogeneous reaction rate constant. On the basis of Jacob [2000], we consider the following heterogeneous reactions on the surface of aqueous particles or cloud/fog droplets in CMAQ-MADRID:

\[
\begin{align*}
\text{(R1)} & \quad \text{HO}_2 & \quad \text{PM} & \quad 0.5\text{H}_2\text{O}_2 \\
\text{(R2)} & \quad \text{NO}_2 & \quad \text{PM} & \quad 0.5\text{HONO} + 0.5\text{HNO}_3 \\
\text{(R3)} & \quad \text{NO}_3 & \quad \text{PM} & \quad \text{HNO}_3 \\
\text{(R4)} & \quad \text{N}_2\text{O}_5 & \quad \text{PM} & \quad 2\text{HNO}_3 \\
\text{(R5)} & \quad \text{N}_2\text{O}_3 & \quad \text{Cloud/Fog} & \quad \text{2HNO}_3 
\end{align*}
\]

Recent experimental data suggest that bulk aqueous-phase chemistry is consistent with the rate of the reaction of \( \text{NO}_2 \) in presence of condensed water [Cheung et al., 2000]. Following discussions with D. Jacob (Harvard University, personal communication, 2001), we elected not to include the heterogeneous reaction of \( \text{NO}_2 \) on droplets; it is, however, included in the CMU bulk aqueous-phase mechanism. In addition, the heterogeneous reactions of \( \text{HO}_2 \) and \( \text{NO}_3 \) on droplets were not included, since the bulk CMU aqueous-phase mechanism already includes the scavenging of \( \text{HO}_2 \) and \( \text{NO}_3 \) by cloud droplets and their subsequent aqueous-phase equilibria and reactions. The first-order rate constant, \( k \), for the heterogeneous loss of a gaseous species \( i \) to the condensed phase is calculated following Jacob [2000]:

\[
k_i = \left( \frac{a}{P_{gi}} + \frac{4}{v_i v_y} \right)^{-1} A, \tag{7}
\]
where \( a \) is the radius of particles or cloud/fog droplets, \( D_{gi} \)
and \( v_i \) are the gas-phase molecular diffusion coefficient and
the mean molecular speed, respectively, of a gaseous species \( i \)
in air, \( \gamma \) is the reaction probability of species \( i \), which
represents the likelihood that a gas molecule impacting the
surface of the condensed phase will undergo reaction, and \( A \)
is the surface area of the condensed phase. For atmospheric
particles, \( A \) is obtained by integrating over the particle size
distribution. The first term on the right-hand side of
equation (7) represents the uptake by diffusion from the bulk
gas phase to the surface of the condensed phase and the
second term represents the uptake by free molecular
collisions of gas molecules with the surface. If \( k_i \rightarrow (D_{gi}/a) \)
(e.g., (R1)), the uptake of a gas molecule by the condensed
phase is diffusion-limited and shows little dependence on the
value of \( \gamma_i \). On the other hand, if \( k_i \rightarrow (\nu_i \gamma_i A/4) \) (e.g., (R2)
and (R3)), the uptake of a gas molecule tends to be limited by free
molecular collision and depends strongly on the magnitude of
\( \gamma_i \). The uptake of species like \( N_2O_5 \) can be either fully
diffusion-limited (e.g., on cloud droplets with diameter
greater than 20 \( \mu \)m and \( \gamma_i = 0.1 \)), partially diffusion-limited
(e.g., for high concentrations of particles with \( \gamma_i = 0.1 \)) or in a
transition regime (for cloud droplets with diameter greater
than 20 \( \mu \)m and \( \gamma_i = 0.01 \) or high concentrations of particles
with \( \gamma_i = 0.01 \)), depending on the diameter and concentrations
of the particle/droplet and the values of the reaction
probabilities used. The values of the reaction probabilities
selected in MADRID are the nominal values listed by Jacob
[2000]; i.e., 0.2, \( 1.0 \times 10^{-4} \), \( 1.0 \times 10^{-3} \) and 0.1 for \( HNO_3 \),
\( NO_2 \), \( NO_3 \) and \( N_2O_5 \), respectively.

There are several important yet uncertain variables in the
calculation of \( k_i \), such as \( \gamma_i \) and \( A \). \( \gamma_i \) has been measured in the laboratory for a number of gases on various con-
densed phases but its values may differ by several orders of
magnitudes for a given species on different types of
surfaces. The surface area of particles depends on the
particle number concentrations and size distribution, which
exhibit high temporal and spatial variabilities. The surface
area of cloud/fog droplets depends on the diameters of
cloud/fog droplets and the cloud/fog liquid water content,
which can be quite different for different types of clouds/
fogs. In addition, \( k_i \) is a function of ambient temperature and
pressure because \( v_i \) is temperature-dependent and \( D_{gi} \)
is temperature- and pressure-dependent.

The particle size distribution simulated by MADRID
is used to calculate the particle surface area. The droplet size
distribution in CMAQ is assumed to be lognormal with a
fixed geometric standard deviation and a variable diameter.
In CMAQ-MADRID, we assumed that cloud or fog droplets are
monodispersive. A droplet diameter of 20 \( \mu \)m was used in the
base calculation. Note that the heterogeneous rates of
diffusion-limited reactions ((R1) and (R5)) are quite sensitive
to the droplet diameter, varying by about one order of
magnitude as the droplet diameter varies from 5 \( \mu \)m to 20 \( \mu \)m.

### 3.3. Atmospheric Deposition

#### 3.3.1. Dry Deposition

New precursor and condensable organic species were
added in CMAQ for the treatment of SOA formation. Condensable gases typically contain multiple functional
groups, such as aldehyde, acid and alcohol groups. Without
information on the identities of the condensable compounds
in the smog chamber experiments of Odum et al. [1997] and
Griffin et al. [1999], deposition velocities of the organic
gases are assumed to be analogous to that of higher
aldehydes. To simulate the dry deposition of particles, we used the algorithm of Venkatram and Pleim [1999]:

\[
V_d = \frac{V_s}{1 - \exp(-r_a r_b V_s)},
\]

where \( V_d \) is the total dry deposition velocity of the particle,
\( V_s \) is the gravitational settling (sedimentation) velocity,
\( r_a \) is the aerodynamic resistance in the lower atmosphere
and \( r_b \) is the resistance in the quasi-laminar layer near the surface.
This approach conserves mass because it accounts for the
fact that the resistance component depends on a concentra-
tion gradient whereas the sedimentation term does not.

The particle dry deposition velocity is calculated for each
particle size section and the dry deposition flux is calculated
accordingly by size section.

#### 3.3.2. Scavenging and Wet Deposition

Wet deposition is treated similarly in the CMAQ-MADRID cloud module as in the original CMAQ cloud module [Roselle and Binkowski, 1999]. If precipitation occurs, the column from the surface to the cloud base is
treated as being in equilibrium between the gas phase and
the droplets. The CMU bulk aqueous-phase chemistry
module calculates the droplet concentrations of dissolved
gaseous species and activated particulate species in the
cloud. Below the cloud, the droplet concentrations of the
soluble gases are calculated using their Henry’s law con-
stants and the particles are assumed to be completely
absorbed into the rain droplets. The column-weighted
droplet concentrations are then multiplied by the precipita-
tion rate to calculate the wet deposition fluxes.

One modification was made to the treatment of
below-cloud scavenging of gases (washout). In the original
CMAQ formulation, the solubility of gases into raindrops is
calculated using the Henry’s law constant. For chemical
species that dissociate in aqueous solutions such as acids
e.g., \( HNO_3 \), \( HCl \) and bases (e.g., \( NH_3 \)), the solubility
is then underestimated. In CMAQ-MADRID, we take into
account the aqueous dissociation reactions by using the
effective Henry’s law constant.

### 4. Application of CMAQ-MADRID

CMAQ-MADRID was applied to simulate the
August 1987 Southern California Air Quality Study
(SCAQS) episode in the Los Angeles basin. Figure 1 shows
the simulation domain and the locations of 38 O3 measure-
ment sites and 8 PM sampling sites in the basin. SCAQS
provides a comprehensive database needed for model inputs
and evaluation. This episode has been used earlier for the
evaluation of PM air quality models [see Seigneur, 2001] and
therefore it provides a convenient benchmark. The PM
measurements include concentrations of total PM\(_{2.5}\) mass,
total PM\(_{10}\) mass, sodium (N\textsubscript{a}), sulfate (SO\textsubscript{2}\textsuperscript{2-})
ammonium (NH\textsubscript{4}\textsuperscript{+}), nitrate (NO\textsubscript{3}\textsuperscript{-}),
chloride (Cl\textsuperscript{-}), elemental carbon (EC, also referred to as black carbon), and organic carbon
(OC). A factor of 1.4 was used to convert the measured OC
to organic material (OM) for model comparison [White and
Roberts, 1977]. Although other inorganic species such as

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dust, calcium and iron were not explicitly measured, they contribute significantly to the total PM mass concentration. For example, they account for up to 28% and 53% of the measured 24-hour average PM$_{2.5}$ and PM$_{10}$ mass at the 8 PM sampling sites. In our model, the simulated PM species include Na, SO$_4^{2-}$, NH$_4^+$, NO$_3^-$, Cl$^-$, EC and OM. All other species such as crustal materials are lumped together as other inorganic species (OI).

