

“Insights from the climate modeling community with special attention to High Resolution Modeling”

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USGS
science for a changing world



Author Lists

- Diffenbaugh, N.S. and M. Scherer, in prep
- Diffenbaugh, N.S., M. Ashfaq and M. Scherer, *JGR*, in review
- Giorgi, F., E. Coppola, E.S. Im, N.S. Diffenbaugh, X. Gao and Y. Shi, *Journal of Climate*, in press
- Diffenbaugh, N.S. and M. Scherer, *Climatic Change Letters*, 2011
- Diffenbaugh, N.S., M.A. White, G.V. Jones and M. Ashfaq, *Environmental Research Letters*, 2011
- Diffenbaugh, N.S., and M. Ashfaq, *GRL*, 2010
- Walker, M.D. and N.S. Diffenbaugh, *Climate Dynamics*, 2009
- Rauscher, S.A., J.S. Pal, N.S. Diffenbaugh and M. Benedetti, *Geophysical Research Letters*, 2008
- Diffenbaugh, N.S., C.H. Krupke, M.A. White and C.E. Alexander, *Environmental Research Letters*, 2008
- Diffenbaugh, N.S., J.S. Pal, F. Giorgi and X. Gao, *Geophysical Research Letters*, 2007
- White, M.A., N.S. Diffenbaugh, G.V. Jones, J.S. Pal and F. Giorgi, *PNAS*, 2006

Objective:

Motivate some climate-related issues that could pose challenges for the energy-land-water nexus and may (or may not) be integratable into the Integrated Assessment Model framework

Climate and the Energy-Land-Water Nexus

Climate change can influence the need for and availability of:

Energy

Land

Water

for a variety of sources and uses

Some Questions

- What processes determine the response of climate to changes in radiative forcing?
- How might natural and human systems be affected?
- How do climate processes and impacts respond to different interventions?

(“changes in radiative forcing” - e.g., changes in global greenhouse gas concentrations)

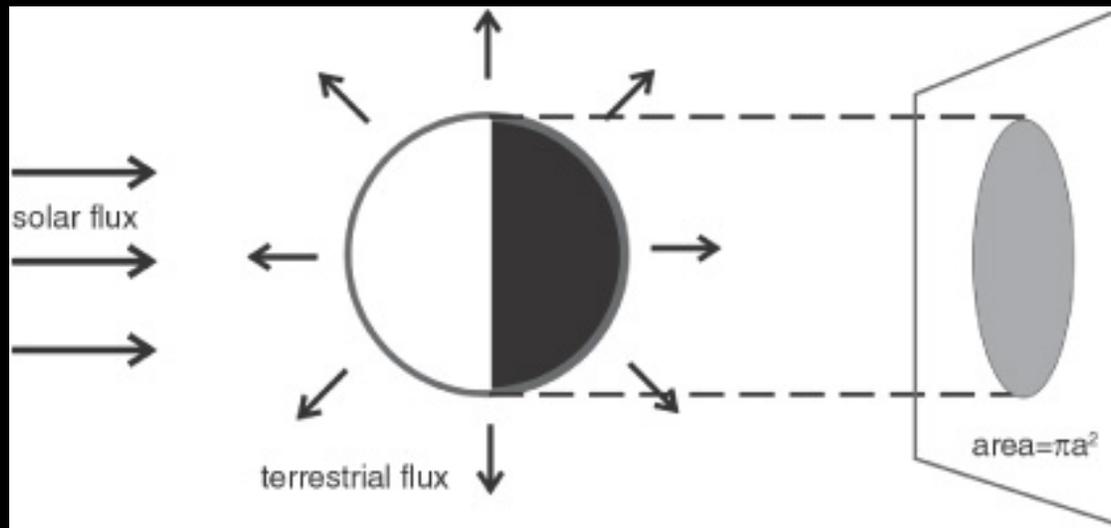


Figure 2.4: The spinning Earth is imagined to intercept solar energy over a disk of radius a and radiate terrestrial energy away isotropically from the sphere.

(Modified from Hartmann, 1994.)

From Marshall and Plumb, Elsevier, 2008

TABLE 2.1. Properties of some of the planets. S_0 is the solar constant at a distance r from the Sun, α_p is the planetary albedo, T_e is the emission temperature computed from Eq. 2.4, T_m is the measured emission temperature, and T_s is the global mean surface temperature. The rotation period, τ , is given in Earth days.

| | r 10^9 m | S_0 W m^{-2} | α_p | T_e K | T_m K | T_s K | τ Earth days |
|---------|-------------------------|----------------------------|------------|------------|------------|------------|----------------------|
| Venus | 108 | 2632 | 0.77 | 227 | 230 | 760 | 243 |
| Earth | 150 | 1367 | 0.30 | 255 | 250 | 288 | 1.00 |
| Mars | 228 | 589 | 0.24 | 211 | 220 | 230 | 1.03 |
| Jupiter | 780 | 51 | 0.51 | 103 | 130 | 134 | 0.41 |

$$T_e = [S_0(1-\alpha_p) / 4\sigma]^{1/4}$$

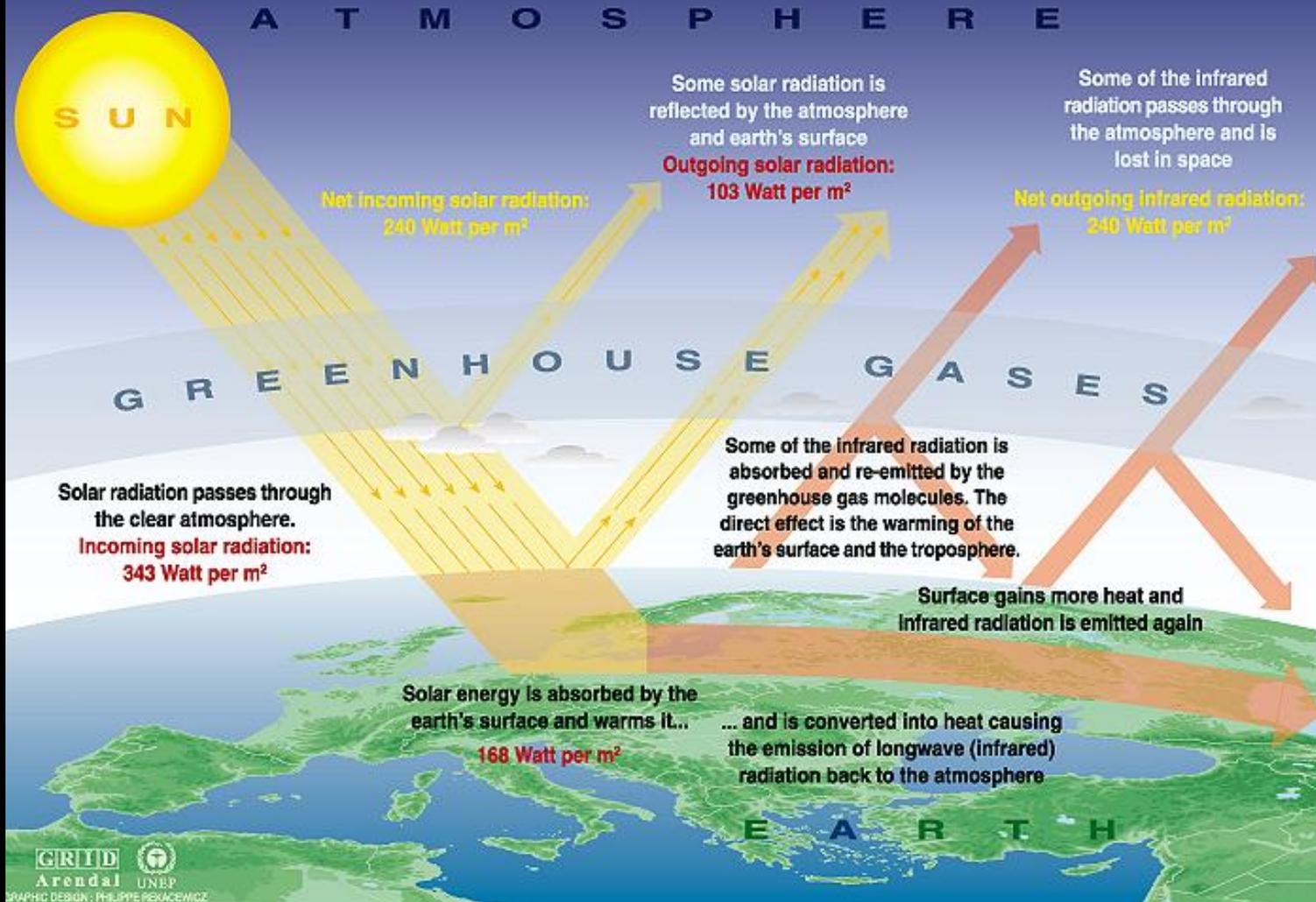
T_e = emission temp

S_0 = solar "constant"

α_p = planetary albedo

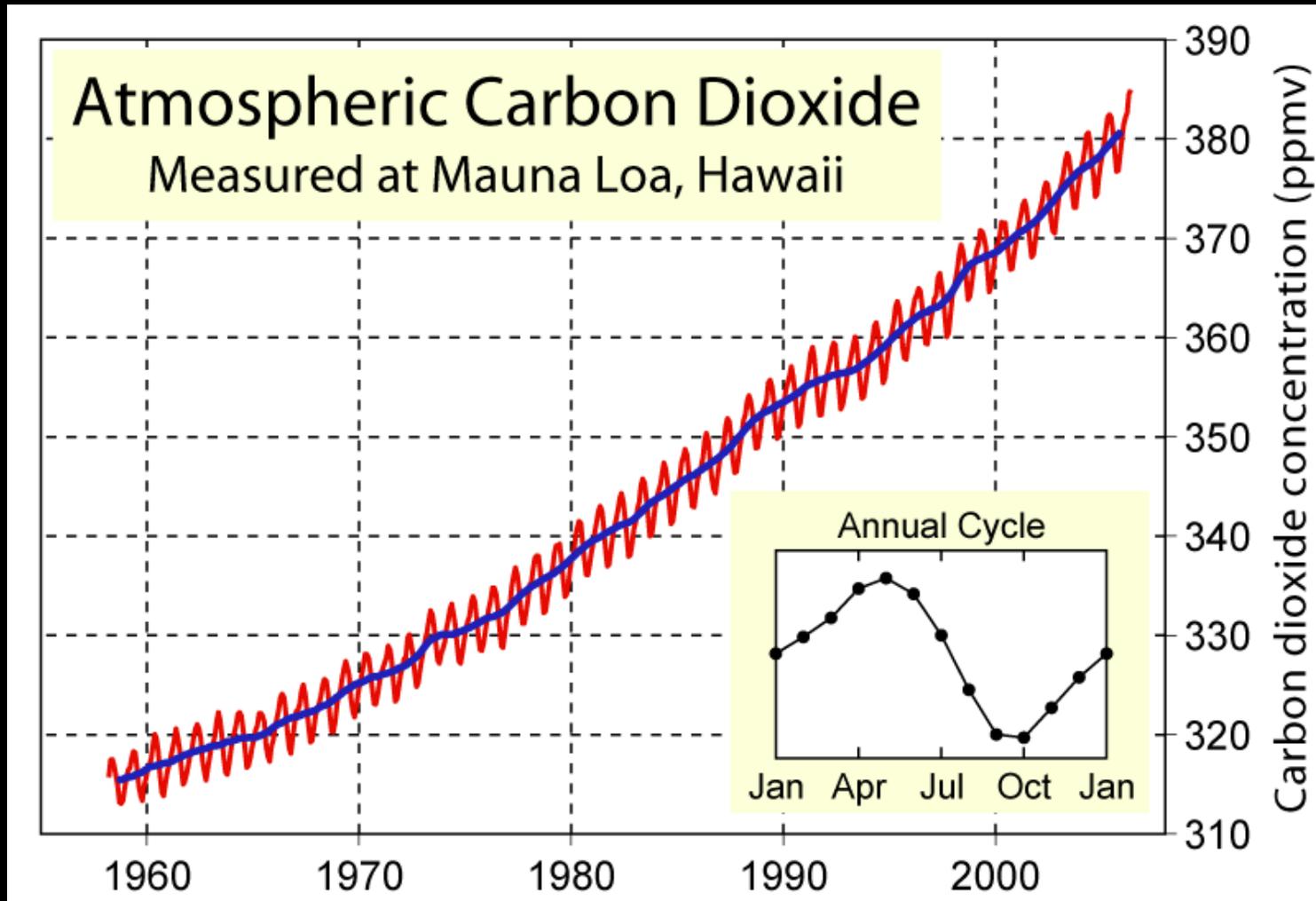
σ = Stefan-Boltzman constant

The Greenhouse effect



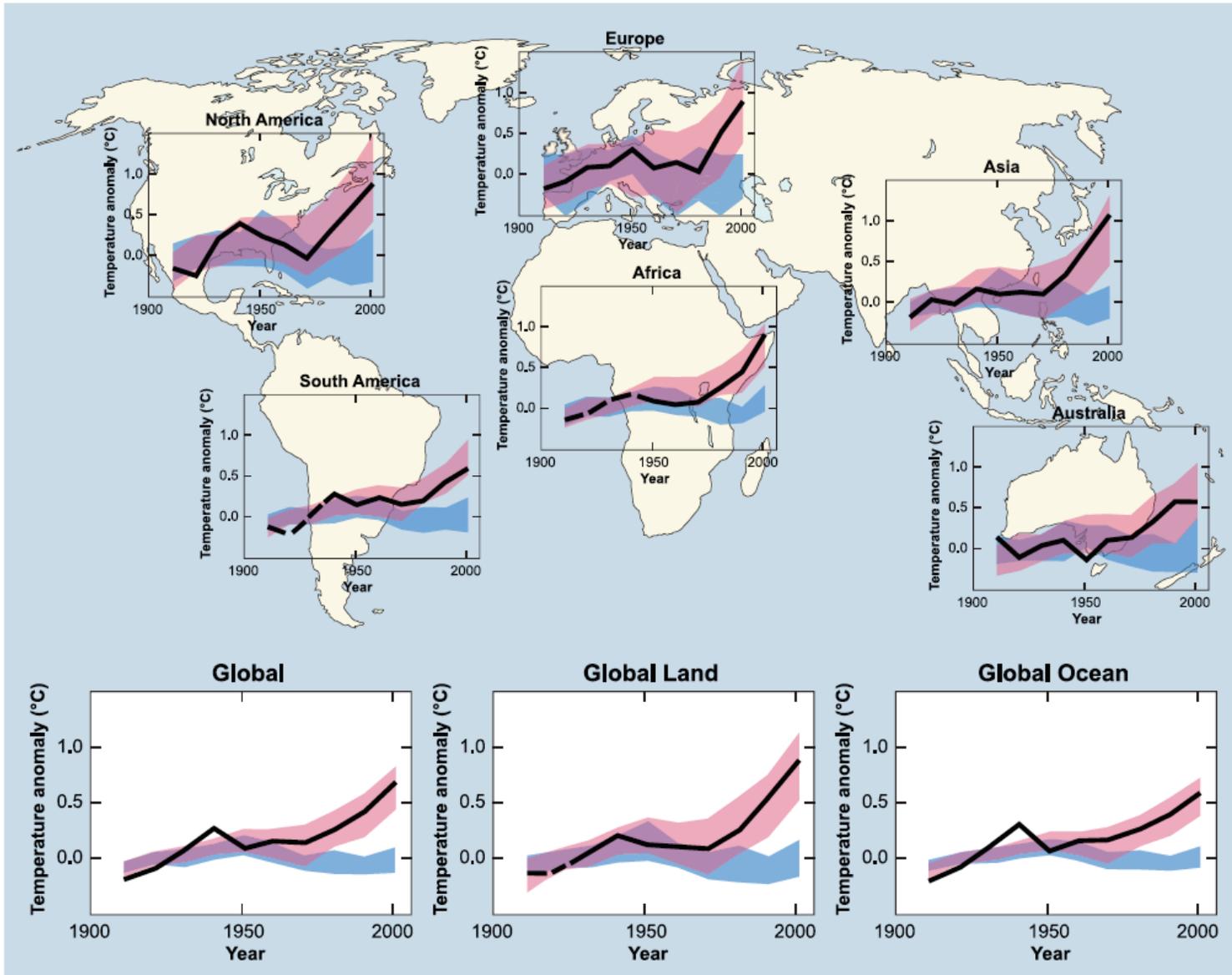
Sources: Okanagan university college in Canada, Department of geography, University of Oxford, school of geography; United States Environmental Protection Agency (EPA), Washington; Climate change 1995, The science of climate change, contribution of working group 1 to the second assessment report of the intergovernmental panel on climate change, UNEP and WMO, Cambridge university press, 1996.