In the SCAQS simulation, the southwest corner of the modeling domain was placed at 33°18'N latitude and 119°24'W longitude. The horizontal grid system consists of $63 \times 28$ grid cells, with a grid resolution of $5 \times 5$ km$^2$. 30 layers of the MM5 grid system were mapped to 15 layers of CMAQ-MADRID, with a one to one mapping near the surface. Two base simulations were conducted, one with 2 size sections and the other with 8 size sections in the size range of 0.0215 and 10 μm to represent particle size distribution. The 2 size sections are 0.0215–2.15 μm and 2.15–10 μm, and the 8 size sections are 0.0215–0.0464 μm, 0.0464–0.1 μm, 0.1–0.215 μm, 0.215–0.464 μm, 0.464–1 μm, 1–2.15 μm, 2.15–4.64 μm, and 4.64–10 μm. The CBM-IV gas-phase chemical mechanism, the CMU bulk aqueous-phase chemical mechanism, and the MADRID 1 aerosol module were used in the two base simulations and in all sensitivity simulations except for a sensitivity simulation in which MADRID 2 was used to evaluate the sensitivity to different SOA formulations. The heterogeneous reactions of HO$_2$, NO$_2$, NO$_3$, and N$_2$O$_5$ on the surface of particles and that of N$_2$O$_5$ on cloud droplets were also accounted for in the base and all sensitivity simulations except one sensitivity simulation in which those heterogeneous reactions were turned off. The particle growth and gas/particle mass transfer were simulated with the moving-center scheme and the CMU hybrid approach, respectively, in the base simulations. Additional simulations were conducted with the finite difference scheme for particle growth and the CIT bulk and a simple bulk equilibrium approaches for gas/particle mass transfer in the sensitivity studies.

4.1. Meteorology

The meteorological fields were simulated using the meteorological Mesoscale Model version 5 (MM5) with four-dimensional data assimilation. This MM5 simulation has been described previously [Hegarty et al., 1998] and used in previous air quality simulations [Pai et al., 2000; Seigneur et al., 2000a, 2000b]. As discussed in these earlier results, the meteorological fields were mispredicted during daytime (particularly on 28 August at inland locations), which led to overestimated vertical mixing. To minimize the impact of such meteorological inputs on the air quality simulations, we added a post-processing step to the MM5 outputs by developing a diagnostic field to the MM5 outputs by developing a diagnostic field.
originally based on the 1982 NH₃ emission inventory of Cass and Gharib [1984].

Larger uncertainties exist in the primary PM emissions, particularly in their size distribution. The emissions of PM₂.₅ and PM₁₀₂.₅ and PM chemical speciation are based on Meng et al. [1998], which was originally created by Lurmann et al. [1997] based on the California Air Resources Board’s original PM emission inventories. The chemical composition of PM emissions includes SO₄²⁻, Na⁺, Cl⁻, EC, OM, and OI (e.g., crustal material) in PM₂.₅ and PM₁₀₂.₅ size ranges. Note that the OM emissions were obtained by adding 40% to the original OC emissions in the CARB’s inventories to account for the oxygen and hydrogen associated with the OC and the original zero OM emission fractions for many emission categories were adjusted by reallocating the original EC fractions between EC and OM [Lurmann et al., 1997]. The adjusted OM emissions are still low compared to those in updated inventories [Eldering and Cass, 1996; Kleeman and Cass, 1998]. Two adjustments were made in this study to the PM chemical speciation used by Meng et al. [1998]. First, 71% of total EC emissions were assigned to the sub-2.5 μm size range. This value is based on the observed mean mass ratio of sub-2.5 μm EC (EC₂₅) to sub-10 μm (EC₁₀) during this SCAQS episode. (EC₂₅ accounts for 78% and 80% of total EC emissions in the work of Meng et al. [1998] and Jacobson [1997b], respectively; however, EC₂₅ concentrations were overpredicted with a bias of 30–35% in both works). Second, we assumed that sea salt emissions are 32 tons day⁻¹ with 10% in the sub-2.5 μm size range. Neither the SCAQMD nor the CARB emission inventory includes sea salt emissions which produce most of the Na⁺ and Cl⁻ mass. Lurmann et al. [1997] and Meng et al. [1998] assumed a total NaCl emission of 75 tons day⁻¹ with 29% in the sub-2.5 μm size range. Lurmann et al. [1997] reported a moderate overprediction of Na⁺ and Cl⁻ with bias of 38% and 24%, respectively. The total oceanic area covered in the simulated domain in the work of Meng et al. [1998] is roughly two times larger than that in the current domain. We scaled down the emission rate of 75 tons day⁻¹ to 52 and 32 tons day⁻¹ in two test simulations. Better agreement between simulated and observed NaCl mass was obtained with 32 tons day⁻¹; this value was therefore used in our simulation. For simulations with 8 size sections, PM emissions were assigned to sections according to the default size distribution of CMAQ [Byun and Ching, 1999].

4.3. Initial and Boundary Conditions

Initial conditions (IC) and boundary conditions (BC) for gases follow Pai et al. [2000]. IC and BC for PM in the fine and coarse size ranges were specified into SO₄²⁻, NO₃⁻, NH₄⁺, Na⁺, Cl⁻, EC, OM, and OI using their observed concentrations in the two size ranges from San Nicholas Island, a “background” site during SCAQS 1987. For simulations with 8 size sections, the mass distribution of IC and BC for PM species in each section was determined using the default size distribution of CMAQ [Byun and Ching, 1999].

4.4. Results and Discussion

All SCAQS simulations were conducted for the period starting at 4:00 PST on 25 August and ending at 4:00 PST on 29 August 1987. The first two days were used as spin-up days to generate IC for 27 August. Results are analyzed and presented for 27 and 28 August.

4.4.1. Predicted O₃ Mixing Ratios at SCAQS Sampling Sites

Figure 2 shows 2-day time series plots of observed and predicted O₃ mixing ratios at 12 monitoring sites selected to represent various parts of the basin. The predicted O₃ mixing ratios from the two base simulations (with 2 and 8 size sections) are very similar, thus only the predictions with 2 size sections are shown. The sites include El Rio (ELRI), Piru (PIRU), Reseda (RESE), and Thousand Oaks (THSO) in the San Fernando Valley in the northwestern basin (Figures 2a, 2b, 2c, and 2d); central and west Los Angeles (CELA and WSLA) in the western urban and suburban areas (Figures 2e and 2f); Pomona (POMA) and Riverside (RIVR) in the central and eastern downwind inland area (Figures 2g and 2h); Crestline (CRES) and Hesperia (HESP) in the northern rural and remote areas (Figures 2i and 2j); and Long Beach (LBCC) and El Toro (TORO) in the middle and southern coastal areas (Figures 2k and 2l).

The observed peak O₃ mixing ratios in the basin typically occurred in the early to midafternoon between noon and 4 p.m. In the San Fernando Valley, the predicted daytime and peak O₃ mixing ratios agreed well with the observations, but the predicted time of peak O₃ was sometimes off by a couple of hours. At the central and west Los Angeles sites, the predicted daytime and peak O₃ mixing ratios on 28 August and the time of peak O₃ on both days matched very well with observations, but the O₃ mixing ratios on 27 August were overpredicted by 69% and 54% at CELA and WSLA, respectively. In the central and eastern locations downwind, the predicted time of peak O₃ was slightly delayed and the peak and some daytime O₃ mixing ratios were underpredicted at POMA (by 32%) on 28 August and at RIVR (by 31% and 53%) on both days. In the northern rural and remote areas (CRES and HESP), the observed daytime O₃ mixing ratios were well predicted with a moderate underprediction (by 22%) in the peak O₃ mixing ratio at CRES on 27 August. In the middle to southern coastal sites (LBCC and TORO), the predicted time of peak O₃ was slightly off. The peak O₃ mixing ratios were overpredicted moderately at LBCC on 28 August (by 36%) and at TORO on both days (by 33% and 18%) and significantly at LBCC on 27 August (by 110%).

Overall, the model simulations reproduce the magnitude and the spatial and temporal variations of O₃ mixing ratios throughout the basin, but tend to overpredict daytime O₃ mixing ratios at a few sites in the western urban areas and in the middle to southern coastal areas on 27 August and underpredict the O₃ mixing ratios at inland sites downwind on 28 August. The underpredictions inland are due, at least in part, to overpredictions of the vertical mixing and wind speeds by the meteorological model inland on that day. The uncertainty in precursor emissions may also contribute to the underpredictions. The model also significantly overpredicts the O₃ mixing ratios at night at all sites except El Rio, Long Beach, El Toro, and Hesperia, which is due in part to an underestimation of NO mixing ratios, as a result of an overestimation of vertical mixing in the surface layer, and in part to a relatively high ceiling of the first model layer (i.e., 60 m) used.
4.4.2. Predicted PM Chemical Composition and Size Distribution

Figures 3–6 show the observed and simulated 24-hour average concentrations for PM$_{2.5}$, PM$_{10}$ and their chemical compositions on 27 and 28 August at four sites: Hawthorne (HAWT), Los Angeles (CELA), Azusa (AZUS), and Riverside (RIVR). They represent an upwind coastal site in the western basin, a downtown area with high motor vehicle emissions, a mid-basin rural/suburban site, and a downwind urban site in the eastern basin, respectively. The model predictions with both two and eight size sections are compared to the observations.