Keeling Curve



Global Warming Art

Global and continental temperature change

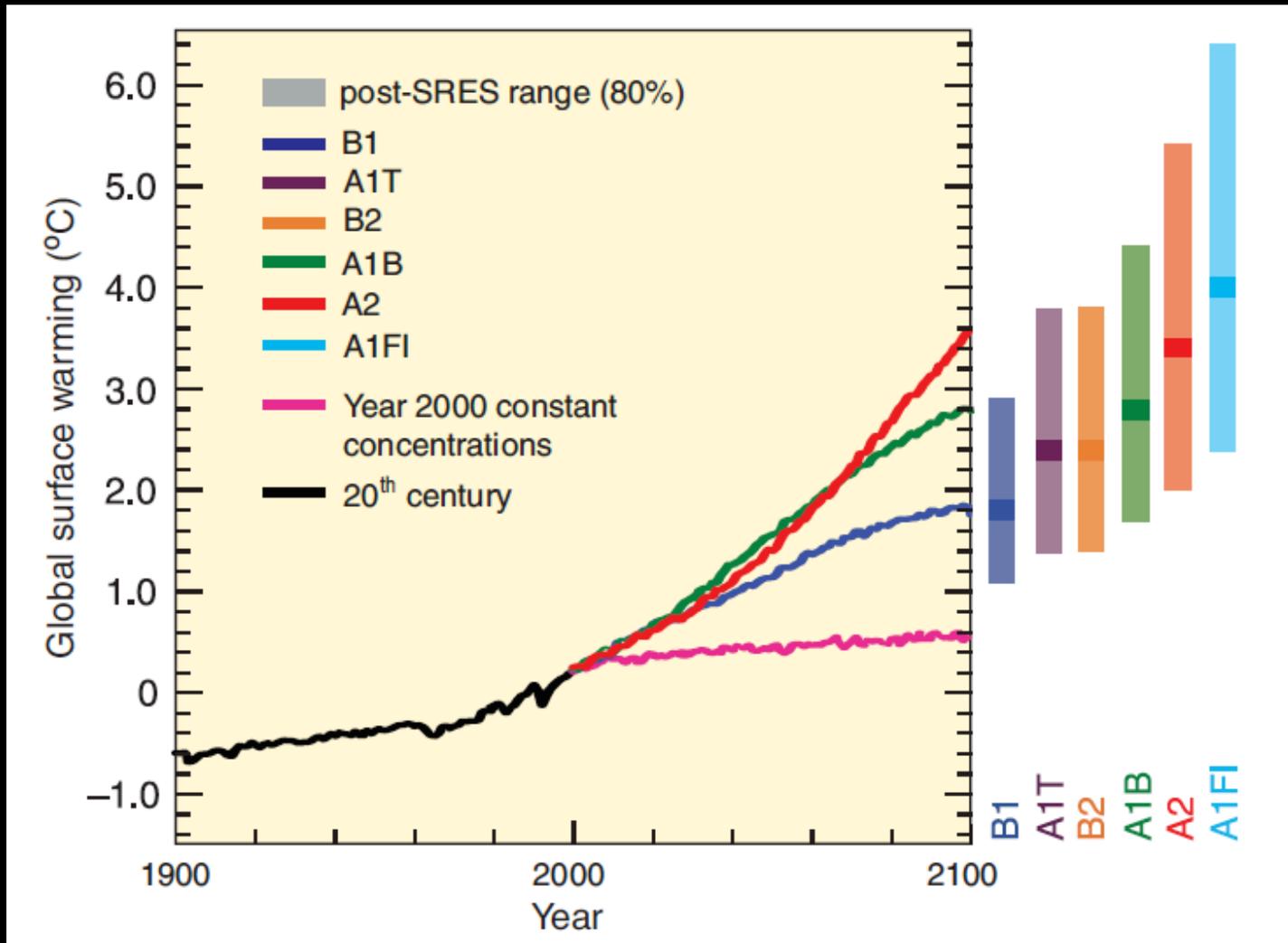


models using only natural forcings
models using both natural and anthropogenic forcings

observations

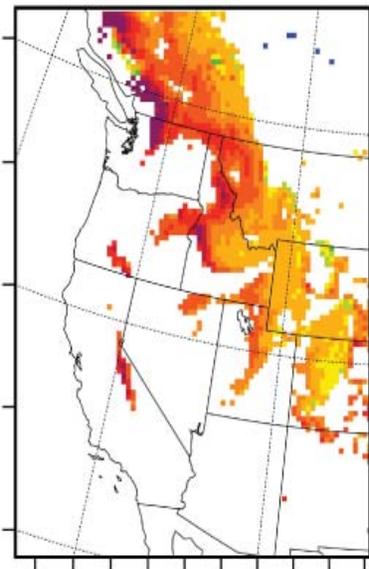
IPCC, 2007

What might happen if climate changes?

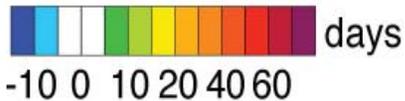


A2: 2071-2099 minus 1961-1989

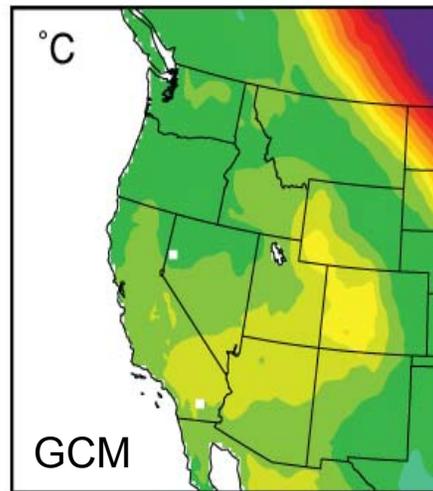
Snow-Dominated
Runoff



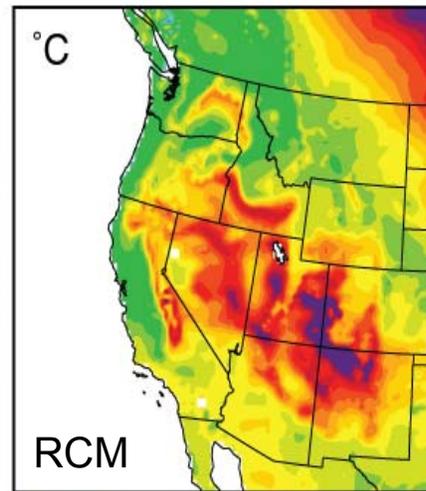
Late 21st Century Change



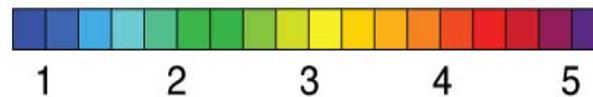
Winter Temperature Change



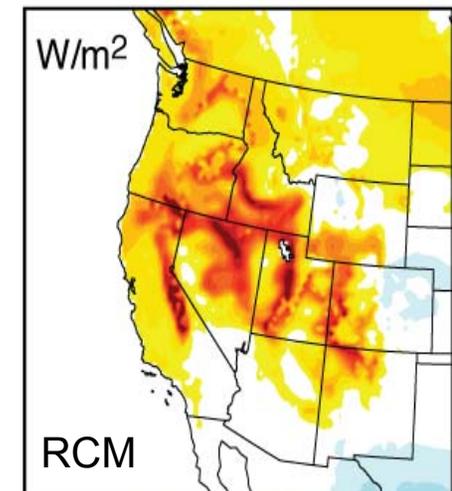
GCM



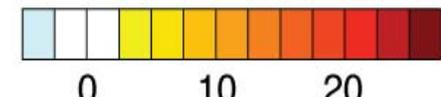
RCM



Absorbed Solar
Radiation

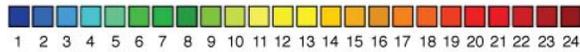
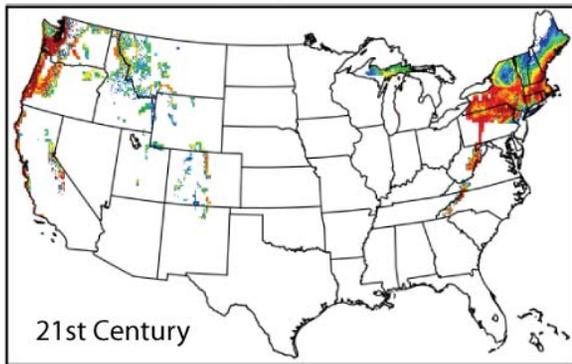
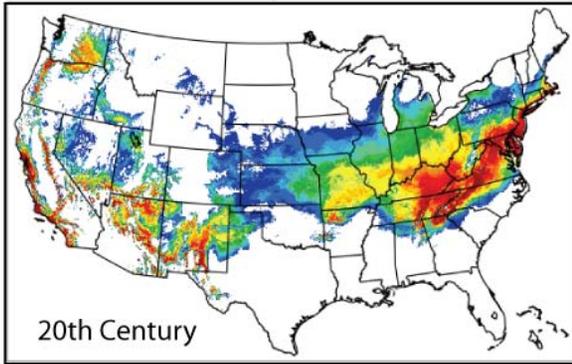


RCM

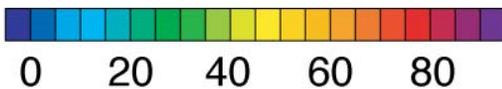
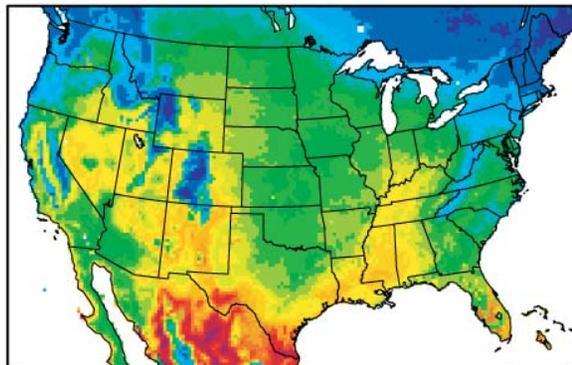


=> The climate change in the high-resolution projection is twice as large as in the GCM because of improved representation of the physical climate system.

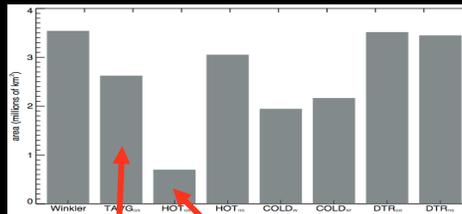
Premium Wine Production



Change in Growing Season Hot Days



Suitable Area



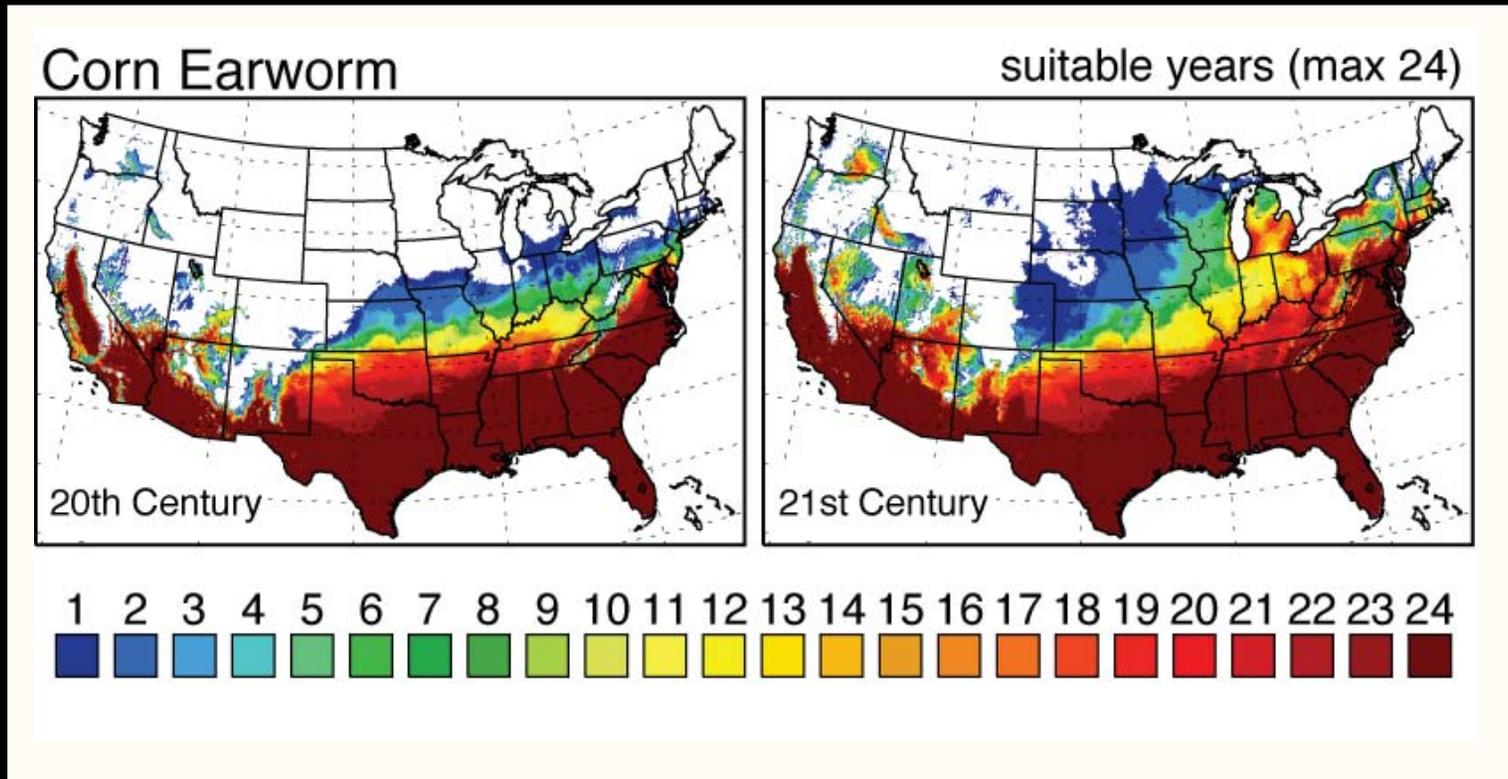
Mean
Seasonal
Temperature

Extreme
Daily
Heat

- **Response** of wine production driven by **response** of severe heat

- **Response** of severe heat varies spatially

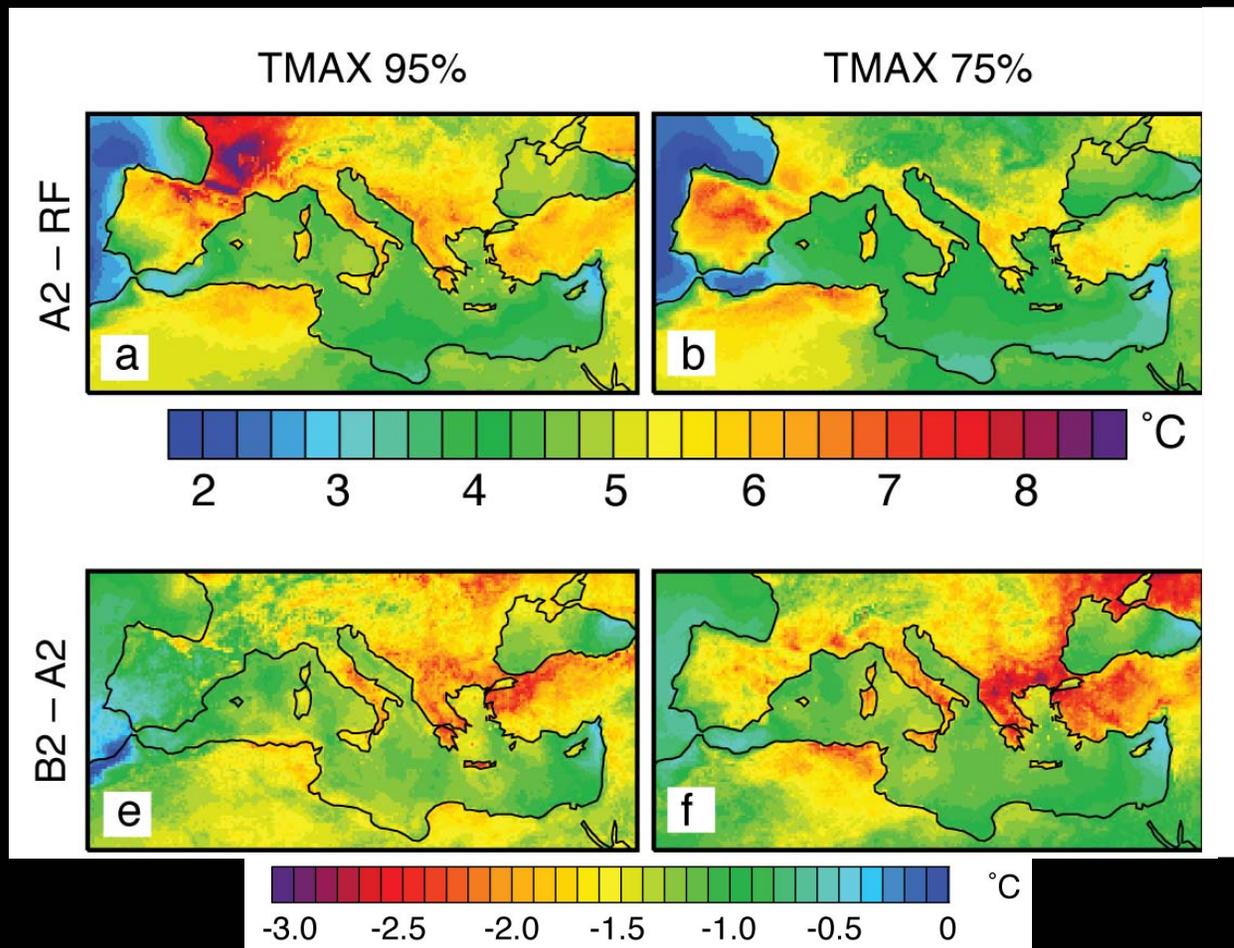
Agricultural Pest Prevalence



- Migratory, cosmopolitan, with strong pesticide resistance
- Expansion of range due to relaxation of severe cold limitation

Diffenbaugh et al., ERL, 2008

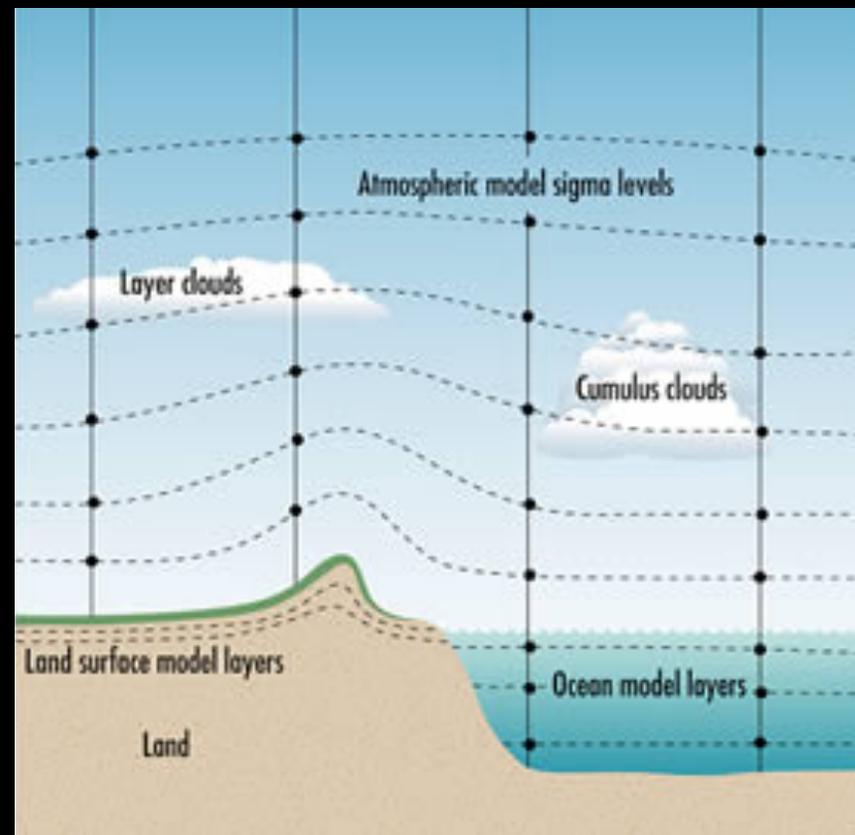
“Deceleration Effect”



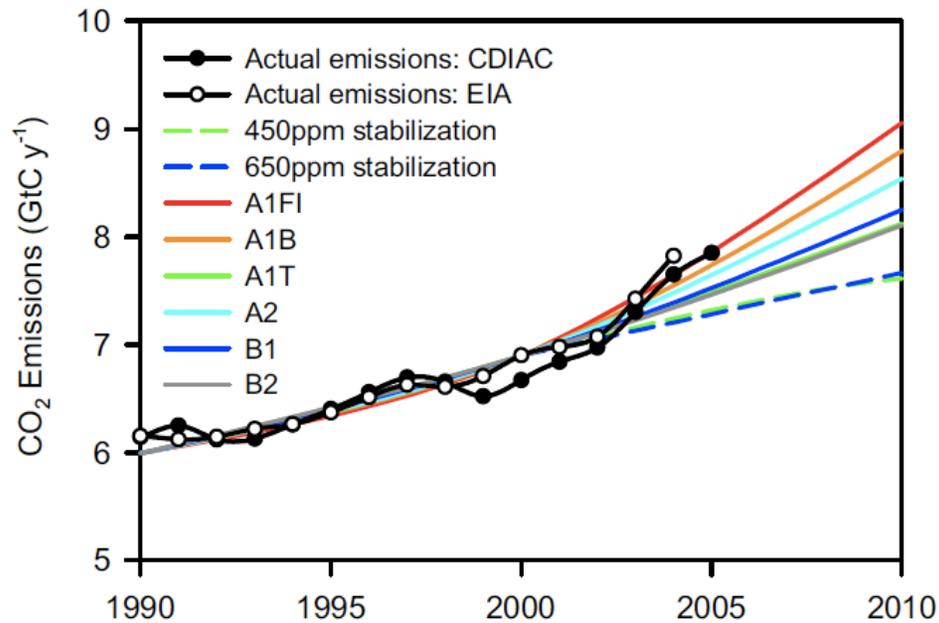
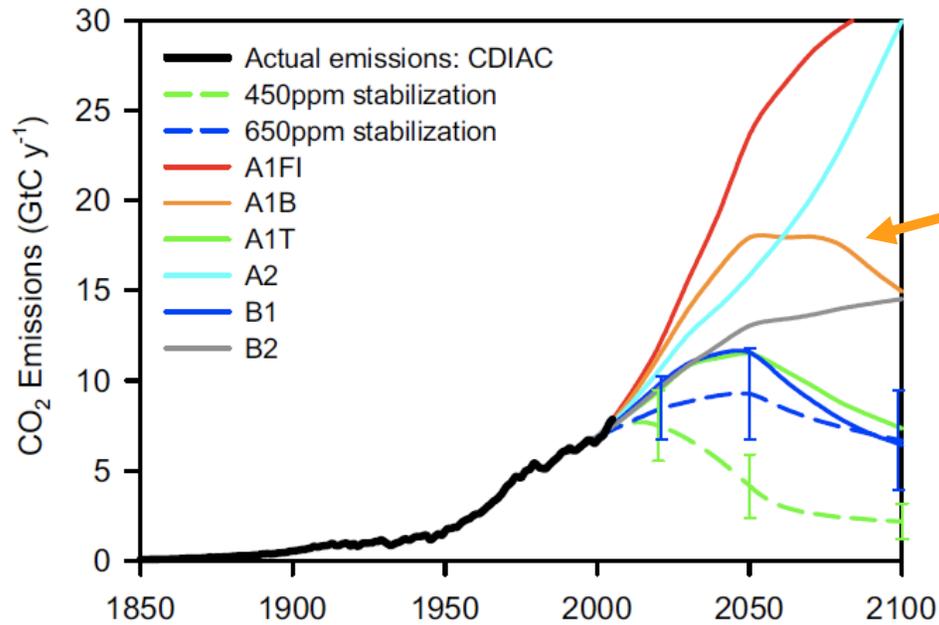
- Reduced mitigation “savings” due to surface moisture feedback
- Value of GHG mitigation is non-uniform, particularly for critical thresholds

**How Quickly Can Severe
Climate Change Emerge?**

General Circulation Model (“GCM”)



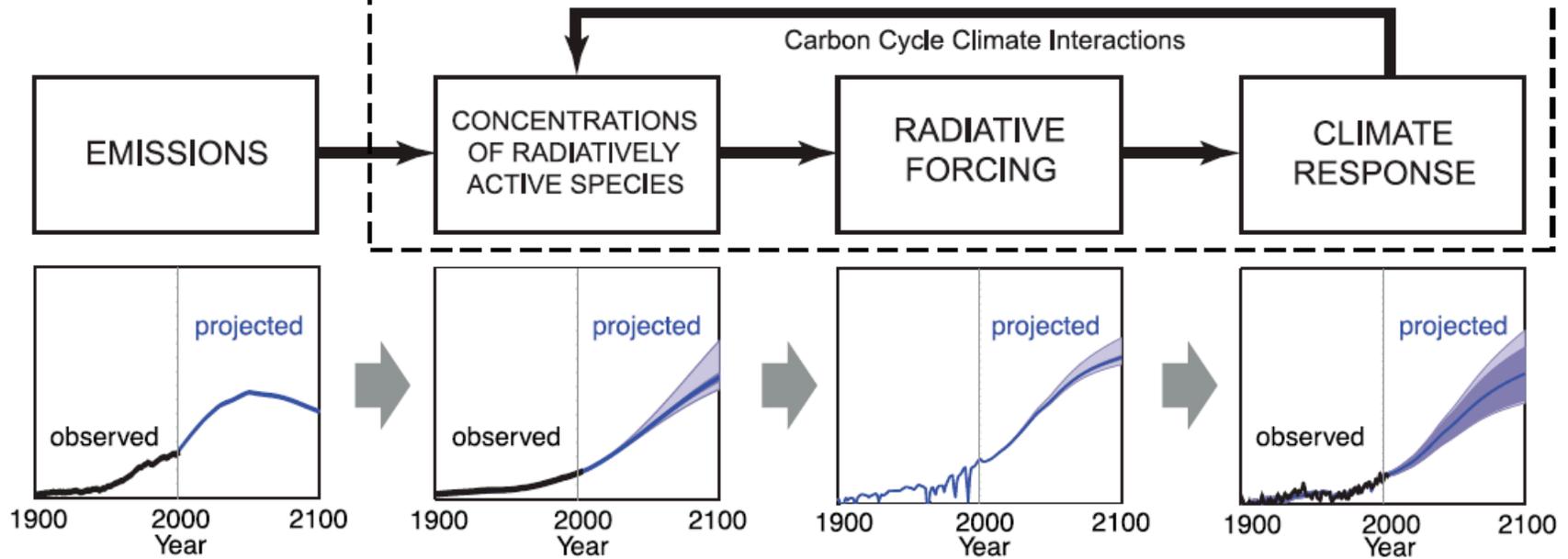
Laws of Physics!!!



A1B

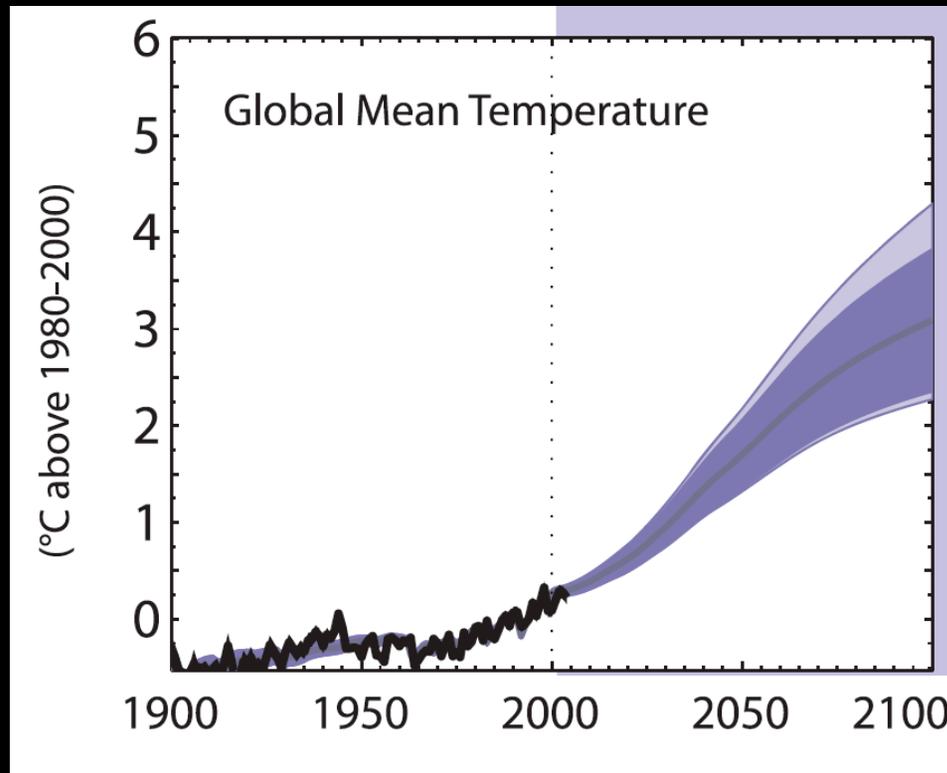
IPCC SRES Scenarios

Comprehensive Climate Model



21st Century Global Warming

Many Models (A1B)



IPCC, 2007

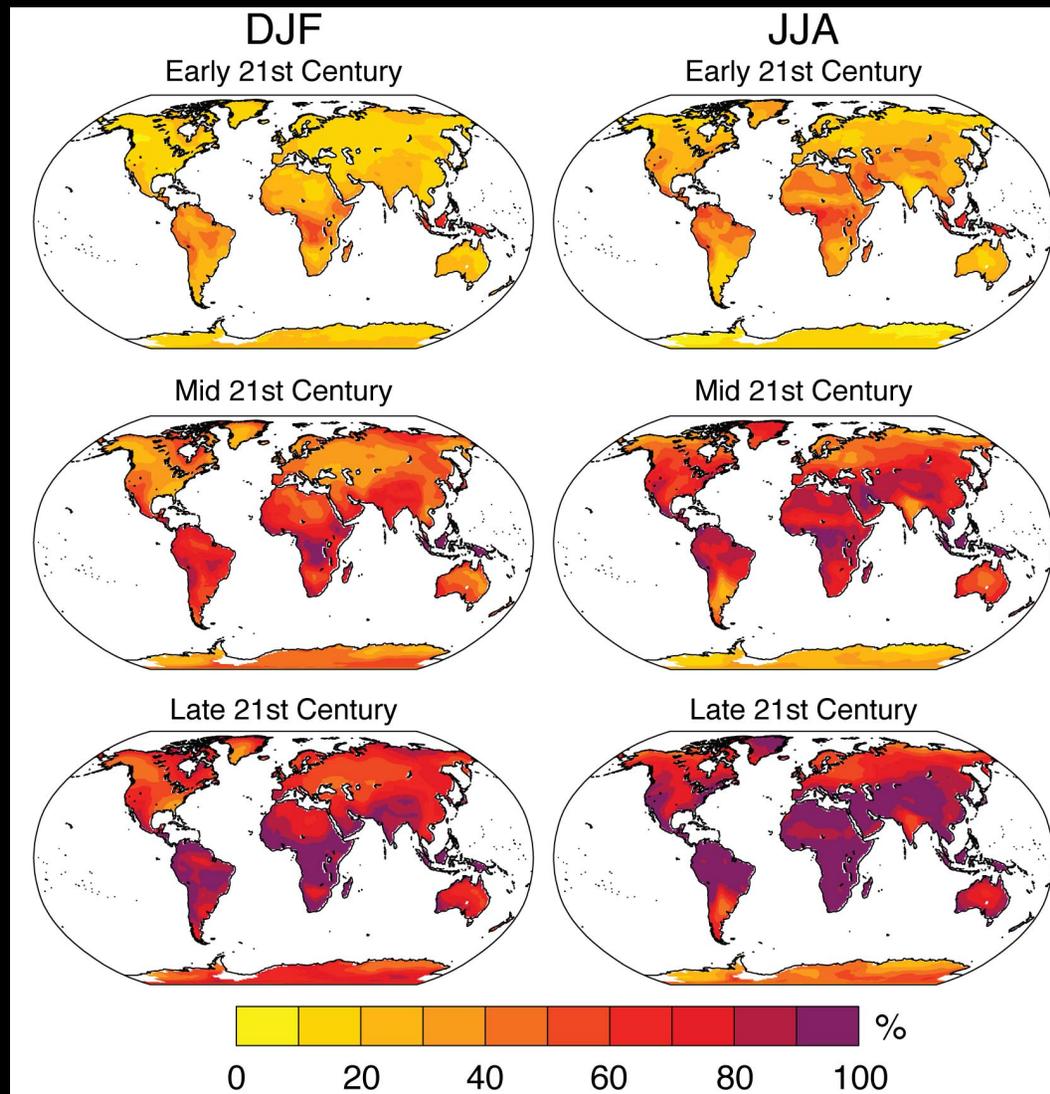
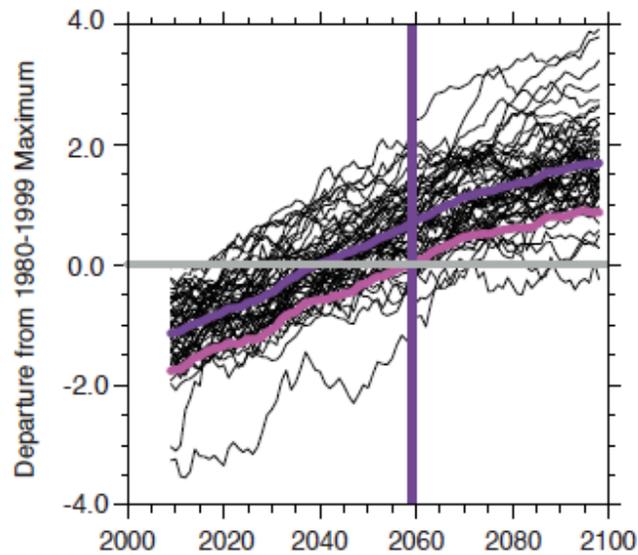


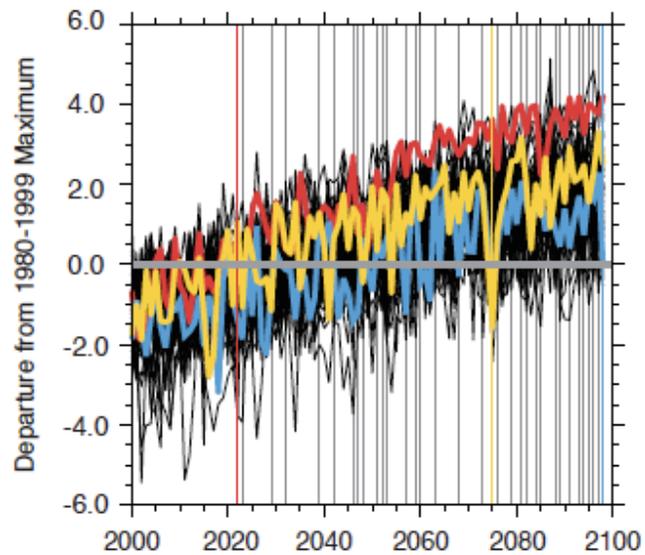
Figure 1. The percentage of seasons in the CMIP3 A1B ensemble for which the surface air temperature exceeds the the warmest season of the 1980-1999 period. The early 20th century period is 2010-2039, the mid 20th century period is 2040-2069, and the late 20th century period is 2070-2098.