The observed concentrations at the four sites show the evolution of PM$_{2.5}$ and PM$_{10}$ across the basin from the coast to the east. The observed PM concentrations were relatively low near the coast but became significantly higher as the air mass was transported across the basin. PM reached its highest level among all monitoring sites at

Figure 2. Two-day time series of observed and predicted O$_3$ mixing ratios at 12 monitoring sites selected to represent various parts of the basin. (a) El Rio (ELRI); (b) Piru (PIRU); (c) Reseda (RESE); (d) Thousand Oaks (THSO); (e) central Los Angeles (CELA); (f) west Los Angeles (WSLA); (g) Pomona (POMA); (h) Riverside (RIVR); (I) Crestline (CRES); (j) Hesperia (HESP); (k) Long Beach (LBCC); (l) El Toro (TORO).
RIVR on both days. The observed sub-2.5 μm and sub-10 μm sulfate (sulfate$_{2.5}$ and sulfate$_{10}$) concentrations ranged across the basin from 5.7 to 10 μg m$^{-3}$ and 6.8 to 12 μg m$^{-3}$, respectively. The observed sub-2.5 μm and sub-10 μm nitrate (nitrate$_{2.5}$ and nitrate$_{10}$) concentrations were relatively low near the coast, but increased significantly downwind of the NH$_3$ source areas in the eastern part of the basin, and the highest ammonium and nitrate concentrations occurred at RIVR. EC and OM concentrations were relatively low near the coast, but increased significantly in the downtown source area and in the northern and eastern basin. AZUS had the highest EC$_{2.5}$ and OM$_{10}$ concentrations on both days, the highest EC$_{10}$ on 27 August, and the highest OM$_{2.5}$ on 28 August. The highest EC$_{10}$ occurred at RIVR on 28 August. The highest OM$_{2.5}$ occurred at CLAR on 27 August (not shown), a northern site close to AZUS (OM$_{2.5}$ at AZUS was the second highest across the basin). Compared to other PM compositions, the observed sub-2.5 μm and sub-10 μm sodium (sodium$_{2.5}$ and sodium$_{10}$) and chloride (chloride$_{2.5}$ and chloride$_{10}$) concentrations were relatively low, with a range of 0.04–0.34, 1.28–2.31, 0.04–0.4 and 0.31–1.08 μg m$^{-3}$ at all 8 sites, respectively. Note that

Figure 2. (continued)
most measured chloride_{2.5} and some measured chloride_{10} concentrations were below the detection limit of 0.5 µg m\(^{-3}\) for Teflon filters [Fraser et al., 1996].

[65] Similar to O\(_3\) predictions, the model tends to overpredict PM concentrations at coastal and western sites on 27 August but underpredict PM concentrations at inland sites downwind on 28 August. This can be partially attributed to the overpredictions on 27 August and the underpredictions on 28 August in the concentrations of precursors for secondary PM as a result of mispredicted wind speeds and vertical mixing. Other factors that are responsible for the discrepancies between observations and predictions are discussed later along with model performance evaluation. Although some discrepancies exist in the magnitudes of observed and simulated PM and its composition at some locations, the simulations with both size resolutions reproduce well the observed evolution of PM and its composition. The predicted sulfate concentrations from the two base simulations match well with the observed values at all sites except at HAWT, where there was a significant overprediction (by 67% for sulfate_{2.5} and 51% for sulfate_{10}) on 27 August and underprediction (by 51% for sulfate_{2.5} and 54% for sulfate_{10}) on 28 August, and at CELA, where there was a significant overprediction (by 74% for sulfate_{2.5} and 55% for sulfate_{10}) on 27 August.
The predicted ammonium and nitrate concentrations from both simulations are lower near the coast and higher at downwind locations, with the highest nitrate concentrations occurring at RIVR on both days. The highest nitrate concentrations predicted at RIVR are 28.9 and 18.3 \( \text{mg m}^{-3} \) with 2 size sections and 29.8 and 18.8 \( \text{mg m}^{-3} \) with 8 size sections on 27 and 28 August. For comparison, the observed nitrate concentrations at RIVR were 33.3 and 39.8 \( \text{mg m}^{-3} \) on 27 and 28 August, respectively. The nitrate concentrations predicted with 2 size sections agree well with the observations at CELA and RIVR on 27 August (with a deviation of 8–18%), but show moderate to significant underpredictions (by 32–88%) on 28 August and at other sites on both days. The nitrate concentrations predicted with 8 size sections are similar to those with 2 size sections except that there is a moderate overprediction (by 41–42% for nitrate\(_{2.5}\) and 47–52% for nitrate\(_{10}\)) at HAWT and CELA on 27 August. The predicted nitrate\(_{2.5}\) and nitrate\(_{10}\) concentrations with 8 size sections are 5.3 and 6.8 \( \text{mg m}^{-3} \) at HAWT and 12 and 13.9 \( \text{mg m}^{-3} \) at CELA. For comparison, the observed values were 3.7 and 4.5 \( \text{mg m}^{-3} \) at HAWT and 8.5 and 9.5 \( \text{mg m}^{-3} \) at CELA, respectively. Such an overprediction seems large in terms of percentage but less significant in terms of the absolute values. Moreover, the observed nitrate\(_{10}\) concentrations are likely underestimated due to mass losses on the Teflon filter [Hering et al., 1997]. The sensitivity of the predicted nitrate concentrations to size resolution has been studied in the work of Koo et al. [2003]. They found that lower nitrate concentrations are predicted by the CMU hybrid approach when a lower size resolution is used, which is consistent with our results. This sensitivity is primarily caused by the mixing of aerosol populations with different composition (e.g., the larger alkaline particles are mixed with the smaller acidic particles in coastal areas) that usually results in an underprediction of the nitrate concentrations when a lower size resolution is used [Capaldo et al., 2000]. For a higher size resolution, more nitrate can be formed in the fine particle size range due to the “bulk equilibrium treatment” and the weighting scheme that is based on surface area, which assigns more condensates to smaller size sections [Koo et al., 2003]. Differences in the removal efficiency and the second-order thermodynamic effects that are size-dependent (thus size resolution-dependent) may also contribute to the predicted sensitivity. Compared to nitrate predictions, better agreement is obtained between the simulated (with both size resolutions) and observed ammonium concentrations at all sites.

EC concentrations are underpredicted by 41–67% with both size resolutions at almost all sites on both days, due mainly to uncertainties in meteorological inputs, EC emissions and the size distribution of emitted EC. The OM concentrations are also underpredicted moderately on 28 August in the western downtown area (e.g., by 37–38% for OM\(_{2.5}\) at CELA) and significantly near the coast (e.g., by 59–60% for OM\(_{2.5}\) at HAWT), in some areas in the northern basin (e.g., by 57–58% for OM\(_{2.5}\) at AZUS but not at CLAR) and downwind in the eastern basin (e.g., by 68–69% for OM\(_{2.5}\) at RIVR). The OM predictions on 27 August...
show a better agreement with observations at all sites (e.g., underpredictions by 18–22%, 12–17%, 37–39% and 52–53% for OM2.5 at CELA, HAWT, AZUS and RIVR, respectively). The predicted sodium2.5 and chloride2.5 concentrations in the two base simulations are in the range of 0.04–0.13 and 0.01–0.31 $\text{mg} \cdot \text{m}^{-3}$ at the four sites, which are generally within the observed range (0.04–0.34 and 0.06–0.4 $\mu \text{g} \cdot \text{m}^{-3}$, respectively). The predicted sodium10 and chloride10 concentrations are in the range of 0.7–1.16 and 0.11–0.82 $\mu \text{g} \cdot \text{m}^{-3}$ at these sites. At most of these sites, the model tends to underpredict the observed sodium10 and chloride10 concentrations (1.28–2.31 and 0.31–1.08 $\mu \text{g} \cdot \text{m}^{-3}$, respectively) by 36–58% and 20–90%.