Time of Emergence



- Individual Realizations
- Mean of Realizations
- Difference Between Mean and Standard Deviation of Realizations

Permanent Exceedence



- Individual Realizations
- First Permanent Exceedence
- Median Permanent Exceedence
- Last Permanent Exceedence

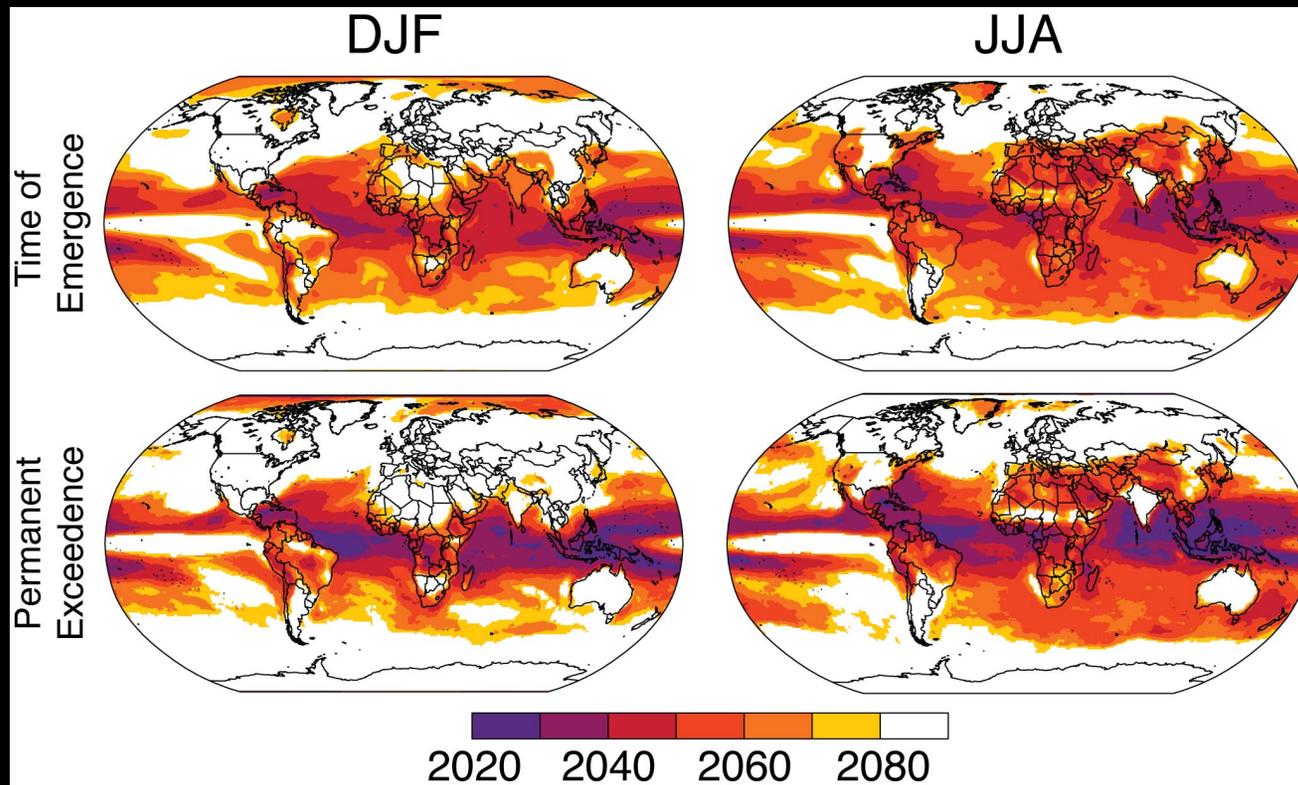


Figure 2. Decade of emergence of extreme surface air temperature. The top row shows the time of emergence (TOE) of the ensemble signal, calculated as the decade in which the ensemble mean seasonal temperature difference from the 1980-1999 maximum becomes permanently greater than the spread (one standard deviation) between the individual member differences from the 1980-1999 maximum. The second row shows the decade of the last occurrence of a season that is cooler than the 1980-1999 maximum, calculated as the median of the values across the CMIP3 ensemble. We cannot confirm whether the exceedence is permanent beyond the end of the 21st century, and therefore eliminate dates after 2080. Further details of both metrics are given in the text, and in Fig. S1.

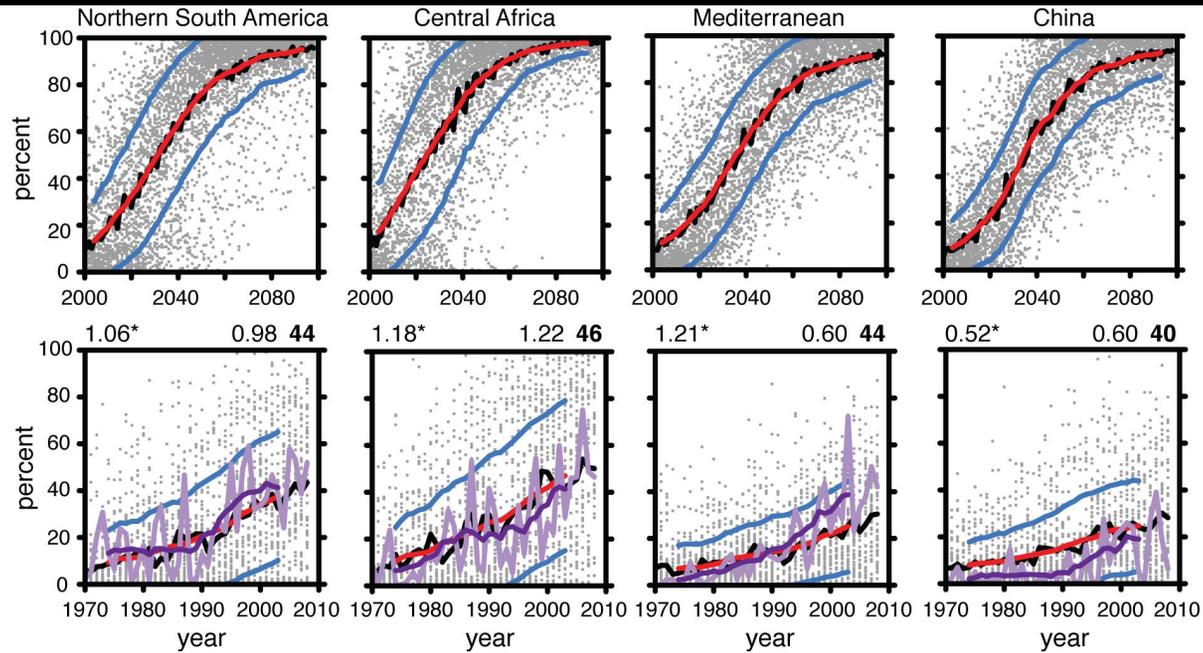


Figure 3. Time series of exceedence of historical temperature extreme. Top row: Time series of the fraction of land grid points with JJA surface air temperatures above the respective 1980-1999 maximum. Bottom row: Time series of the fraction of land grid points with JJA surface air temperatures above the respective 1952-1969 maximum. The grey points denote this fraction for each CMIP3 model realization. The black curves show the mean across all realizations. The red curves show the 10-year running mean of the mean across all realizations. The blue curves show the 10-year running mean of a one-standard-deviation range across the mean of all realizations. The light purple shows the observational timeseries. The dark purple curves show the 10-year running mean of the observational timeseries. The regions cover land areas in Northern South America (25S-3N, 82-33W), Central Africa (1S-13N, 18W-51E), the Mediterranean (28-50N, 12W-38E), and China (21-43N, 97-123E). The trend in the observational timeseries (% per year) is shown in the upper left of the bottom panels, with those trends that are statistically significant identified with a *. The mean of the trends in the CMIP3 realizations is shown in the upper right of the bottom panels, with the number of realizations that exhibit a statistically significant trend shown in bold.

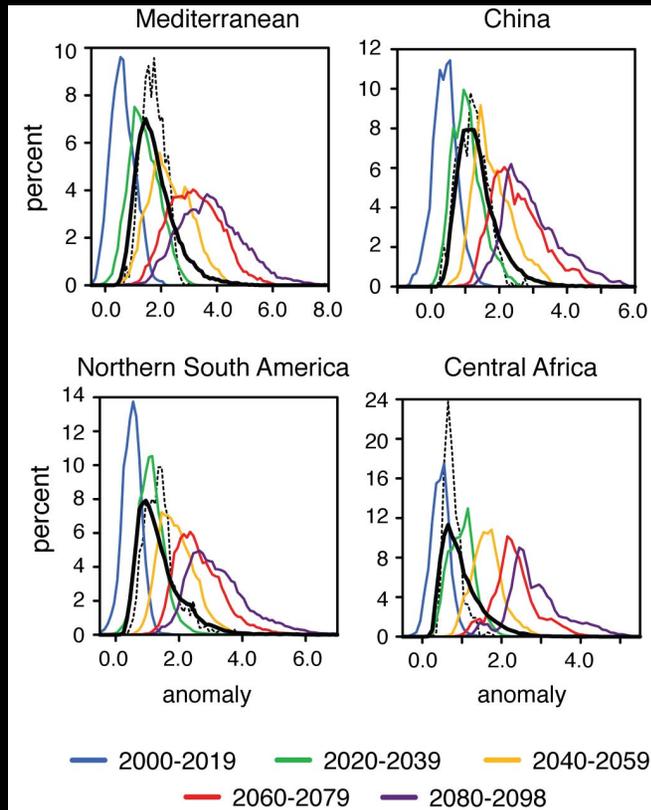


Figure 4. Probability density function (PDF) of JJA surface air temperature anomalies over land for four regions (Mediterranean, China, Northern South America and Central Africa). Each colored PDF represents anomalies for the CMIP3 ensemble for a 2-decade period of the 21st century (relative to the 1980-1999 period). The black lines show the PDF of the differences between the mean and maximum for the 1980-1999 period, with solid black lines showing the CMIP3 values and dashed lines showing the observed values. Regions are as in Figure 3.

Near-Term Global Warming

Many Models (A1B)

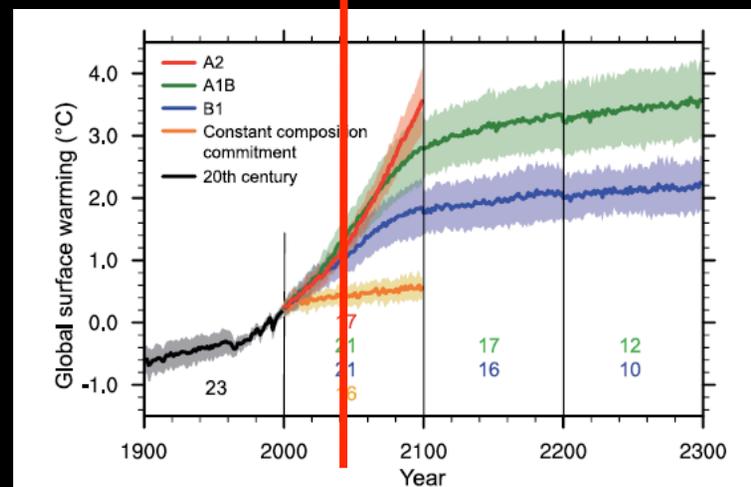
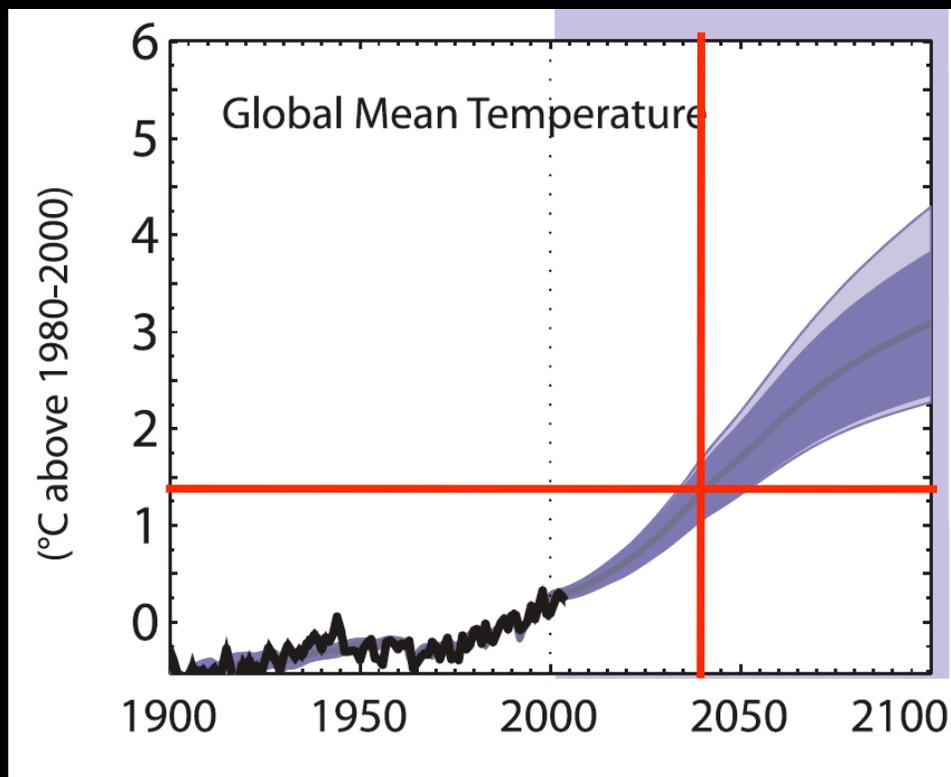


Figure 10.4. Multi-model means of surface warming (relative to 1980–1999) for the scenarios A2, A1B and B1, shown as continuations of the 20th-century simulation. Values beyond 2100 are for the stabilisation scenarios (see Section 10.7). Linear trends from the corresponding control runs have been removed from these time series. Lines show the multi-model means, shading denotes the ± 1 standard deviation range of individual model annual means. Discontinuities between different periods have no physical meaning and are caused by the fact that the number of models that have run a given scenario is different for each period and scenario, as indicated by the coloured numbers given for each period and scenario at the bottom of the panel. For the same reason, uncertainty across scenarios should not be interpreted from this figure (see Section 10.5.4.6 for uncertainty estimates).

1 °C additional global warming

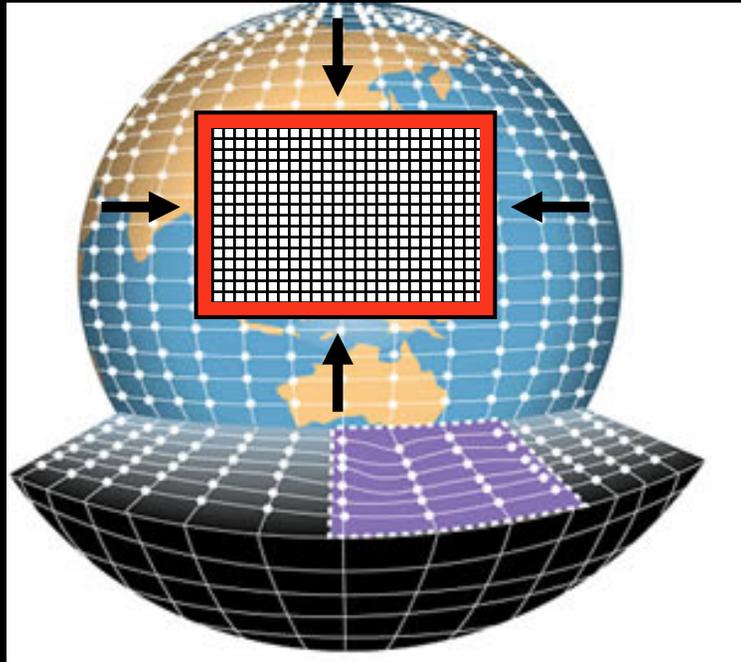


> 50% chance of every warm season being hotter than late-20th century maximum over much of the globe

Question 1: Why is 2°C the “Target”?