[68] Particle size distributions of individual inorganic ions and organic species were measured with the electrical aerosol analyzer (EAA), the laser optical particle counter (OPC), the 9-stage Berner impactors (BI, for inorganic ions) and the micro-orifice uniform deposit impactor (MOUDI, for organic ions) in 11 summer time intensive sampling days at Claremont and Riverside during 1987 SCAQS [e.g., John et al., 1990; Hering et al., 1997]. The measured particle aerodynamic diameter range is 0.03–0.3 $\mu \text{m}$ for EAA, 0.1–3 $\mu \text{m}$ for OPC, 0.075–16.5 $\mu \text{m}$ for BI and 0.1–3 $\mu \text{m}$ for MOUDI [Hering et al., 1997]. Hering et al. [1997] obtained a representative particle mass distribution over the size range of 0.03 to 3 $\mu \text{m}$, measured by EAA, OPC and impactors (BI and MOUDI), for Claremont and Riverside by averaging all summer samplings to compare PM measurements with different techniques and to study the characteristics of PM mass size distribution in southern California. The observed representative PM mass distributions summed those of individual PM species including $\text{SO}_4^{2-}$, $\text{NO}_3^-$, $\text{NH}_4^+$, EC and OM (= 1.4 × measured OC). Figure 7 compares the predicted PM mass size distribution for the above five PM components on 28 August 1987 with the impactor average PM mass size distribution obtained by Hering et al. [1997] at Claremont and Riverside. We discuss here the base 8-section simulation results obtained with the moving-center scheme and the CMU hybrid approach. Results with other combinations of algorithms for particle growth and gas/particle mass transfer are also shown in Figure 7 but will be discussed later in the sensitivity study.

[69] The model correctly predicts a unimodal distribution for accumulation mode PM that is typical at Riverside during the summer time but it fails to reproduce the typical bimodal distribution of accumulation mode PM at Claremont (a bimodal distribution was, however, predicted for some time periods at Claremont and in other locations, see section 4.5.3). The predicted peak value at Riverside is higher than that at Claremont and both peak values occur in the same size range, consistent with observations. However, the PM mass concentration peaks in the size section of 0.215–0.464 $\mu \text{m}$ at both sites, which is somewhat off from the diameter of the observed peak values (i.e., 0.52 $\mu \text{m}$). The peak values are underpredicted by 30% at Claremont and 16% at Riverside. In addition to the aforementioned factors (e.g., mispredictions in meteorology and uncertainties in emissions of primary PM species and precursors of

![Figure 4. Observed and predicted 24-hour average concentrations for PM2.5, PM10 and their chemical compositions on 27–28 August 1987 at central Los Angeles (CELA), CA.](image-url)
secondary PM species) that contribute to differences between observed and simulated PM mass and composition, the discrepancy in observed and predicted particle size distribution indicate that a finer size resolution (>8 sections) is needed to accurately simulate PM size distribution.

4.4.3. Model Performance Evaluation

The model performance is evaluated following the guidance developed by Seigneur et al. [2000c]. Our evaluation focuses on the mean normalized gross error (MNGE) and mean normalized bias (MNB) in the O3 and PM predictions at the sampling sites. Table 2 shows the MNGE and MNB for 1-hour average O3 mixing ratios and 24-hour average concentrations of PM2.5, PM10 and their components averaged over all sampling sites on 27 and 28 August 1987. Table 2 also shows the ranges of MNGE and MNB obtained from simulations of SCAQS episodes with other 3-D air quality models including GATOR [Jacobson, 1997b], CIT [Harley et al., 1993, 1997; Meng et al., 1998], and UAM-AERO [Lurmann et al., 1997]. GATOR and CIT were applied to the same episode whereas UAM-AERO was applied to a June 1987 episode. Only statistics obtained from the CMAQ-MADRID simulation with 2 sections are presented in Table 2. In the following analysis and discussion, we focus on results from the simulation with 2 size sections. The differences between the results in the two base simulations will be discussed in the relevant sections below.

The MNGE in O3 predictions with both size resolutions at 38 sites are 36% and 20% for 27 August and 31% and 3% for 28 August. The statistical values were calculated for each day using a cut-off value of 40 ppb for O3. The MNGE and MNB are within the range of other work [e.g., Jacobson, 1997b; Meng et al., 1998; Lurmann et al., 1997].

The performance statistics for PM2.5, PM10 and their components were calculated at 8 PM sampling sites in the modeling domain. The predicted mean PM2.5 and PM10 mass concentrations averaged over all 8 locations are 51.8 and 91.7 μg m⁻³ on 27 August, which overpredict the mean observed values of 42.3 and 75.9 μg m⁻³ by 21–22%. The predicted mean PM2.5 and PM10 concentrations averaged over all 8 locations are 43.4 and 81.3 μg m⁻³ on 28 August, which compare well with the observed values of 48.1 and 85.3 μg m⁻³. The MNGE and MNB in the predicted PM2.5 concentrations are 42% and 30% on 27 August and 47% and 2% on 28 August. The MNGE and MNB in PM10 concentrations are slightly higher (53% and 33% for 27 August and 56% and 8% for 28 August). For the simulation with 8 size sections, the MNGE and MNB in the predicted PM2.5 concentrations are very similar to those with 2 size sections on 28 August but slightly higher on 27 August (45% and 32%, respectively). These performance statistics for both PM2.5 and PM10 mass concentrations are consistent with those obtained with the other three models.

The predicted sulfate2.5 and sulfate10 concentrations match the observations well at all sites, with MNGE and
MNB of 27% to 49% and −27% to 49% for PM$_{2.5}$ sulfate and 35% and −35% to 32% for PM$_{10}$ sulfate. The MNGE and MNB are lower with 8 size sections for PM$_{2.5}$ and PM$_{10}$ sulfate (40% and 38% for PM$_{2.5}$ sulfate and 31% and 25% for PM$_{10}$ sulfate) on 27 August but slightly greater (30% and −30% on for PM$_{2.5}$ sulfate and 37% and −37% for PM$_{10}$ sulfate) on 28 August. These values are commensurate with those obtained in other SCAQS simulations. The predicted mean sulfate$_{2.5}$ and sulfate$_{10}$ concentrations are 9.5 and 10.2 µg m$^{-3}$ on 27 August and 6.3 and 6.8 µg m$^{-3}$ on 28 August. The corresponding observed values are 6.5 and 7.9 µg m$^{-3}$ on 27 August and 8.8 and 10.6 µg m$^{-3}$ on 28 August. In addition to direct emissions within the basin, particulate sulfate concentrations are also affected by the upwind boundary conditions and the formation of H$_2$SO$_4$ via SO$_2$ oxidation. Large uncertainties exist in the particulate sulfate emissions and the upwind sulfate boundary concentrations, which may contribute to the moderate overprediction on 27 August and underprediction on 28 August in particulate sulfate concentrations. Large uncertainties exist in the particulate sulfate emissions and the upwind sulfate boundary concentrations, which may contribute to the moderate overprediction on 27 August and underprediction on 28 August in both sulfate$_{2.5}$ and sulfate$_{10}$. Upwind boundary concentrations of 2.1 and 2.6 µg m$^{-3}$ were used for sulfate$_{2.5}$ and sulfate$_{10}$, respectively. The prevailing surface wind for this episode was westerly and blew from the western coast to the inland mountains throughout the basin [Lu et al., 1997]. The 8 PM sites are either located near the coast (e.g., HAWT) or inland along the trajectory of the prevailing wind, the upwind boundary conditions were found to contribute significantly (up to 35% contribution) to sulfate formation at those sites. The oxidation of SO$_2$ is slow in the absence of clouds but the gas-phase SO$_2$ oxidation produces up to 1.0–1.5 ppb (4–6 µg m$^{-3}$) of H$_2$SO$_4$ at the 8 PM sampling sites, which may contribute up to 75% and 33% of hourly and 24-hour average concentrations of sulfate$_{2.5}$, respectively, for this episode. The moderate overprediction on 27 August and underprediction on 28 August in particulate sulfate concentrations may therefore also result from the gas-phase formation of H$_2$SO$_4$ via SO$_2$ oxidation by OH radicals. OH is primarily produced through the photolytic reaction of O$_3$ and subsequent hydrolysis reaction of O($^1$D) with H$_2$O. The predicted O$_3$ and water vapor concentrations on 28 August are lower than those on 27 August at most sites in the northern and eastern portions of the SCAQS domain (e.g., BURK, CELA, RIVR), resulting in relatively high total gas-phase oxidizing capacity (i.e., higher OH levels) on 27 August and relatively lower oxidizing capacity (i.e., lower OH levels) on 28 August. At coastal sites (e.g., HAWT and
LBCC) and one northern site (i.e., AZUS), the OH levels on 28 August are similar to those on 27 August, but the predicted SO2 mixing ratios are lower than those on 27 August due possibly to some biases in other predicted meteorological variables such as wind fields that affect diffusion and transport of species. This could also contribute to lower H2SO4 formation on 28 August at those sites.