**What processes cause intensification
in different regions?**

Nested High-Resolution Climate Model

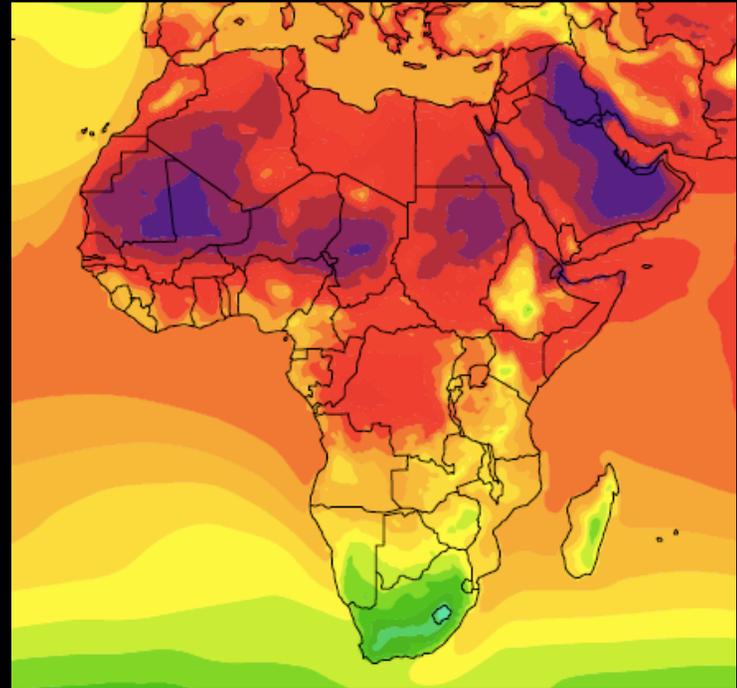
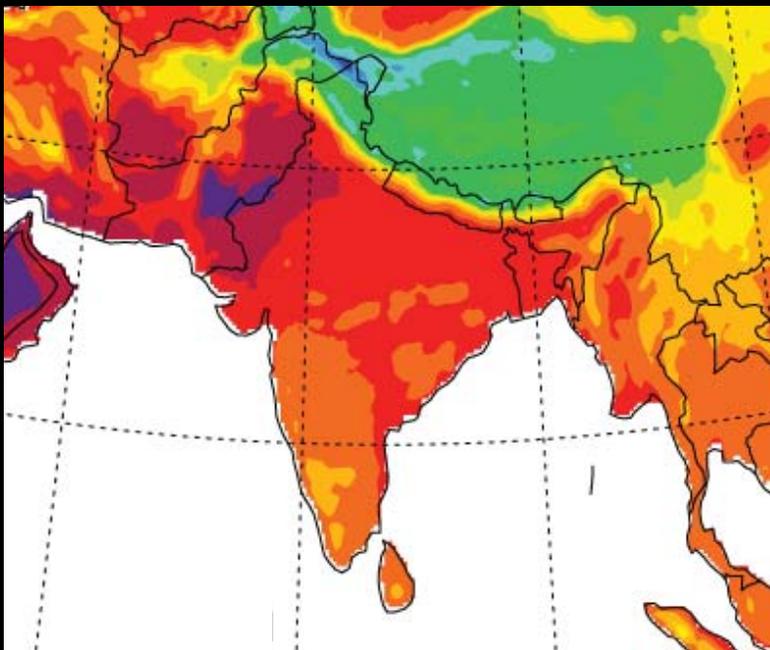
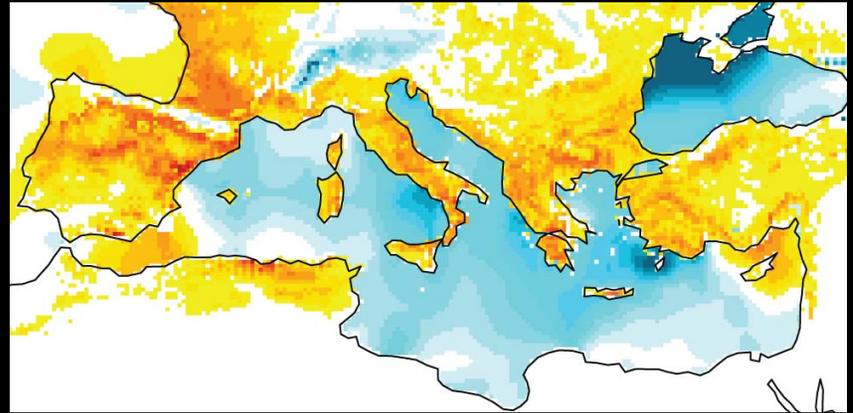
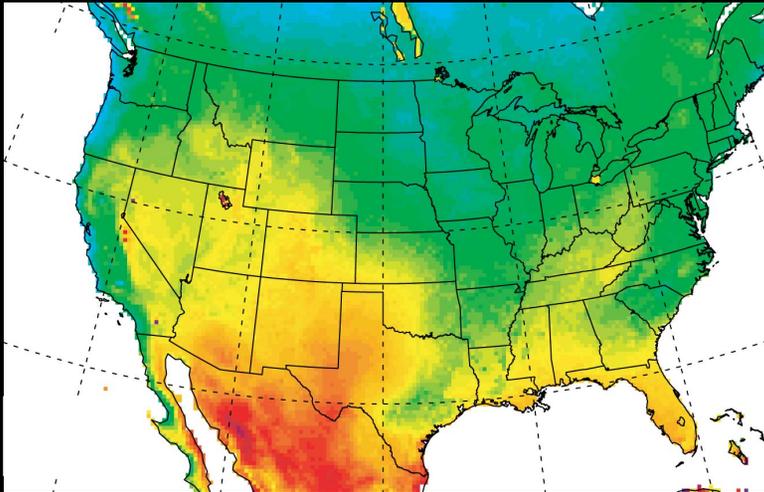


Global (“GCM”):
~125 km horizontal resolution

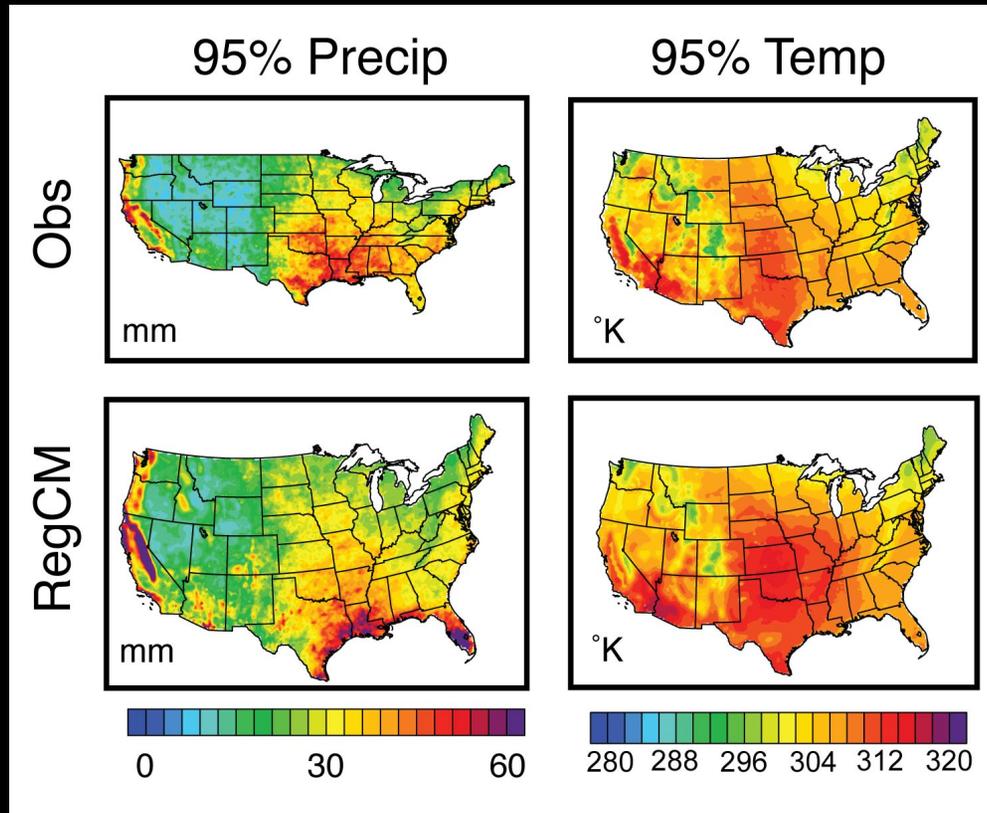
Nested (“RCM”):
~25 km horizontal resolution

Large-scale variables passed to nested model at lateral boundaries every 6 model hours

Multiple Regions

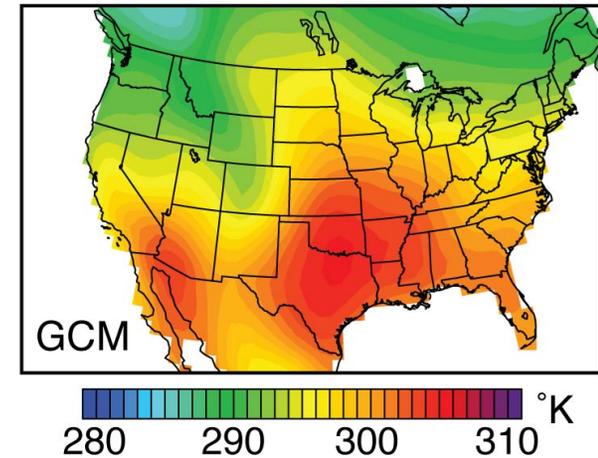
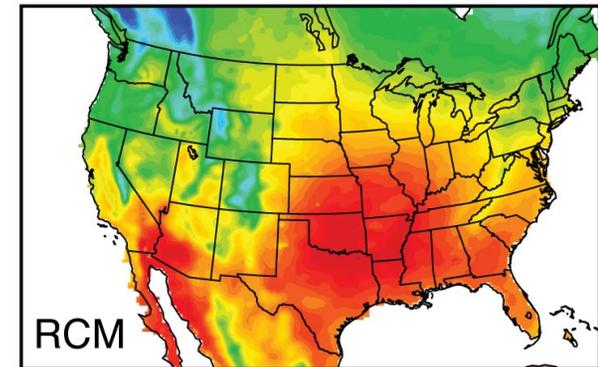
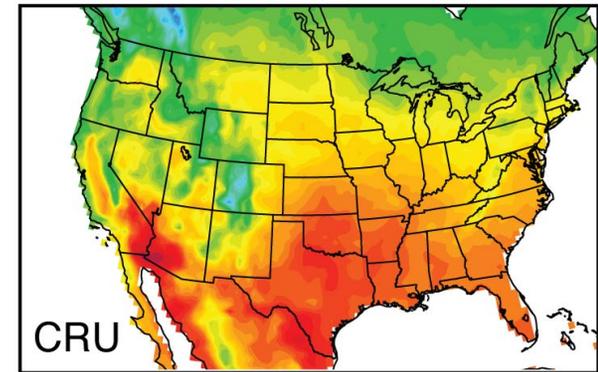