Two techniques were used to measure particulate nitrate during SCAQS. The Teflon filter method was used to measure both nitrate2.5 and nitrate10, and the denuder difference method was used to measure nitrate2.5 only. Large discrepancies were found between the observed nitrate2.5 concentrations obtained with these two techniques, and the denuder difference method is believed to be more accurate than the Teflon filter method [Hering et al., 1997]. The observed nitrate10 concentrations obtained with the Teflon filter were sometimes even lower than the nitrate2.5 concentrations measured with the denuder difference method, suggesting that the Teflon filter method may underestimate nitrate10 mass. Therefore the nitrate2.5 measurements obtained with the denuder difference method were used here. The predicted mean nitrate2.5 and nitrate10 concentrations are 10 and 11.6 μg m⁻³ on 27 August and 7.5 and 8.4 μg m⁻³ on 28 August. The corresponding observed values are 13.2 and 12.1 μg m⁻³ on 27 August and 16.6 and 13.5 μg m⁻³ on 28 August. The MNGE and MNB in predicted nitrate2.5 concentrations are 29% and 25% on 27 August and 60% and 51% on 28 August. The corresponding values in predicted nitrate10 concentrations are 25% and −1% on 27 August and 59%...
and −27% on 28 August. While the model tends to overpredict nitrate mass at night, it underpredicts nitrate mass during the day, resulting in an underprediction in 24-hour average nitrate$_{2.5}$ concentrations on both days and in nitrate$_{10}$ concentrations on 28 August. For the simulation with 8 size sections, nitrate$_{2.5}$ concentrations are somewhat overpredicted (instead of underpredicted) at HAWT and LBCC on 27 August (e.g., see Figure 3), resulting in a net overprediction, with MNGE and MNB of 44% and 10% for nitrate$_{2.5}$ and 47% and 27% for nitrate$_{10}$ on 27 August. The statistics on 28 August with 8 size sections are slightly improved for nitrate$_{2.5}$ with MNGE and MNB of 58% and −47% but slightly worse for nitrate$_{10}$ with MNGE and MNB of 62% and −29%.

[75] The particulate nitrate can be formed through dissolution or condensation of HNO$_3$ on the surface of particles and the heterogeneous reactions of nitrogen species such as N$_2$O$_5$ and NO$_3$ on the surface of wet particles. Ammonium nitrate (NH$_4$NO$_3$) and sodium nitrate (NaNO$_3$) are the most common nitrate salts. The formation of particulate nitrate can be limited by the abundance of HNO$_3$ (i.e., NO$_2$-limited), NH$_3$ (i.e., NH$_3$-limited) or ambient water vapor (i.e., H$_2$O-limited), depending on the chemical and meteorological conditions at a specific location. Therefore, the accuracy in nitrate predictions depends not only on the accuracy of meteorology and emissions of precursors (e.g., NO$_x$ and NH$_3$) but also on the accuracy of the gas- and aqueous-phase chemistry as well as gas-to-particle conversion processes simulated.

[76] Few clouds were present during this SCAQS episode; therefore HNO$_3$ formation was governed primarily by gas-phase reactions and to a lesser extent heterogeneous reactions. Gas-to-particle conversion was simulated with the CMU hybrid approach in the base simulations, in which ISORROPIA is used to calculate the particulate-phase concentrations. ISORROPIA has been shown to produce better fine NH$_4$ and NO$_3$ predictions than SEQUILIB [Nenes et al., 1999] (SEQUILIB was used in the simulations of SCAQS 24–26 June episode of Lurmann et al. [1997]). In addition, we tested the performance of ISORROPIA under 400 cases typical of the atmospheric conditions in the eastern and the western U.S. and found that the predictions of total PM and its chemical composition by ISORROPIA are comparable with those predicted by the more comprehensive equilibrium modules such as SCAPE2 and EQUISOLV II. The underpredictions in the NO$_3$ concentrations are therefore unlikely related to the model treatment of thermodynamic equilibrium in MADRID. Other work has also shown that given the correct total concentrations, the PM nitrate can be reproduced reasonably well by ISORROPIA, considering the measurement error and the bulk equilibrium assumption [Nenes et al., 1999; Ansari and Pandis, 1999].

[77] The formation of nitrate was unlikely NO$_2$-limited since there were generally abundant HNO$_3$ on both days at many locations. Underpredictions in NH$_3$ mixing ratios may contribute to the underpredictions of NO$_3$ concentrations at some western and central sites such as HAWT and ANAH on 28 August when the formation of NH$_4$NO$_3$ was NH$_3$-limited. The predicted low NH$_3$ mixing ratios during daytime may also contribute to low NO$_3$ formation at other inland locations, especially in the eastern domain. For example, the observed 24-hour average NH$_3$ mixing
ratio for 28 August was 1.4 ppb at HAWT and 15.1 ppb at RIVR. For comparison, the predicted 24-hour average NH₃ mixing ratio for the same day was 0.7 ppb at HAWT and 8.8 ppb at RIVR. Underestimation of the ambient RH during daytime is likely another factor that contributed to the under-prediction of NO₃. For example, the observed daytime RH values during 27–28 August range from 65–83% at HAWT and 28–60% at RIVR, while the predicted daytime RH values range from 44–75% at HAWT and 16–45% at RIVR. Therefore the predicted low RH limited to some extent the dissolution and condensation of HNO₃, resulting in lower NO₃ formation during daytime on both days.

[75] The biases in predicted meteorological fields may also have affected NH₃ predictions that depend on the accuracy of the precursor emissions and the predicted meteorological conditions. The predicted mean sub-2.5 μm and sub-10 μm ammonium (ammonium₂.₅ and ammonium₁₀) concentrations are 6.2 and 6.3 μg m⁻³ on 27 August, which are higher by 51% and 40% than the observed values of 4.1 and 4.5 μg m⁻³. The predicted ammonium₂.₅ and ammonium₁₀ concentrations are 4.6 and 4.7 μg m⁻³ on 28 August, which are lower by 18% and 19% than the observed values of 5.6 and 6.1 μg m⁻³. The MNGE and MNB at all 8 sites are 61% on 27 August and 44% and −11% on 28 August for predicted ammonium₂.₅ concentrations and 50% and 47% on 27 August and 44% and −19% on 28 August for predicted ammonium₁₀ concentrations. The NH₃ emission inventory used here was based on 1982 NH₃ emissions estimated by Cass and Gharib [1984]; therefore it may not accurately reflect the actual total NH₃ emissions and their regional distributions in 1987. Also, since the sulfate₂.₅ concentrations are dominated by (NH₄)₂SO₄, the moderate overprediction on 27 August and underprediction on 28 August in SO₄⁺ and NH₄⁺ are highly correlated. The underprediction in NH₄⁺ on 28 August also correlates with the underprediction in NO₃, since nitrate₂.₅ concentrations are dominated by NH₄NO₃ in the particulate phase. The overpredictions on 27 August and underpredictions on 28 August in NH₃ concentrations are most likely caused by biases in the predicted meteorological variables such as wind fields.

[76] The predicted mean EC₂.₅ and EC₁₀ mass are, respectively, 1.8 and 2.3 μg m⁻³ on 27 August and 1.7 and 2.2 μg m⁻³ on 28 August. These values are lower than the observed values of 2.5 and 3.3 μg m⁻³ on 27 August and 2.8 and 4.0 μg m⁻³ on 28 August. The statistics obtained with both 2 and 8 size sections are quite similar. As discussed above, large uncertainties exist in the PM emission inventories that affect the predicted EC concentrations.

[77] The predicted mean OM₂.₅ concentrations are 7.3 and 6.8 μg m⁻³ on 27 and 28 August; these values are lower by 22% and 40% than the observed values of 9.4 and 11.4 μg m⁻³. The predicted OM₁₀ concentrations are 13.9 and 13.0 μg m⁻³ on 27 and 28 August; the corresponding observed values are 13.6 and 16.9 μg m⁻³. The MNGE and MNB at all 8 sites are 38% and −14% on 27 August and 61% and −28% on 28 August for OM₂.₅ concentrations and 54% and 17% on 27 August and 65% and −7% on 28 August for OM₁₀ concentrations. The simulation with 8 size sections predicted similar MNGE and MNB, with 40% and −17% on 27 August and 61% and −29% on 28 August for OM₂.₅ concentrations and 53% and 14% on 27 August and 65% and −8% on 28 August for OM₁₀ concentrations. Those values are generally consistent with those obtained in earlier SCAQS simulations; except for slightly higher MNGE for OM₂.₅ and OM₁₀ predictions on 28 August and a slightly higher MNB for OM₁₀ predictions on 27 August.