Extreme Events



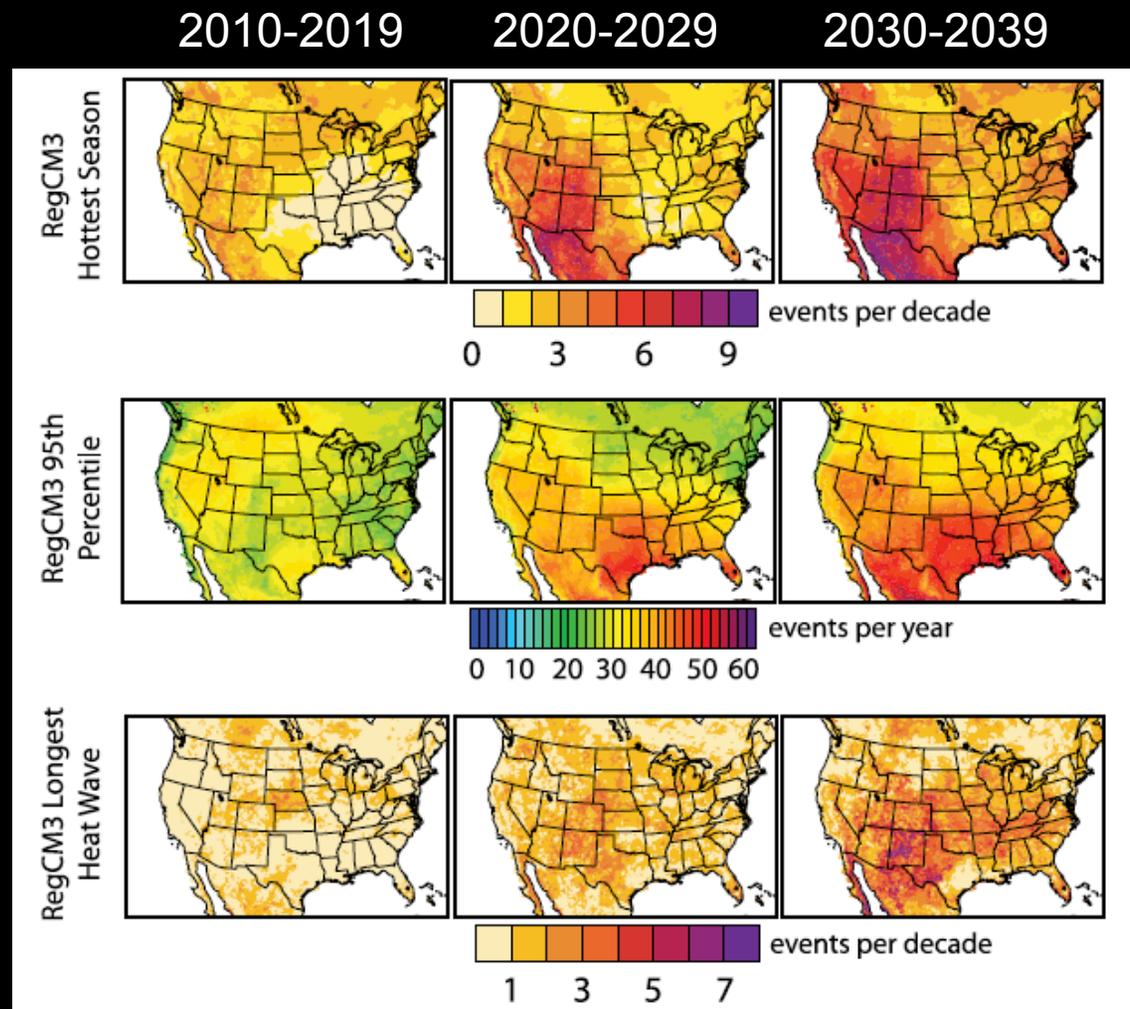
Walker and Diffenbaugh, 2009

1951-1999 Hottest Season



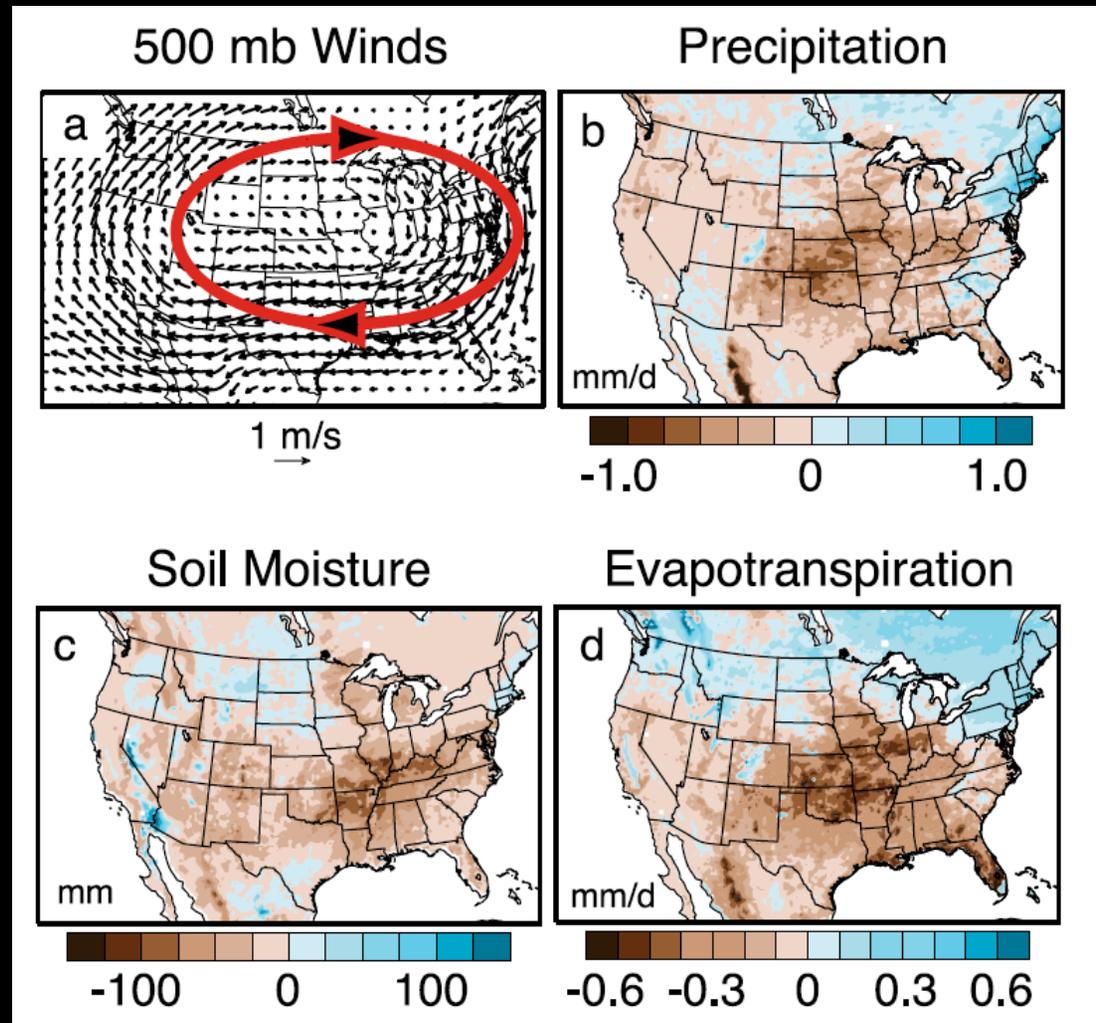
Diffenbaugh and Ashfaq, 2010

Intensification of Hot Extremes



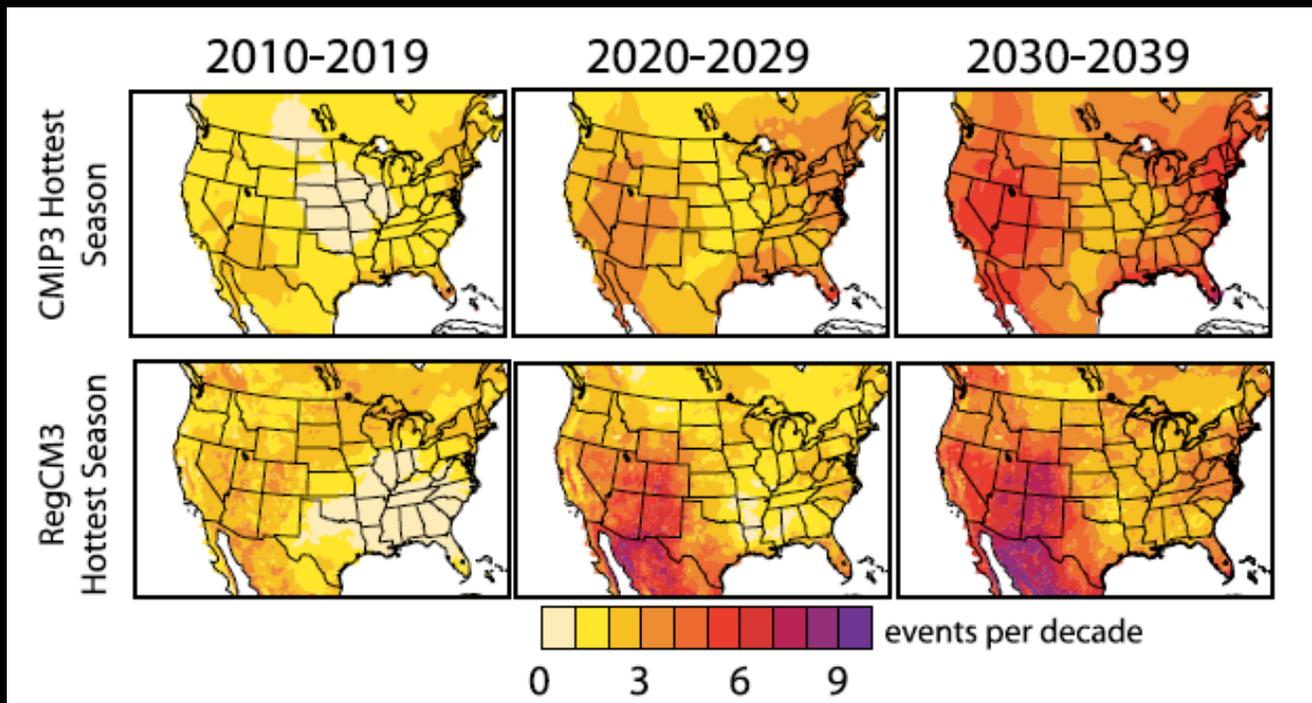
Summer Drying

2030-2039 minus 1980-1999



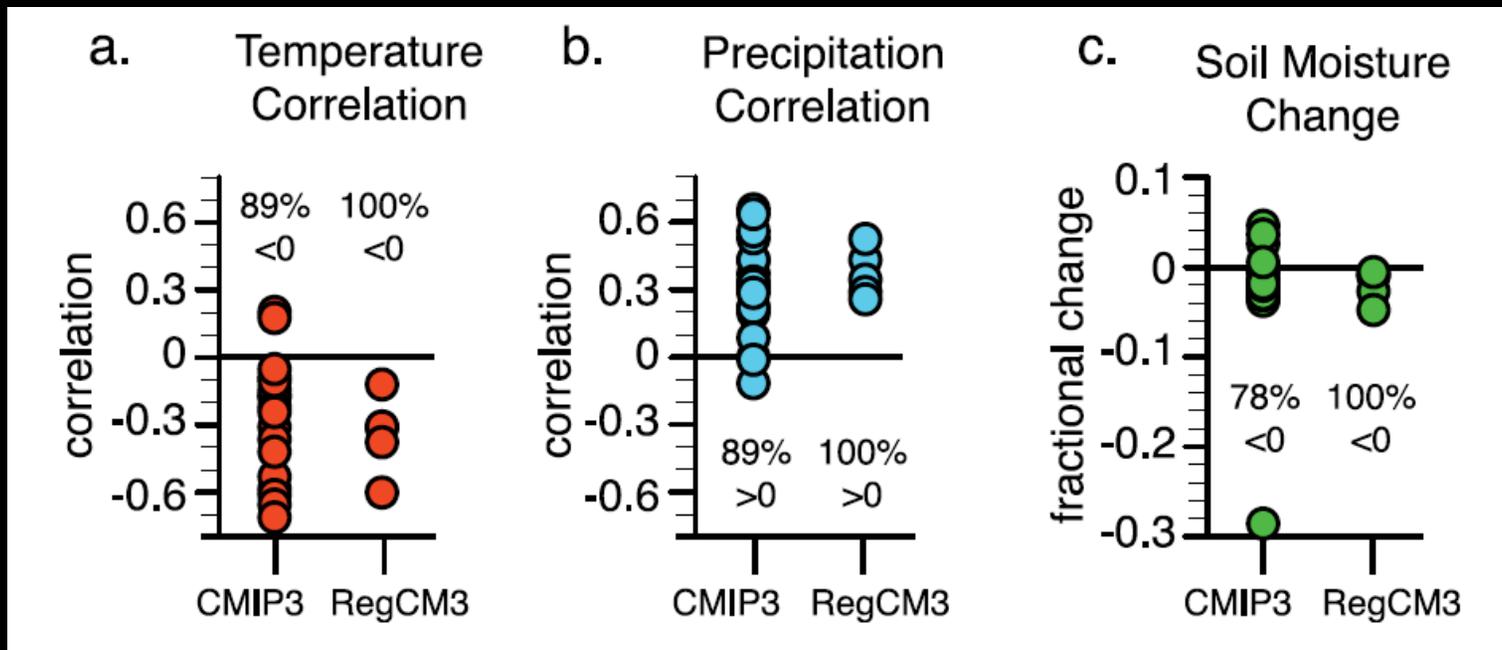
Diffenbaugh and Ashfaq, 2010

Comparison with “IPCC” GCMs



Summer Soil-Moisture/Temperature/ Precipitation Coupling

2030-2039 minus 1980-1999



What might be the impacts of such changes?

One Example: Temperature Suitability for Premium Winegrape Cultivation

Environ. Res. Lett. 6 (2011) 024024

N S Diffenbaugh *et al*

Table 2. Temperature screening criteria.

| Temperature variable | Minimum allowable | Maximum allowable |
|---|-------------------|-------------------|
| Growing season ^a growing degree days (GDD) ^b | 850 GDD | 2700 GDD |
| Growing season ^a diurnal temperature range (DTR) | — | 20 °C |
| Ripening season ^c diurnal temperature range (DTR) | — | 20 °C |
| Fall ^d , winter ^e and spring ^f severe cold days ^g | — | 14 days |
| Growing season ^a severe hot days ^h | — | 15–45 days |
| Growing season ^a mean temperature | 13 °C | 20–22 °C |

^a 1 April to 31 October. ^b 15 August to 15 October. ^c Base 10 °C. ^d 1 September to 30 November.

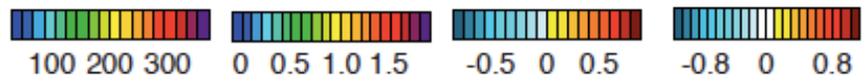
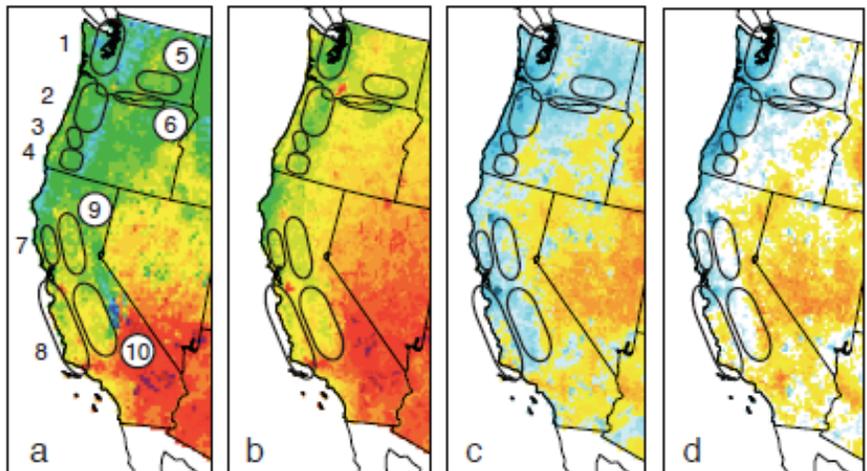
^e 1 December to 28 February. ^f 1 March to 31 May.

^g Total of fall days below -6.7 °C, Winter days below -12.2 °C, and Spring days below -6.7 °C.

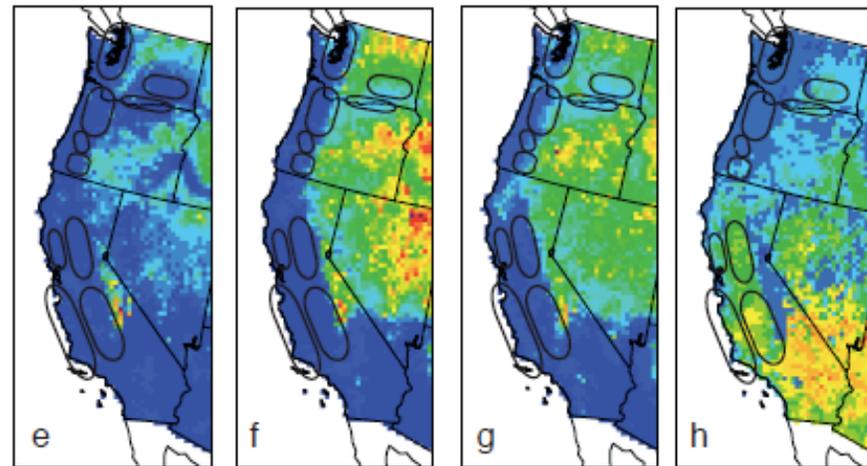
^h Days above 35 °C.

- sensitivity to heat and cold; mean and extremes
- variations within and at margins of suitability range

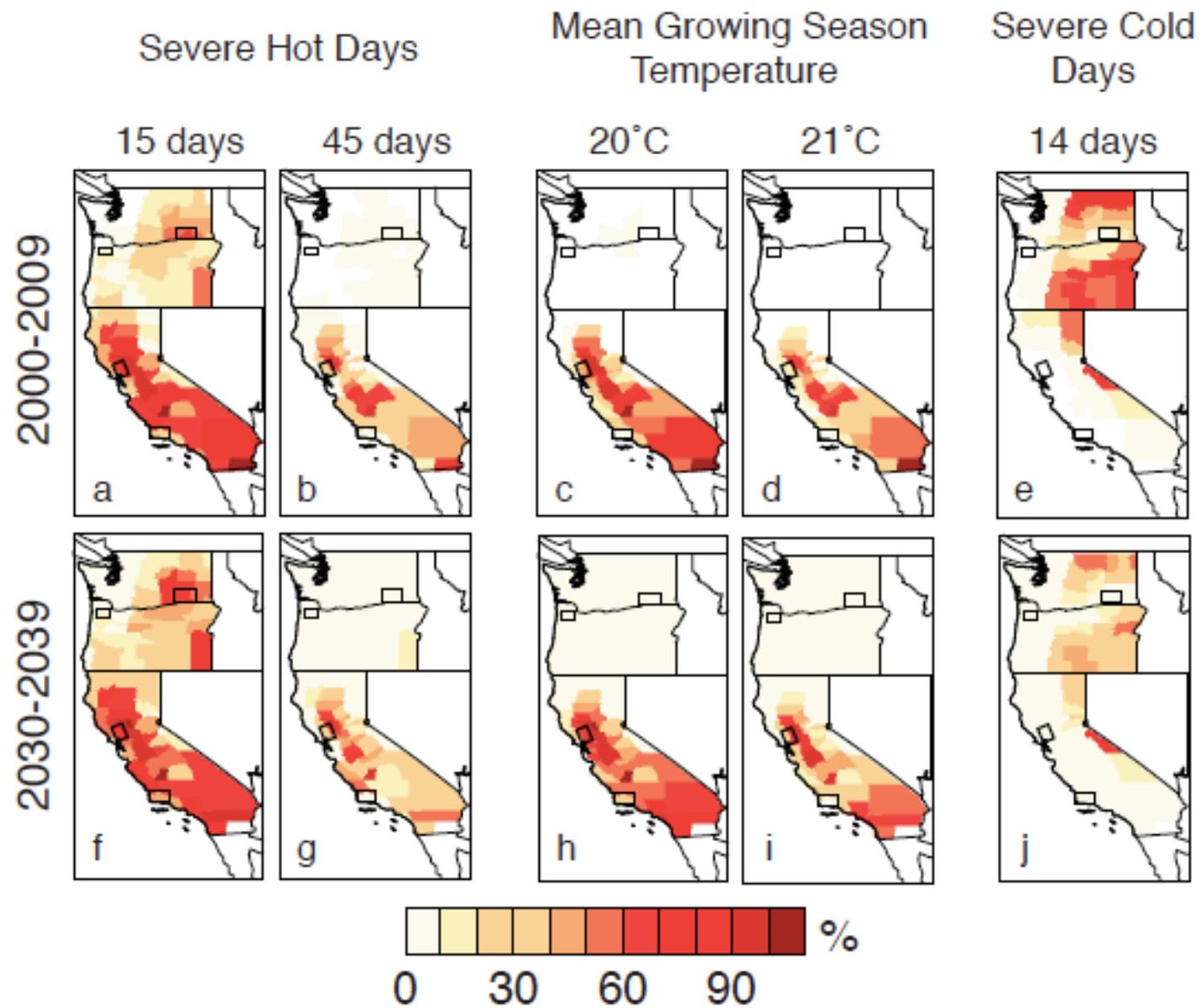
Winkler Summation (GDD) Growing Season Temperature (°C) Growing Season DTR (°C) Ripening Season DTR (°C)

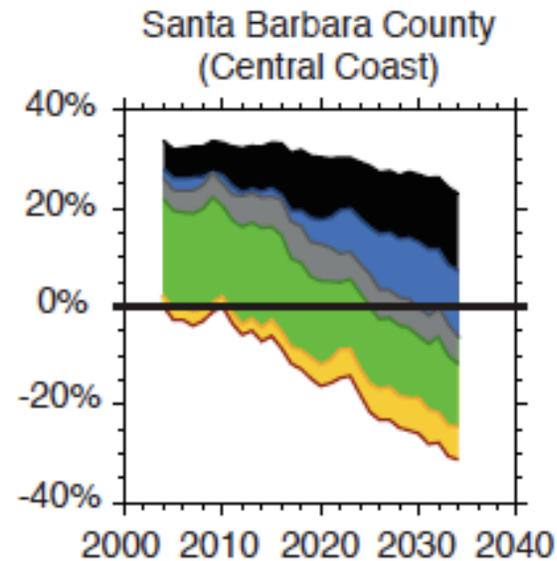
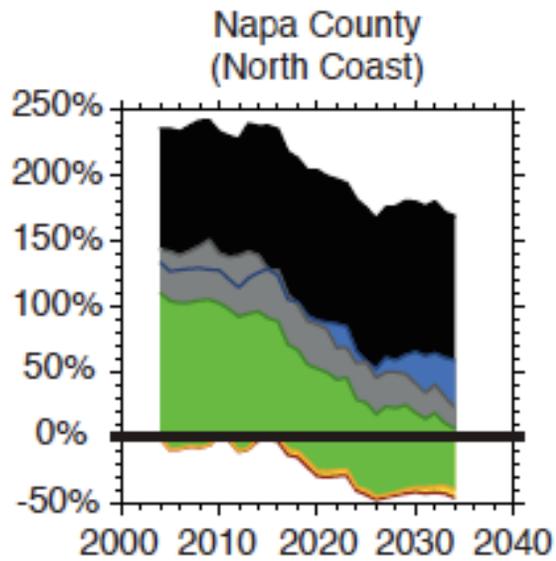
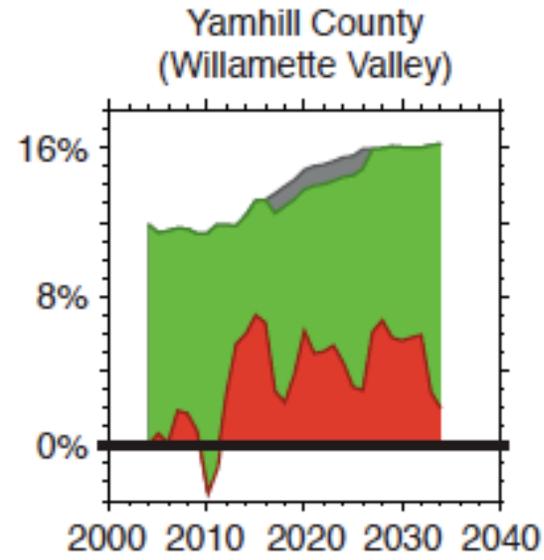
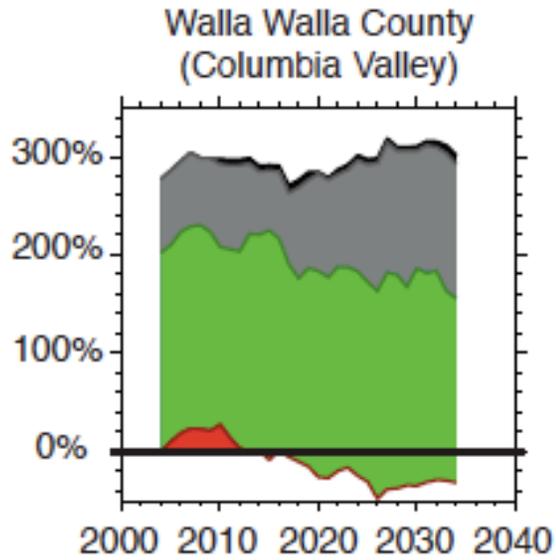


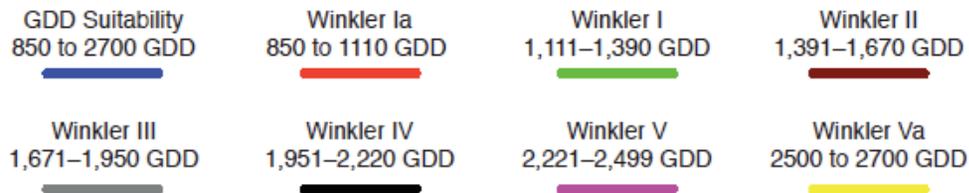
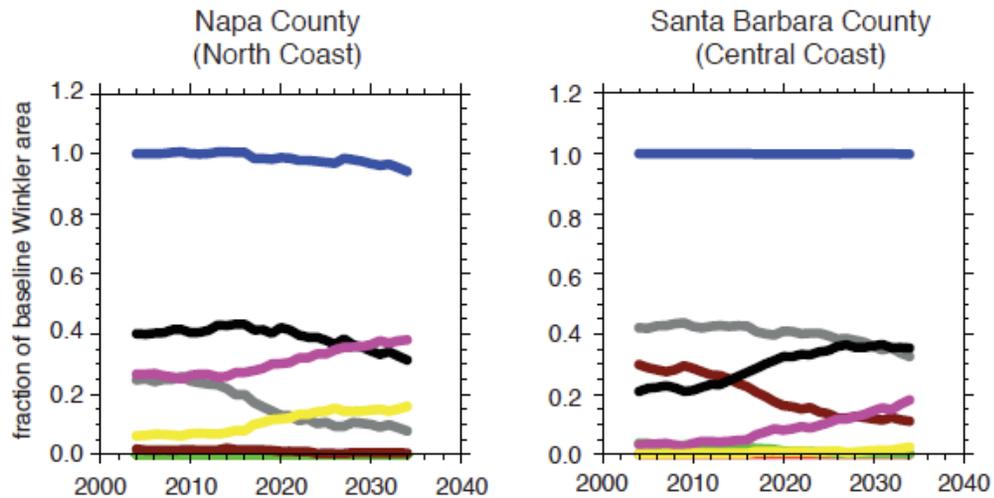
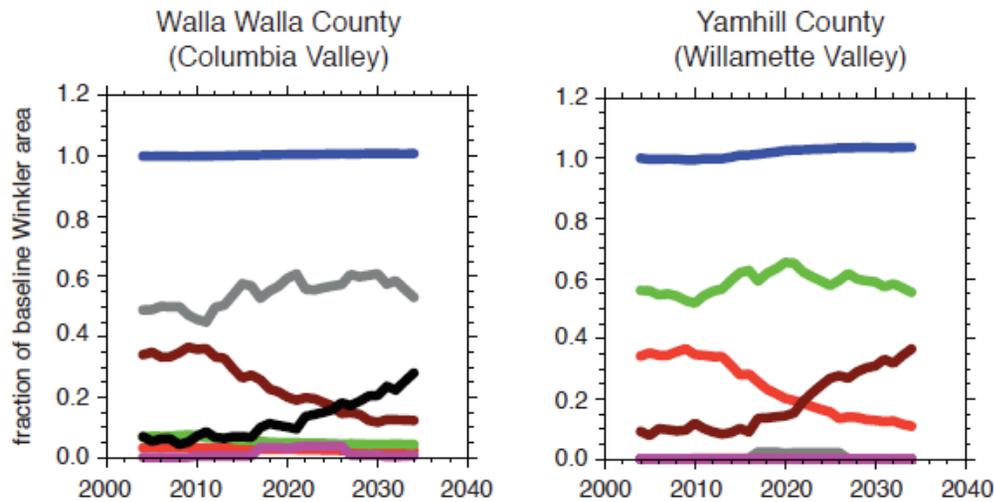
Winter Days Below -12.2 °C Spring Days Below -6.7 °C Autumn Days Below -6.7 °C Growing Season Days Above 35 °C



- heterogeneity in changes in temperature criteria, including extremes

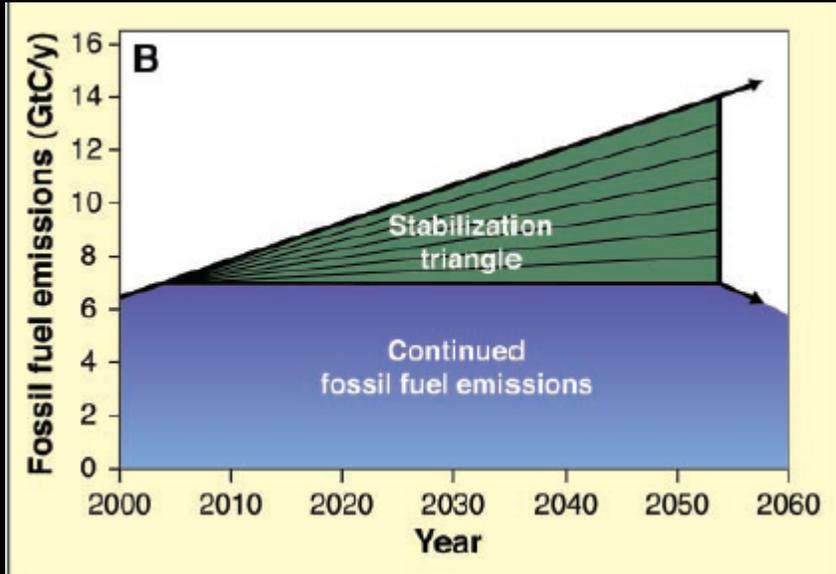




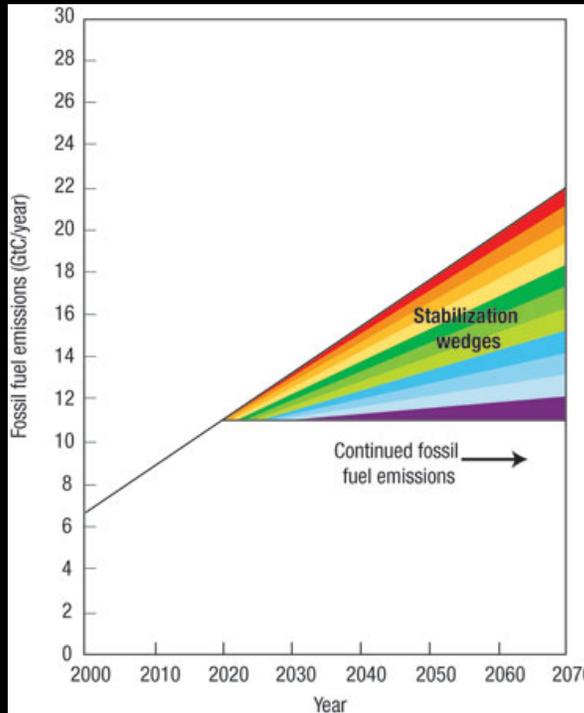


- potentially important changes in quality within suitability range

Pacala and Socolow, 2004



Wedges?



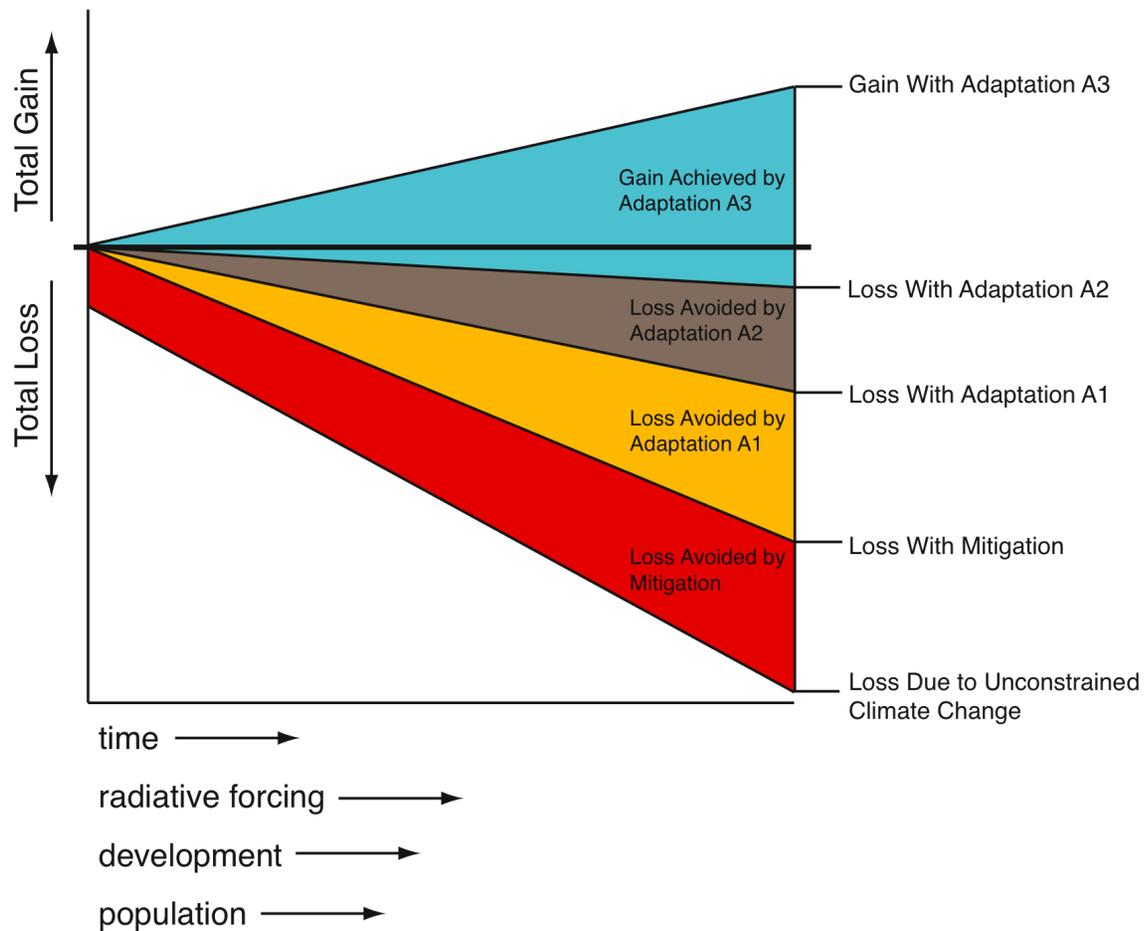
- Coal: 800 gigawatt-sized plants with all the carbon captured and permanently sequestered
- Nuclear: 700 new gigawatt-sized plants (plus replacement plants)
- Concentrated solar thermal electric: 1,600 gigawatts peak power
- Solar photovoltaics: 3,000 gigawatts peak power
- Efficient buildings: savings totalling 5 million gigawatt-hours
- Efficient industry: savings totalling 5 million gigawatt-hours, including co-generation and heat recovery
- Wind power: 1 million large wind turbines (2 megawatts peak power)
- Vehicle efficiency: all cars 60 miles per US gallon
- Wind for vehicles: 2,000 gigawatts wind, with most cars plug-in hybrid electric vehicles or pure electric vehicles
- Cellulosic biofuels: using up to one-sixth of the world's cropland
- Forestry: end all tropical deforestation

Romm, 2008

What is optimal combination of domains?

- Minimize total cost?
- Maximize total avoided damage?
- Maximize avoided damage/cost ratio?

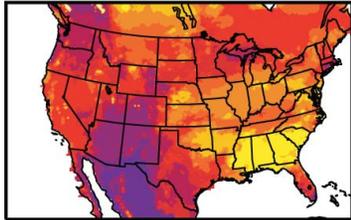
- What is total cost of each domain?
- What is avoided damage of each domain?



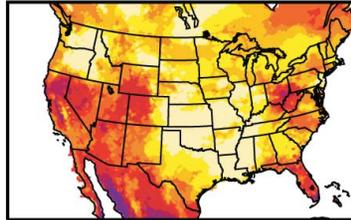
How Robust Is This
Emergence?

2030-2039

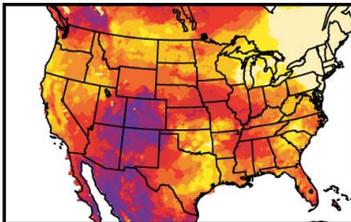
bES



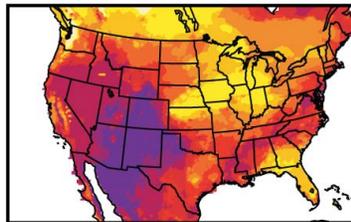
c



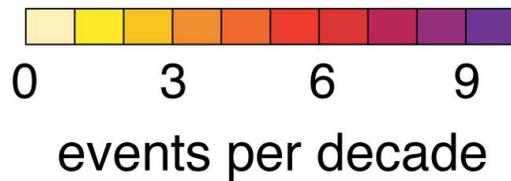
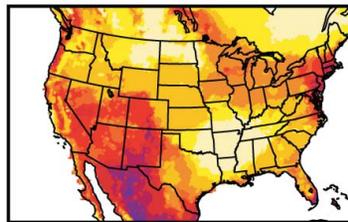
e



fES



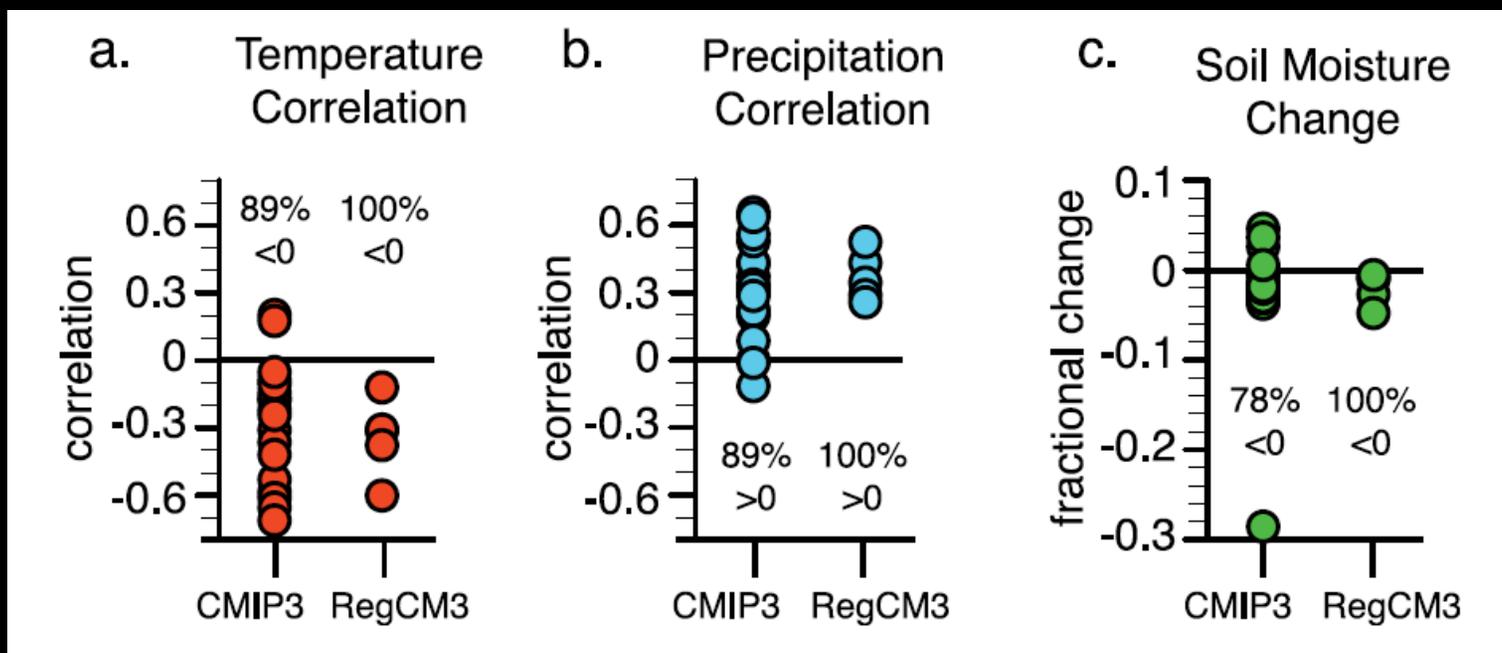
gES



- Variation within “physically uniform ensemble”

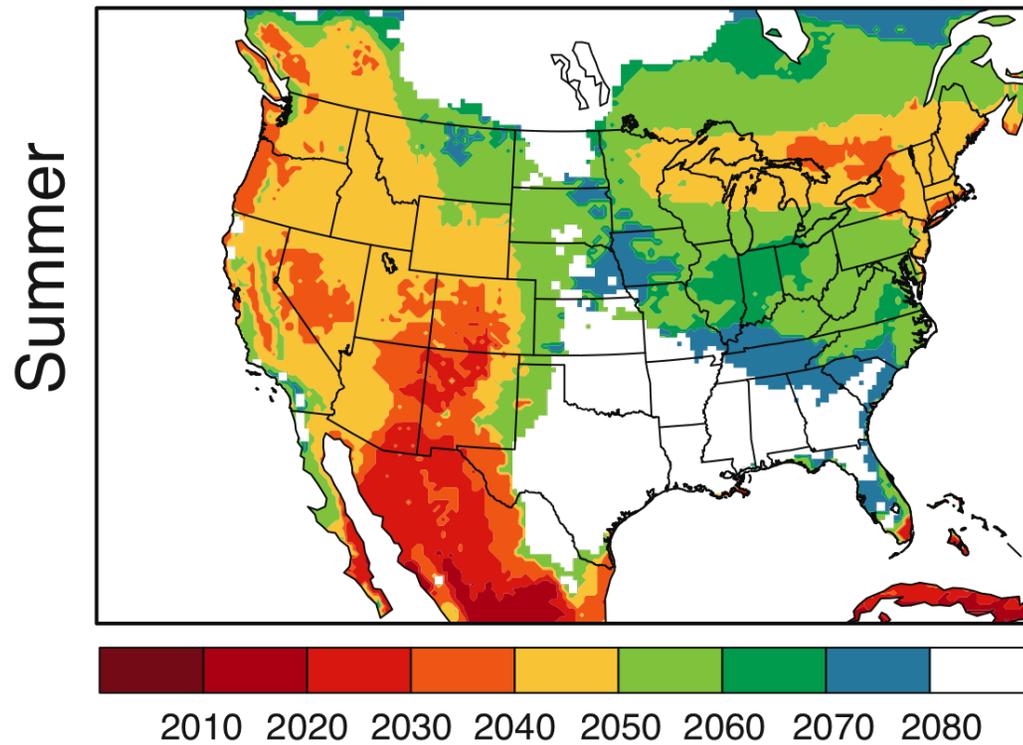
Summer Soil-Moisture/Temperature/ Precipitation Coupling