[81] The observed OM comprises both primary and secondary organics, but the ambient monitoring data do not separate them. A direct evaluation of SOA predictions is thus not possible. The underpredictions in OM₂.₅ concentrations are likely due to underestimates in both the primary OC emissions and SOA formation. The SOA percentages in OM₁₀ predicted in the base simulations with MADRID 1 range from 1.7% to 10.9% at the sampling sites; these values are significantly lower than the estimated values of 10–75% for the SCAQS domain [Pandis et al., 1993], particularly on 28 August. In addition to the overpredicted wind speeds and vertical mixing that led to an underestimation of VOC precursor mixing ratios of SOA, several factors may contribute to this underestimate. First, large uncertainties exist in the emissions of SOA precursor VOC species in the SCAQS inventory as discussed previously. Different adjustments have been applied to the original CARB emission inventories for SCAQS 1987 in various modeling studies, resulting in quite different VOC emissions used in these studies. For example, the adjusted daily emissions of TOL and XYL used in the work of Jacobson et al. [1996b], Lu et al. [1997], Meng et al. [1998] and this work are 239 and 201 tons, 378 and 301 tons, 214 and 174 tons, and 314 and 226 tons, respectively. Most previous gas-phase simulations [e.g., Jacobson et al., 1996b; Harley et al., 1993, 1997] and this work underpredicted O₃, whereas Lu et al. [1997] predicted O₃ mixing ratios in excellent agreement with observations. This implies that the VOC precursor emissions used in all these work except Lu et al. [1997] may be somewhat underestimated. Emissions of other anthropogenic SOA precursors such as long-chain alkanes and alkenes, cresols and phenols have little impact on the predicted SOA in MADRID 1 because (1) those species are not considered as SOA precursors in MADRID 1, (2) they are represented by separate mechanism species in CBM-IV, thus not included in TOL and XYL. Second, uncertainties may also exist in the partition coefficients that were obtained at experimental temperatures of 301 K to 316 K and were corrected for temperature dependence at other ambient temperatures. The SOA formation is highly sensitive to the selection of the enthalpy of vaporization in equation (2) that was used to account for the temperature dependence. Although the temperature dependence of the partition coefficients was taken into account, the partition coefficients used in MADRID 1 are roughly lower by a factor of 2 than those calculated with the formulation of Pankov [1994] in MADRID 2. Third, the SOA module in MADRID 1 only includes two anthropogenic surrogate aromatic precursors, namely, TOL and XYL, while other anthropogenic VOC (e.g., long-chain alkanes and alkenes, cresol and phenols) may yield to SOA formation. Pandis et al. [1992] simulated SOA formation for the same SCAQS episode and found that the CBM-IV surrogate aromatic species (i.e., toluene, xylenes and alkylbenzenes) contributed to at most 55% of SOA formation. Biogenic SOA formation was found to be negligible in our simulation in the SCAQS domain due to low biogenic emissions. This is consistent
with the low biogenic SOA formation in another Los Angeles basin simulation [Griffin et al., 2002]. Considering additional VOC precursors of SOA will reduce the SOA (thus OM) underpredictions. A sensitivity simulation was thus conducted with a more detailed SOA module (i.e., MADRID 2), which includes 42 anthropogenic condensable VOC precursor species (see section 4.5.2).

Since the measured fine chloride concentrations were below the detection limit and highly uncertain, we thus focus our evaluation on sodium$_{10}$ and chloride$_{10}$. The predicted sodium$_{10}$ concentrations agree well with observations at all sites, with MNGE and MNB of 37 to 39% and −35 to −34%, respectively, whereas the statistics deteriorate somewhat for the predicted chloride$_{10}$ mass with MNGE and MNB of 49 to 80% and −34 to 3%, respectively. The MNGE and MNB predicted with 8 size sections are slightly worse for sodium$_{10}$ (40 to 49% and −41 to −38%, respectively), and significantly worse for chloride$_{10}$ (63 to 101% and 8 to 49%, respectively), which exemplifies the difficulty in simulating the volatility of chloride measurements and the difficulty in simulating volatile chlorine species.

4.5. Sensitivity Analyses

Since sensitivity studies to uncertainties in VOC and NO$_x$ emissions in the SCAQS inventory have been previously performed by many investigators [e.g., Pandis et al., 1992; Lu et al., 1997; Harley et al., 1997]. We focus here on the sensitivity of model predictions to several unique treatments and formulations in CMAQ-MADRID including heterogeneous reactions of HO$_2$, NO$_x$, O$_3$ and N$_2$O$_5$ on the surface of particles and droplets; two SOA modules that treat different numbers of condensable VOC precursor species with different partitioning mechanisms; two numerical schemes for simulating particle growth by condensation [or shrinkage by volatilization]; and three approaches for simulating gas/particle mass transfer. As shown below, these treatments/formulations can be major sources of uncertainties in 3-D O$_3$ and PM modeling.

4.5.1. Sensitivity to Heterogeneous Reactions

Figure 8 shows the predicted gas-phase mixing ratios of HO$_2$, H$_2$O$_2$, HNO$_3$ and O$_3$ at a coastal site (HAWT) with and without heterogeneous reactions. The heterogeneous reactions provided a source for H$_2$O$_2$ and HNO$_3$ throughout the simulation period, a sink for HO$_2$ during all the period except in the morning. The O$_3$ mixing ratios at HAWT decreased slightly (<4%) throughout the simulation period except between the sunrise and noon when there was an appreciable 0.3–7.8 ppb increase (by 0.4–16.7%). At this coastal site, the reactions that contribute to O$_3$ production include O($^1D$) + H$_2$O, NO$_2$ + hv, and CO + OH, CH$_4$ + OH, and HO$_2$ + NO. The reactions that contribute directly or indirectly to O$_3$ destruction include O($^1D$) → O($^3P$), OH + NO$_2$, O$_3$ + NO, C$_2$O$_3$ + NO$_2$, HO$_2$ + HO$_2$ + H$_2$O and HO$_2$ + HO$_2$. The heterogeneous reaction of NO$_2$ continued to be an important sink for NO$_2$ in the morning, resulting in lower NO$_2$ (by 4–7%) and lower NO (by 6–19%) mixing ratios.

While a lower NO$_2$ mixing ratio can lead to a lower O$_3$ in...
the morning, a lower NO and higher OH and HO₂ mixing ratios can lead to a higher O₃ mixing ratio by reducing the consumption rate of O₃ by NO and increasing the oxidation rate of CO and CH₄ by OH and that of NO by HO₂. The latter effect dominated in the morning, resulting in a net increase in O₃ mixing ratios at HAWT.

[85] The changes in the gas-phase species due to heterogeneous reactions caused corresponding changes in PM₁₀⁻, PM₂.₅⁻ and their components such as SO₄²⁻ and NO₃⁻ at HAWT, as shown in Figure 9. The 24-hour average sulfate₂.₅⁻ and sulfate₁₀⁻ concentrations with heterogeneous reactions were higher by 0.55 and 0.68 μg m⁻³ (by 6.7 and 7.5%, respectively) than those without heterogeneous reactions on 27 August at HAWT. This resulted from a higher aqueous-phase oxidation of SO₂ by H₂O₂ in cloud droplets between 1–6 a.m. and a higher gas-phase oxidation of SO₂ by OH between sunrise and noon. The changes in the 24-hour average nitrate₂.₅⁻ and nitrate₁₀⁻ concentrations at HAWT were opposite those of SO₄²⁻, namely, 1.25 and 1.21 μg m⁻³ (by 19.1 and 15.1%, respectively) lower on 27 August. The heterogeneous reactions of NO₂, NO₃ and N₂O₅ can either increase or decrease NO₃⁻ formation, depending on the physical and chemical conditions. These reactions may produce more HNO₃ that remain primarily in the gas-phase in the absence of NH₃ and would then condense on particles to form NH₄NO₃ due to the availability of NH₃. On the other hand, NO₃⁻ formation may be reduced in the presence of these heterogeneous reactions due to a lower gas-phase oxidation of NO₂ by OH (due to lower NO₂ and OH at night), a lower aqueous-phase formation rate (due to lower dissolved NO₂ and NO₃ radicals) between 1 a.m. and 6 a.m. and a lower oxidation of NO₂ by OH (due to lower NO₃⁻) between sunrise and noon. The latter effect dominated, resulting in lower 24-hour average NO₃⁻ concentrations on 27 August. As a net result of a higher SO₄²⁻ and a lower NO₃⁻ on 27 August, the 24-hour average PM₂.₅⁻ and PM₁₀⁻ mass concentrations were lower by 0.77 and 0.47 μg m⁻³, respectively, on 27 August.

[86] The effect of heterogeneous reactions on hourly concentrations of PM₂.₅⁻, PM₁₀⁻ and their components is more significant than that for their 24-hour average concentrations. For example, the hourly nitrate₁₀⁻ concentrations increased by up to 7.3 μg m⁻³ (~19%) at RIVR and 20 μg m⁻³ (~47%) at ANAH when heterogeneous reactions were included. The changes in the predicted hourly concentrations of sulfate₁₀⁻, ammonium₁₀⁻ and nitrate₁₀⁻ were up to 21%, 25% and a factor of 12, respectively, at the 8 PM

\[ \text{Figure 8. Predicted gas-phase mixing ratios of (a) HO₂, (b) H₂O₂, (c) HNO₃ and (d) O₃ at Hawthorne (HAWT) on 27–28 August 1987 with and without heterogeneous reactions.} \]

\[ \text{Figure 9. Observed and predicted (with and without heterogeneous reactions) 24-hour average mass concentrations of PM₂.₅⁻ and its chemical compositions at Hawthorne (HAWT) on 27 August 1987.} \]
sampling sites. Uncertainties are associated with the reaction probabilities of the heterogeneous reactions used in the simulations and with the simulated surface areas of particle and droplets.