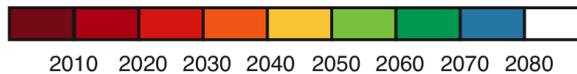
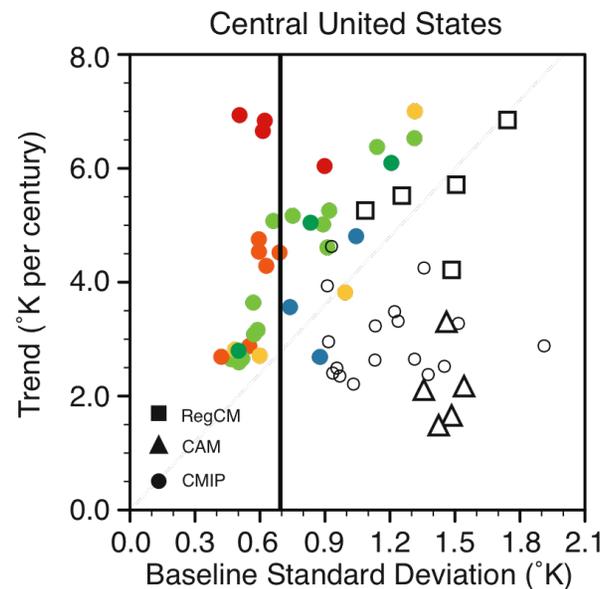
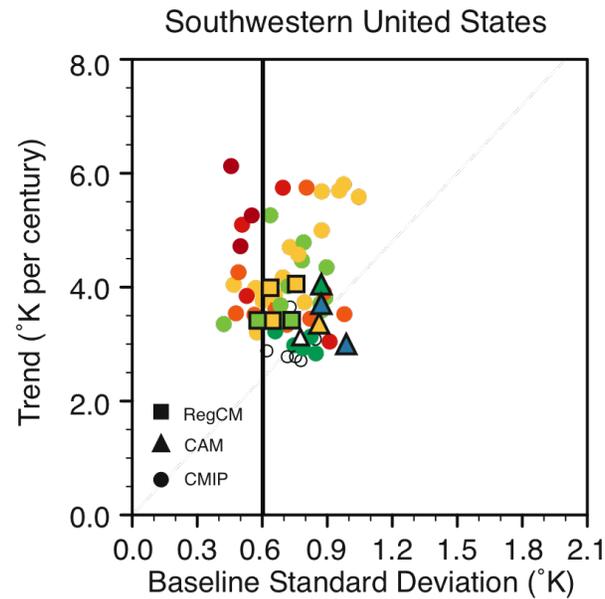
2030-2039 minus 1980-1999



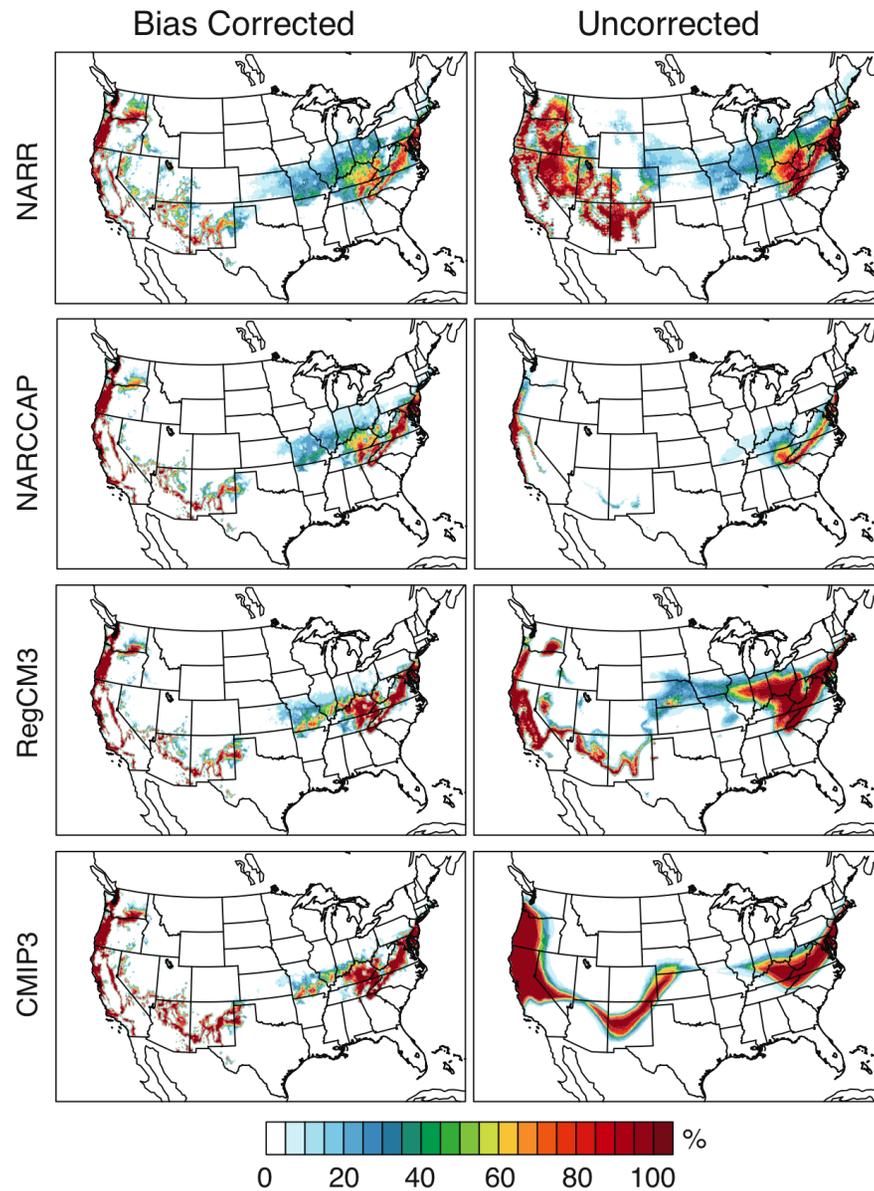
- Variation within “robust” physical response

Time of Emergence of Seasonal Warming





- Variation between models in both 21st century trend in mean and 20th century bias in variability

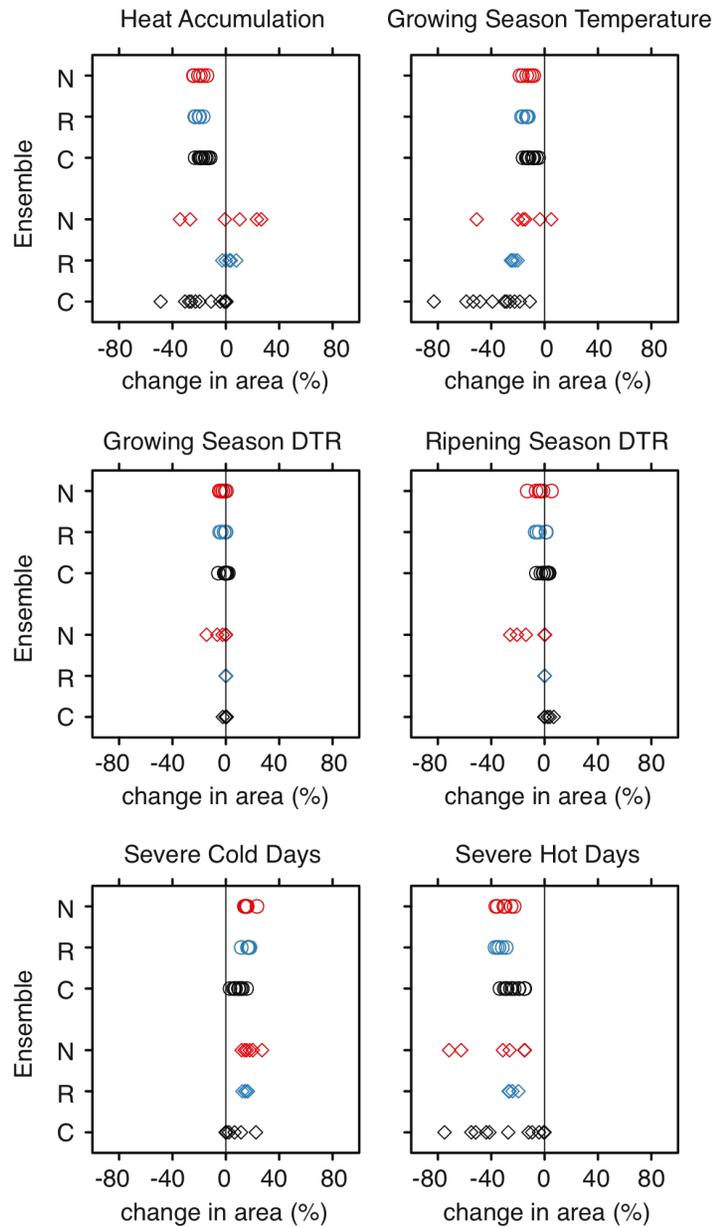


Premium Wine Suitability

- Bias correction reduces spread between model ensembles

(by construction for baseline)

California

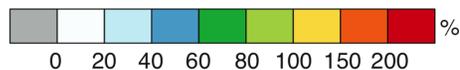
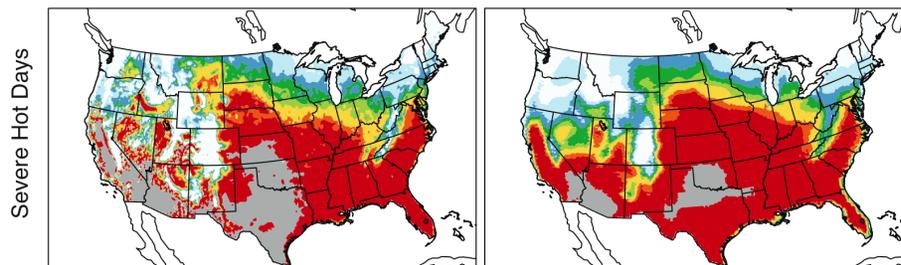


- Bias correction reduces spread within and between model ensembles (although less for single-model ensemble)

NARCCAP

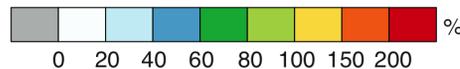
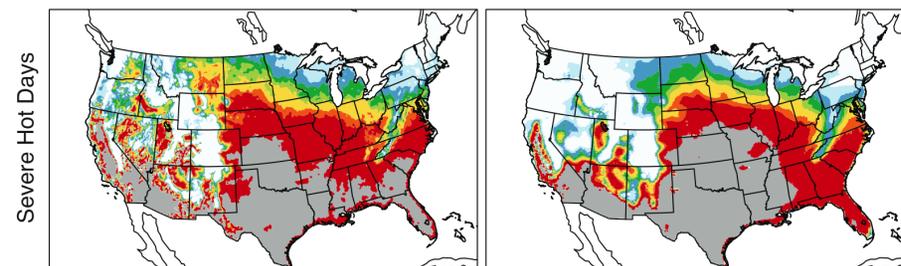
Bias Corrected

Uncorrected

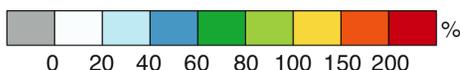
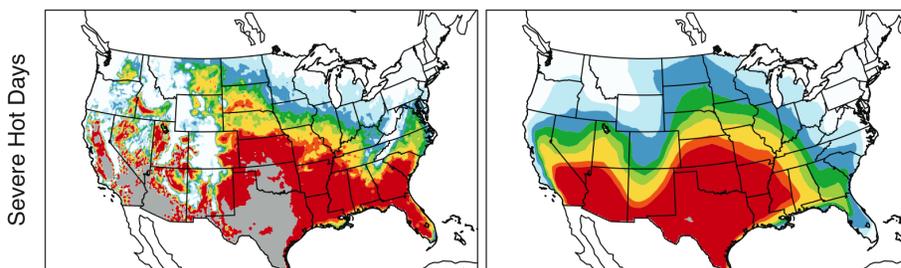


Change as
% of
“suitability
margin”

RegCM3

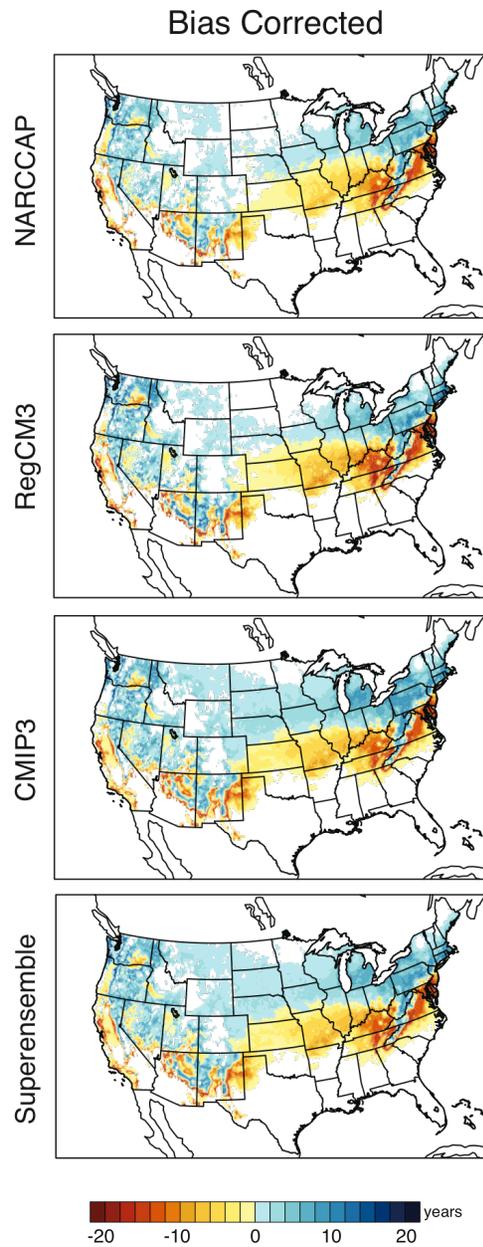


CMIP3

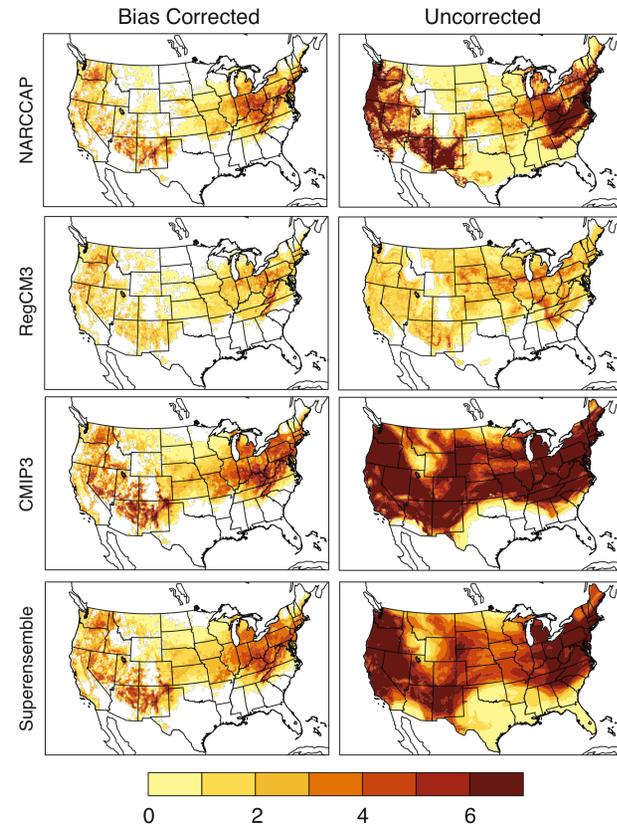


*Diffenbaugh and
Scherer, in prep*

Mid-21st Century Change



Intra-ensemble S.D.



Diffenbaugh and Scherer, in prep

What Does This All Mean For The E-L-W Nexus?

- Fine-scale climate processes can dictate
 - the magnitude and spatial variability of climate change,
 - where and at which levels of forcing critical thresholds are crossed,
 - and the magnitude and spatial variability of climate change impacts
- These processes can create substantial differences (in space and time) in impact and uncertainty
- Therefore, these fundamental climate dynamics must be integrated into policy frameworks in order to create effective interventions
- But... the near-term time scales and regional/local spatial scales that can be highly relevant for policy considerations can be at best difficult from a scientific perspective



Woods Institute for the Environment
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STANFORD UNIVERSITY
SCHOOL OF EARTH SCIENCES

Snowmelt-Dominated Runoff