### 4.5.2. Sensitivity to SOA Formulations

[87] A simulation was conducted with MADRID 2 of Pun et al. [2002], which represents the 42 condensable organic products of the CACM using 10 surrogate compounds: 5 hydrophobic organic compounds that partition into the particulate-phase and 5 surrogate hydrophilic organic compounds that dissolve into an aqueous phase to form SOA. Compared to the SOA formulation in MADRID 1, MADRID 2 treats 28 hydrophobic aromatic SOA precursors and their second- and third-generation products such as benzene-based aromatics with low and high volatility, naphthalene-based and aliphatic compounds. In addition, MADRID 2 treats 14 water-soluble organic compounds and their second- and third-generation products that are dissociative and nondissociative.

[88] The 24-hour average SOA concentrations predicted with MADRID 2 at the sampling sites increased from 0.06–0.6 to 1.6–6.1 μg m⁻³ (by a factor of 9 to 45). Compared to the results obtained with MADRID 1, the predicted SOA percentage in OM₁₀ increased from 2–11% to 15–62%.

For example, the 24-hour average SOA concentrations predicted by MADRID 2 at Claremont on 27 and 28 August are 6.1 and 3.2 μg m⁻³, which are a factor of 10 and 12 higher than those predicted by MADRID 1. With increased SOA formation, the simulation with MADRID 2 improved the performance statistics for OM₂₅ and OM₁₀ concentrations, turning an underprediction on both days into an overprediction on 27 August and a lower underprediction on 28 August. The predicted mean OM₂₅ concentrations averaged over 8 PM sampling sites increased from 7.3 to 11.3 μg m⁻³ on 27 August and 6.8 to 9.0 μg m⁻³ on 28 August. For comparison, the observed OM₂₅ concentrations were 9.4 and 11.4 μg m⁻³ on 27 and 28 August, respectively. The predicted mean error and bias with MADRID 2 are 46% and 34% on 27 August and 53% and −7% on 28 August for OM₂₅. The higher SOA predictions by MADRID 2 can be attributed to several factors: (1) more condensable organic products are treated in MADRID 2; (2) hydrophilic compounds, in addition to hydrophobic compounds, are treated in MADRID 2 but not in MADRID 1. Those hydrophilic condensable compounds formed primarily at night due to high RHs, they contributed about 5–15% of SOA formation on daily average; (3) larger (by roughly a factor of 2) partition coefficients employed in MADRID 2 than those in MADRID 1.

### 4.5.3. Sensitivity to Condensational Growth Algorithms

[89] A sensitivity simulation was conducted with 8 size sections and a simple finite difference approach (i.e., a semi-Lagrangian technique) for particle growth and the results were compared to the base case results with the moving-center scheme to demonstrate the model sensitivity to different particle growth schemes. In both simulations, the CMU hybrid approach was used to simulate gas/particle mass transfer. The simple finite difference scheme used here is based on that in the 1998 version of CIT [Meng et al., 1998], which was coupled with the CIT bulk equilibrium approach or the simple bulk equilibrium approach to approximate mass transfer in CIT. In this simple finite difference approach, the particle mass concentration after condensation or volatilization is adjusted using a factor that is calculated based on an incremental particle diameter corresponding to the incremental change in the particle mass (or volume) concentration. A fundamental difference between the moving-center scheme and the finite difference scheme is that the former predicts both PM mass and number concentrations, whereas the latter predicts only the PM mass concentrations and diagnoses the PM number concentrations from the predicted PM mass and the fixed PM mean diameters.

[90] The particle size distribution predicted by the moving-center and the finite difference schemes is shown at HAWT, LBCC, CELA and BURK on 27 August in Figure 10 and at CLAR and RIVR on 28 August in Figure 7 along with the observed average particle size distribution during the SCAQS summer sampling periods. The particle size distribution is obtained by summing those for all PM compositions (i.e., SO₄²⁻, NO₃⁻, NH₄⁺, Na⁺, Cl⁻, EC, OM and OI) in Figure 10 and all but Na⁺, Cl⁻ and OI in Figure 7. The finite difference scheme tends to predict a diffusive type of distribution for PM₂₅, with high concentrations in size sections 1 and 2 (0.0215–0.0464 μm and 0.0464–0.1 μm, respectively) (i.e., upstream diffusion) at both sites. The moving-center scheme predicts either one or two modes, with a PM₂₅ concentration peak in size sections 3 or 4 (0.1–0.215 μm and 0.215–0.464 μm). The PM mass fractions in the first size section (0.0215–0.0464 μm) predicted by the finite difference scheme account for 10 to 30% total PM₂₅ at the 8 PM sampling sites, which is significantly higher than those with the moving-center scheme (i.e., 0.02 to 0.2%). For comparison, the observed PM mass concentrations below 0.05 μm during the 1987 SCAQS study were negligible [e.g., Wall et al., 1988; John et al., 1990; Hering et al., 1997; Meng et al., 1998; Kleeman and Cass, 1998]. Accordingly, the finite difference scheme severely underpredicts a PM mass concentration of accumulation mode particles. In contrast, the moving-center scheme predicts a size distribution that is closer to the observed one in terms of the magnitudes and general shape of size-resolved PM composition, although the peak PM mass is somewhat underpredicted and is off the observed size for peak PM mass. The finite difference scheme moves mass from section to section during each timestep, causing significant numerical diffusion. Whereas, the moving-center scheme only moves mass from section to section when the section center grows out of the section, therefore minimizing the numerical diffusion.

[91] The numerical diffusion of the finite difference scheme artificially increases the 24-hour average PM₂₅ and PM₁₀ mass at many sites on both days. Compared to the base results, the 24-hour average mass concentrations of PM₂₅ predicted by the finite difference scheme increase significantly (by 8–30%) at all sites except the two downwind sites (CLAR and RIVR) on 27 August. The mass concentrations of PM₁₀ also increase accordingly (by 8–27%) at ANAH, AZUS, BURK, HAWT and LBCC on both days and CELA on 27 August. The increase in mass concentrations in the first two sections significantly increases the concentrations of particle number and the surface area available for heterogeneous reactions, which
in turn led to a significant increase in the heterogeneous reaction rates of HO2, NO2, NO3, and N2O5 and the mixing ratios of relevant gas-phase species including H2O2, NO2, O3, NO, HNO3, OH and HO2 at these locations. For example, at LBCC, the particle number concentrations and H2O2 mixing ratios increased by 46–80% and 11–65%, respectively. The HNO3 mixing ratios decreased by up to a factor of 3.9 at night, then increased by up to 16% between sunrise and noon. The O3 mixing ratios increased by up to 10% (13.1 ppb) in the morning but decreased in the afternoon by up to 10% (11.4 ppb).

4.5.4. Sensitivity to Gas-to-Particle Mass Transfer Approaches

Figure 11 compares the total particle mass size distribution predicted with the CMU hybrid, the CIT bulk and a simple bulk equilibrium approaches at HAWT and CELA on both days. A similar comparison is shown in Figure 7 at CLAR and RIVR on 28 August along with the observed average particle size distribution during the summer sampling periods. Although the simple bulk equilibrium approach gave mass concentrations for PM10 and its components similar to those of the CIT bulk equilibrium approach, it predicted a diffusive type of distribution, with lower mass concentrations in sections 2–5 between 0.1 and 1.0 μm, but much higher concentrations in larger sections, particularly in sections 6 and 8 (i.e., downstream diffusion). The total mass size distributions predicted by the CIT bulk equilibrium and the CMU hybrid approaches are quite similar (i.e., with similar amount of mass in each section and the peak occurring in the same section) at CLAR and ANAH (not shown) on 27 August and AZUS (not shown), BURK (not shown), CELA, CLAR (see Figure 7) and HAWT on 28 August. They both predicted a two-mode distribution at BURK (see Figure 10d) and HAWT on 27 August and a one-mode distribution at RIVR on 27–28 August (e.g., see Figure 7) and ANAH (not shown) and LBCC (not shown) on 28 August, although the mass concentrations in each section were somewhat different and the peak concentrations occurred in different sections. The CMU hybrid approach predicted a two-mode distribution at CELA and LBCC (not shown) with peaks in sections 3 and 6 (0.1–0.215 μm and 1.0–2.15 μm, respectively) on 27 August, whereas the CIT bulk equilibrium approach predicted a one-mode distribution at these locations with the peak in sections 4 and 3 at the two locations, respectively. Compared to the CIT bulk equilibrium approach, the CMU hybrid approach treats the gas/particle mass transfer explicitly for coarse particles, predicting higher mass concentrations in coarse sections, which in turn affects the mixing ratios of gas-phase species that are in equilibrium with the fine size sections. Although both approaches distribute the particle mass changes into fine size sections based on particle surface area weighting, the equations for the calculation of the weighting factors are somewhat different; in particular, the mass accommodation coefficient used in the CMU hybrid approach is 0.1, which is a factor of 10 higher than that used in the CIT bulk equilibrium approach. Therefore the CMU hybrid approach tends to give more weight to smaller size sections, with a peak concentration occurring most likely in smaller sections than the CIT bulk equilibrium approach at many sites such as BURK, CELA, HAWT, AZUS, CLAR and LBCC.

The CMU hybrid approach and the CIT bulk equilibrium approach predicted quite different mass size distributions of individual components, as shown in Figure 12. Both approaches predict a two-mode distribution for SO42− with similar total mass and the same peak sections (i.e., between 0.1–0.215 μm and 0.464–1.0 μm), but the highest sulfate mass concentration occurred between 0.1–0.215 μm for the CMU hybrid approach and between 0.464–1.0 μm for the CIT bulk equilibrium approach. Both approaches
predict similar total mass of NH$_4^+$ but with a more spread-out distribution over size sections in the CIT bulk equilibrium approach. Both approaches predict similar mass of Na$^+$ in section 8 (4.64–10 µm), but the CMU hybrid approach predict higher mass concentrations of Cl$^-$ than those of the CIT bulk equilibrium approach (0.19 versus 0.01 µg m$^{-3}$ at HAWT and 0.6 versus 0.13 µg m$^{-3}$ at CELA), which are in better agreement with the observed values of 1.0 and 0.7 µg m$^{-3}$ at HAWT and CELA, respectively.

The most significant differences between the two approaches lie in the total mass and size distribution of NO$_3^-$. The CMU hybrid approach predicts a much higher total NO$_3^-$ mass than the CIT bulk equilibrium approach (6.8 versus 1.9 µg m$^{-3}$ at HAWT and 13.9 versus 9.9 µg m$^{-3}$ at CELA, respectively). As a result of a larger total mass with a faster growth rate in fine sections and an explicit mass transfer in coarse sections, the CMU hybrid approach predicts higher mass concentrations for both fine and coarse NO$_3^-$ than the CIT bulk equilibrium approach. For example, the mass concentrations of coarse NO$_3^-$ predicted by the CMU hybrid and CIT bulk equilibrium approaches are 1.5 and 0.19 µg m$^{-3}$ at HAWT and 1.93 and 0.77 µg m$^{-3}$ at CELA, respectively.

**Figure 11.** The total particle mass size distribution predicted with the CMU hybrid, the CIT bulk equilibrium and the simple bulk equilibrium approaches at Hawthorne (HAWT) (a and b) and central Los Angeles (CELA) (c and d) on 27–28 August 1987.

**Figure 12.** The mass size distributions of Na$^+$, NH$_4^+$, SO$_4^{2-}$, NO$_3^-$ and Cl$^-$ predicted by the CMU hybrid and the CIT bulk equilibrium approaches at HAWT (a and b) and CELA (c and d) on 27 August 1987.
at CELA, respectively, on 27 August. For comparison, the observed mass concentrations of coarse NO$_3$ (nitrate$_{10-2.5}$) are 0.73 and 1.0 μg m$^{-3}$ at HAWT and CELA, respectively. Note that the observed nitrate$_{10-2.5}$ mass concentrations are obtained by subtracting the observed nitrate$_{2.5}$ mass from the observed nitrate$_{10}$ mass and they may be underestimated, due to mass losses occurring during nitrate$_{10}$ sampling with the Teflon filter method. Therefore the accuracy in nitrate$_{10}$ predicted by the CMU hybrid and CIT bulk equilibrium approaches can not be determined because of the uncertainties in the nitrate measurements. It is clear that the CIT bulk equilibrium approach tends to underpredict the mass concentrations of coarse nitrate and chloride at or near coastal sites; it is, however, applicable for ambient conditions with low concentrations of coarse sea salt and alkaline dust particles.

Despite the above differences, both the CMU hybrid and the CIT bulk equilibrium approaches can predict particle size distribution for PM$_{2.5}$ that is reasonably close to the observed size distribution as shown in Figure 7, considering the large uncertainties in meteorology and emissions inventories used, a relatively coarse size resolution of 8 size sections used and difficulties in accurately simulating particle size distribution due to many complex microphysical processes that govern PM chemical composition and size distribution. One limitation is that both the CIT bulk equilibrium and the CMU hybrid approaches employ equilibrium assumptions on gas/particle mass transfer (except for coarse particles in the CMU hybrid approach). The dynamic approach that explicitly treats mass transfer [e.g., Meng and Seinfeld, 1995; Meng et al., 1998; Jacobson, 1997a, 1997b; Sun and Wexler, 1998a, 1998b] should give more accurate results in theory than both simplified approaches but will require higher computational costs. Such an explicit mass transfer approach should be used when computational resources are available.

5. Conclusion

We have presented the development and initial application of a new 3-D air quality model for PM, CMAQ-MADRID. This model combines a state-of-the-science representation of the major processes that govern the chemical composition and size distribution of PM in the atmosphere with numerical robustness of the corresponding algorithms.

CMAQ-MADRID was applied to simulate an air pollution episode in the Los Angeles basin. Model performance for both O$_3$ and PM predictions was shown to be consistent with existing guidance. The model simulations reproduce the magnitude and the spatial and temporal variations of O$_3$ mixing ratios throughout the basin, but tend to overpredict daytime O$_3$ mixing ratios at a few sites in the western basin on 27 August and underpredict the O$_3$ mixing ratios at inland sites downwind on 28 August. The overpredictions or underpredictions are due, at least in part, to mispredictions in the meteorological inputs such as wind speeds and vertical mixing. The evolution of the chemical composition of PM from the coastal areas to the inland areas was well reproduced except that EC and OM were underpredicted and particulate nitrate formation was sometimes underpredicted. The underprediction in EC may be due to uncertainties in meteorological inputs, EC emissions and the size distribution of emitted EC. The underpredictions in OM may be due to uncertainties in meteorological inputs, emissions of condensable VOC species and primary OC emissions, the partition coefficient used for condensable OC precursors, as well as an incomplete inclusion of the condensable VOC species that contribute to SOA formation. The underpredictions in nitrate are due mainly to overpredictions in vertical mixing, underpredictions in RH and uncertainties in the emissions of primary pollutants such as VOC, NO$_x$, and NH$_3$. These results indicate the importance of accurate meteorological inputs, emissions (both gases and primary PM species), characterization of emitted PM size distribution as well as model representation of various atmospheric processes for PM air quality simulations.

The sensitivity of model predictions was evaluated with respect to several major areas of uncertainties in PM modeling including the treatment of heterogeneous reactions, and different modules/algorithms for SOA formation, particle growth due to condensation (or shrinkage due to volatilization) and gas/particle mass transfer. The predicted gas-phase species mixing ratios and particulate-phase species concentrations are sensitive to heterogeneous reactions of HO$_2$, NO$_2$, NO$_3$ and N$_2$O$_5$ on particles and droplets. Heterogeneous reactions taking place at the surface of particles and droplets are shown to potentially affect hourly O$_3$ mixing ratios by up to 17% and hourly concentrations of sulfate$_{10}$, ammonium$_{10}$, and nitrate$_{10}$ by up to 21%, 25%, and a factor of 12, respectively. Such an effect may even cause changes in 24-hour average PM$_{2.5}$, PM$_{10}$ and their compositions (e.g., up to 3%, 7% and 19% for 24-hour average PM$_{2.5}$, sulfate$_{2.5}$ and nitrate$_{2.5}$ concentrations). Uncertainties are associated with the reaction probabilities of the simulated heterogeneous reactions and the particle and droplet surface areas used.

The treatment of SOA formation was investigated with a sensitivity simulation. A SOA module with a mechanistic representation (i.e., MADRID 2) provides OM concentrations that are in better agreement with observations than a SOA module that is based on smog chamber data (i.e., MADRID 1). The treatment of SOA formation is still an area of ongoing research and large uncertainties currently exist for this PM component in all existing air quality models. The yields of condensable organic compounds and the temperature-dependent gas/particle partition coefficients of those compounds are still major sources of uncertainties that will require additional fundamental studies to improve our ability to predict SOA concentrations.

The predicted fine and total particle mass and their size distributions are sensitive to the numerical algorithms for simulation of particle growth and gas/particle mass transfer. The moving-center scheme for particle growth is shown to predict more accurate particle size distributions than other semi-Lagrangian and Lagrangian schemes such as the finite difference scheme, which typically cause an upstream numerical diffusion. For gas/particle mass transfer, a realistic particle size distribution can be predicted with the CMU hybrid approach under most ambient conditions and with the CIT bulk equilibrium approach under conditions with negligible reactive coarse particles (e.g., sea salt and dust). In contrast, the simple bulk equilibrium approach tends to cause a downstream numerical diffusion in the predicted particle size distribution. These sensitivity simulations demonstrate a need for careful selection of numerical
algorithms for simulating PM thermodynamics and dynamics for 3-D air quality models that are based on a sectional representation of the particle size distribution.

[10] Accurate model inputs (e.g., emissions and meteorology), realistic representations of various atmospheric processes (e.g., SOA formation), appropriate numerical algorithms for PM dynamics (e.g., condensational growth and gas/particle mass transfer) and fine particle size resolution (e.g., ≥8 size sections) will greatly reduce the major model uncertainties and thus improve the accuracy for 3-D PM modeling.

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